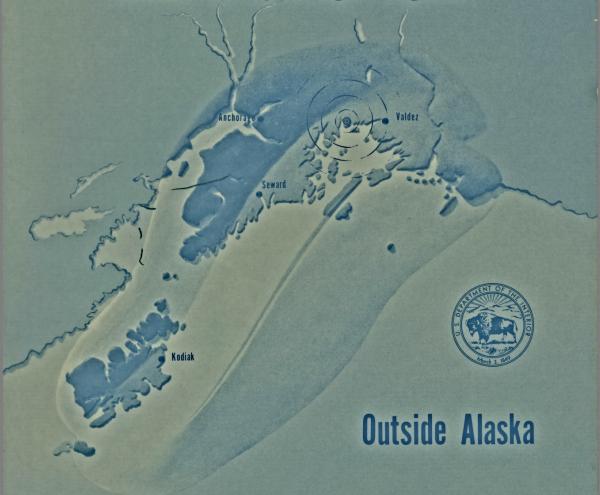
The Alaska Earthquake

March 27, 1964

Effects on Hydrologic Regimen



GEOLOGICAL SURVEY PROFESSIONAL PAPER 544-C

THE ALASKA EARTHQUAKE, MARCH 27, 1964: EFFECTS ON THE HYDROLOGIC REGIMEN

Hydrologic Effects of the Earthquake Of March 27, 1964 Outside Alaska

By ROBERT C. VORHIS

With Sections on

HYDROSEISMOGRAMS FROM THE NUNN-BUSH SHOE CO. WELL WISCONSIN

By ELMER E. REXIN and ROBERT C. VORHIS

and

ALASKA EARTHQUAKE EFFECTS ON GROUND WATER IN IOWA By R. W. COBLE

GEOLOGICAL SURVEY PROFESSIONAL PAPER 544-C

UNITED STATES DEPARTMENT OF THE INTERIOR

STEWART L. UDALL, Secretary

GEOLOGICAL SURVEY

William T. Pecora, Director



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THE ALASKA EARTHQUAKE SERIES

The U.S. Geological Survey is publishing the results of investigations of the Alaska earthquake of March 27, 1964, in a series of six Professional Papers. Professional Paper 544 describes the effects of the earthquake on the hydrologic regimen. Other Professional Papers describe the history of the field investigations and reconstruction effort; the effect of the earthquake on communities; the regional effects of the earthquake; and the effects on transportation, communications, and utilities.

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HYDROLOGIC EFFECTS OF THE EARTHQUAKE OF MARCH 27, 1964, OUTSIDE ALASKA

By Robert C. Vorhis

ABSTRACT

The Alaska earthquake of March 27, 1964, had widespread hydrologic effects throughout practically all of the United States. More than 1,450 water-level recorders, scattered throughout all the 50 States except Connecticut, Delaware, and Rhode Island, registered the earthquake. Half of the water-level records were obtained from ground-water observation wells and half at surfacewater gaging stations. The earthquake is also known to have registered on water-level recorders on wells in Canada, England, Denmark, Belgium, Egypt, Israel, Libya, Philippine Islands, South-West Africa, South Africa, and Northern Territory of Australia.

The Alaska earthquake is the first for which widespread surface-water effects are known. The effects were recorded at stations on flowing streams, rivers, reservoirs, lakes, and ponds. The 755 surface-water stations recording effects are spread through 38 States, but are most numerous in the south-central and southeastern States, especially in Florida and Louisiana. Most of the fluctuations recorded can be referred to more precisely as seismic seiches; however, a few stations recorded the quake as a minor change in stage. The largest recorded seiche outside Alaska was 1.83 feet on a reservoir in Michigan. The

next largest was 1.45 feet on Lake Ouachita in Arkansas.

The largest fluctuation in a well was 23 feet registered by a pressure recorder near Belle Fourche, S. Dak. Fluctuations of more than 10 feet were reported from wells in Alabama, Florida, Georgia, Illinois, Missouri, and Pennsylvania. A 3.40-foot fluctuation was recorded in a well in Puerto Rico.

The Alaska earthquake was registered by about seven times as many water-level recorders as recorded the Hebgen Lake, Mont., earthquake of August 19, 1959.

INTRODUCTION

The hydrologic response to the Alaska earthquake of March 27, 1964, was the most widespread of all previously registered seismic events. Some 716 wells in the United States recorded water-level fluctuations caused by the quake. Outside the United States, wells in Canada, England, Belgium, Denmark, Libya, Israel, South-West Africa, South Africa, and Australia recorded the earthquake. This worldwide response results

from the great magnitude of the quake—the largest to occur in North America in this century. Of the previous large earthquakes recorded widely in the United States, the Assam, India, earthquake of August 15, 1950, affected at least 161 wells, and the Hebgen Lake, Mont., earthquake of August 17, 1959, was recorded in 185 wells.

Another important response was registered as water-level fluctua-

tions on streams, reservoirs, ponds, and lakes. At most of the gaging stations the charts show upward and downward motions that were about equal and that generally recovered to a normal level within a few minutes. On some lakes and reservoirs the fluctuations continued for an hour or more.

A third effect caused by the earthquake was roiling or muddying of well and spring water. This phenomenon, when reported by well users, generally occurred in wells that required long-continued pumping to clear the water at the time the well was drilled. The roiling is presumably limited to wells and springs tapping aquifers that contain considerable colloidal material.

The purpose of this report is to assemble the hydrologic effects of the Alaska earthquake that were recorded outside Alaska. The nature and geographic distribution of the hydrologic effects are described, the seismic fluctuations in water wells in the United States are tabulated, and data on seismic seiches in North America and well fluctuations outside the United States are summarized. Thus, the report is a compilation of both published and unpublished data on known hydrologic effects throughout the world. Furthermore, a background of previous work, a discussion of several water-level recording instruments, and observations of fluctuations during other earthquakes are presented in order to provide a suitable basis for future interpretive studies. It is hoped that this framework will encourage further studies so that the discrepancies that exist between earthquake theory and observed effects can be narrowed and ultimately bridged.

DEFINITION OF TERMS

The term "hydroseism" is here introduced as convenient to include all seismically induced water-level fluctuations other than tsunamis. Although this type of fluctuation has been described in many previous papers, no one term nor one phrase has been used consistently. Terms and phrases which have been used to describe hydroseisms in wells include:

1. Pressure fluctuations produced by seismic waves.

- 2. Seismically induced fluctuations of water level.
- 3. Water-level fluctuations.
- 4. Earthquake-induced fluctuations.
- 5. Fluctuations in well water levels (the title under which hydroseismic data have been published annually in the U.S. Coast and Geod. Survey series "United States Earthquakes").

The following terms have been used to describe hydroseisms in surface-water bodies:

- 1. "Disturbances [as recorded] at stream-gaging stations" (U.S. Coast and Geod. Survey, 1946, p. 26).
- 2. "Range of stage recorded in stilling wells * * * as the result of earth tremors" (Stermitz, 1964, p. 144).
- 3. Seismic seiches.

The term "hydroseism" is derived from the Greek words νδωρ meaning water and σεισμος meaning earthquake. As defined and used in this report, hydroseism applies to seismically induced fluctuations in wells, streams, lakes, ponds, and reservoirs. As such, it is identical in meaning with any and all of the expressions listed above.

"Hydroseismic data" includes both the charts that record hydroseisms and the information taken from the charts.

"Hydroseismogram" is a hydroseism recorded at an expanded time scale.

"Seiche" is a term first used in Switzerland by Forel (1895) to apply to standing waves set up on the surface of Lake Geneva by wind and by changes in barometric pressure. Richter (1958, p. 109) points out that seiches may occur not only in closed bodies of water but also in partially closed bodies such as harbors or channels and as lateral oscillations in

rivers, canals, or ditches. It is only necessary that the geometry of the water boundary define a natural period of oscillation. Where there are currents, part of a seiche may be transformed into a progressing wave.

To restrict the term "seiches" to those events caused by earthquakes, Kvale (1955) qualified the phenomenon as seismic seiches. Following his usage, in this paper "seismic seiches" refer to symmetrical fluctuations (that is, those fluctuations where the water-level rise is exactly equal to the waterlevel decline) typical of standing waves set up on rivers, reservoirs, ponds, and lakes at a time corresponding with the passage of seismic waves from the Alaska earthquake. Where the record does not correspond to what would be expected from a standing wave, the more general term "hydroseism" is used.

ACKNOWLEDGMENTS

The material on which this report is based has come from many sources and has been the work of hundreds of persons engaged in various aspects of water-resource investigations. Most of these people are known to the writer only by their affiliation with the organization which transmitted the data.

The following organizations outside the United States have furnished data on hydrologic effects of the Alaska earthquake in their countries, and their help is gratefully acknowledged:

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Canada, Ontario Water Resources Commission

Canada, Saskatchewan Department of Agriculture

United Kingdom of Libya, Ministry of Agriculture

State of Israel, Hydrological Service of Ministry of Agriculture

South-West Africa, Geological Survey

Republic of South Africa, Geological Survey

Commonwealth of Australia, Northern Territory Administration

New South Wales, Snowy Mountains Hydro-Electric Authority

Victoria, State Electricity Commission

Belgian Ministry of Public Works, Hydraulic Research Laboratory

United Kingdom, Geological Survey and Museum, Water Department

Denmark Geological Survey Republic of the Philippines, Bureau of Public Works

United Arab Republic, U.S. AID Mission

Most of the records of hydroseisms in the United States were compiled and sent to the author by members of the Water Resources Division of the U.S. Geological Survey. The following organizations in the United States have also made their records and data available for use, and the author is grateful to them for this help:

> Orange County [Calif.] Water District

> Los Angeles County [Calif.] Flood Control District

> Idaho Bureau of Mines and Geology

> Illinois State Water Survey Gas storage companies in Illinois, Iowa, and Ohio

> Missouri Geological Survey and Water Resources

Division of Water, Ohio Department of Natural Resources

Tennessee Department of Conservation, Division of Water Resources

Virginia Geological Survey

The author is also grateful to the many organizations which examined their recorder charts for the period of the Alaska earthquake and verified that no fluctuations caused by the earthquake had been recorded. The countries and names of the organizations in which recorded charts were found to show no apparent reaction to the quake are the following:

Afghanistan:

Water and Soil Survey Authority

U.S. AID Mission

Australia:

Sydney Metropolitan Water, Sewerage, and Drainage Board

Irrigation and Water Supply Commission

South Australia Engineering and Water Supply Department

South Australia Director of Mines

Victoria Rivers and Water Supply Commission

Austria:

Hydrografisches Zentral-

British Guiana: Geological Survey Department

Cyprus:

Ministry of Commerce and Industry

Geological Survey Department

Denmark: Meteorological Institute

Ethiopia: Water Resources Department

Ghana: National Construction Corp.

Greece: Institute for Geology and Subsurface Research

Hungary: Research Institute for Water Resources

Jamaica: Geological Survey Department

Indonesia: Geological Survey

Kenya: Water Development Department

Mozambique: Service for Geology and Mines

Nepal:

Hydrological Survey Department

U.S. AID Mission

Netherlands: Archives of Ground-Water Levels

New Guinea: Australia Department of Public Works New Zealand:

Ministry of Works Geological Survey

Nigeria:

Geological Survey U.S. AID Mission

Norway: Water Resources and Electricity Board

Papua: Australia Department of Public Works

Portugal: Geological Survey Rhodesia:

> Ministry of Water Development, Hydrological Branch

> Ministry of Mines and Lands, Geological Survey Office

Saudi Arabia: Ministry of Petroleum and Mineral Resources

Spain: Institute of Geology and Mines

Sudan: Geological Survey Switzerland: Federal Office

for Water Economy Syria: Ministry of Industry Taiwan: Geological Survey

Tasmania:

Rivers and Water Supply Commission

Hydro Electric Commission

Turkey: State Hydraulic \mathbf{Works}

Uganda: Water Development Department

Union of Soviet Socialist Republics: Hydrometeorological Service

Zambia: Ministry of Lands and Natural Resources

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PREVIOUS INVESTIGATIONS

Hydrologic effects of earthquakes have been previously compiled from observation wells and at surface-water gaging stations. In some studies, recorders with expanded time scales have been operated to ascertain the types of seismic waves that produce hydrologic effects. The Richter scale of earthquake magnitudes has been applied to hydroseisms, and theoretical studies have been developed to account for hydrologic responses to seismic waves. No single study, however, has considered both the effects recorded by observation wells and those recorded at surface-water gaging stations.

HYDROSEISMS IN WELLS

Data on hydroseisms in wells are published annually by the U.S. Coast and Geodetic Survey (1945, 1946, 1951–65). The years for which these data are now available include 1943, 1944, and 1949 through 1963. Only one publication (da Costa, 1964) lists hy-

droseisms in wells throughout the United States from a single major earthquake, the Hebgen Lake, Mont., earthquake of August 17, 1959.

Other published reports describe hydroseisms that have been recorded in a single well. Parker and Stringfield (1950, p. 456-458) list earthquakes and their hydroseisms as recorded in a limestone well at Miami, Fla. Eaton and Takasaki (1959) made a similar study of a well in basalt at Honolulu, Hawaii, and were the first to attempt to correlate earthquake magnitudes with hydroseismic data. Vorhis (1964a), using the same approach, made a study of hydroseisms recorded in a well penetrating crystalline rocks in Dawson County, in the Piedmont of northern Georgia.

Hydroseisms from one or a few selected earthquakes as recorded in wells scattered through a rather restricted geographic area have been described in many publications: Austin (1960) for wells in New Jersey; Davis, Worts, and Wilson (1955) for hydroseisms in California wells caused by the Kern County earthquake of 1952; Hopkins and Simpson (1960) for the effects of the Hebgen Lake earthquake in mine-water pools in Pennsylvania; LaMoreaux (1953) for the effects in Alabama wells of the Kamchatka earthquake of November 4, 1952; LaRocque (1941) for hydroseisms in California wells caused by five different quakes from 1933 to 1940; Leggette and Taylor (1935) for hydroseisms in Utah wells; Piper (1933) for hydroseisms recorded in the Mokel, umne area of California from the December 20, 1932, earthquake at Lodi, Calif.; Piper, Thomas, and Robinson (1939) for later hydroseisms recorded in the Mokelumne area of California; Stearns (1928) and Stearns, Robinson, and Taylor (1930, p. 145–151) for hydroseisms in California wells; Stewart (1958) for hydroseisms in Georgia wells; Thomas (1940) for hydroseisms recorded in California by earthquakes of November 10, 1938 (in Alaska), and January 24, 1939 (in Chile); and Vorhis (1955) for hydroseisms in a Wisconsin well.

The theoretical effect of earthquake waves on a water well was originally studied from a seismological point of view by Blanchard and Byerly (1935). A reexamination of this theory was made by Rexin, Oliver, and Prentiss (1962) who related the magnification in the well to the period of the seismic surface waves causing the fluctuation. The Alaska earthquake reawakened interest in the subject. Subsequent reports hold much promise for providing a theoretical understanding of hydroseismic data. Papers containing a development of theory with emphasis on the hydrologic controls have been prepared by Cooper and others (1965) and Bredehoeft and others (1965). The Richter scale of earthquake magnitudes has been applied to hydroseisms (Eaton and Takasaki, 1959; Vorhis, 1964a, 1965a), but the many variables relating to a well, to earthquake waves, and to their travel paths prevent a rigorous application of this method, at least for the present.

Hydrologic effects of the Alaska earthquake have been discussed or mentioned in many publications. Fluctuations in Canadian wells were described for Alberta by Gabert (1965) and for Ontario and Saskatchewan by Scott and Render (1965). Widespread hydrologic effects of the quake as recorded in the United States were published by Peterson (1964), Waller, Thomas, and Vorhis (1965), and Vorhis (1965b). Local hydrologic effects within the United States were

discussed by Coble (1967), Fellows (1965), Fuller (1964), Wilson (1964), Hassler (1965), Miller and Reddell (1964), The Cross Section (1964), Mills (1964), Rexin (1963, 1964a, b), Cooper and others (1965), and Bredehoeft and others (1965).

HYDROSEISMS ON SURFACE-WATER BODIES

Earthquake effects on river gages were first mentioned by Piper (1933, p. 475, fig. 2). Two of six gages on the Mokelumne River in California showed pronounced dots on the traces of the water surface that were caused by the December 20, 1932, earthquake at Lodi, Calif. Two other gages on a nearby diversion canal showed double amplitudes of 0.08 and 0.04 foot for the same quake. Although they were not designated as such by Piper, they can properly be called seismic seiches.

The U.S. Coast and Geodetic Survey (1946, p. 26) lists 18 "Disturbances at stream-gaging stations in New York on September 5, 1944, from an earthquake in the St. Lawrence Valley." Most of these may be seismic seiches but, not knowing whether they are standing waves, it seems safer to refer to them as hydroseisms.

For the Hebgen Lake earthquake of August 17, 1959, Stermitz (1964, p. 144, table 10) lists 54 records for "Range of stage recorded in stilling wells * * * as the result of earth tremors."

The hydroseisms mentioned above were all recorded within a few hundred miles of the epicenter of the quakes. In Stermitz's list, the most distant event recorded was 340 miles northwest of the epicenter. The events discussed below are unusual in that they were recorded thousands of miles distant from the epicenter.

Kvale (1955) discusses seismic seiches in 29 fiords and lakes in Norway caused by the Assam earthquake of August 15, 1950, but does not mention that any seiches were recorded in rivers.

Seismic seiches from the Alaska quake, recorded on the west coast of Canada in lakes, rivers, and tidal waters, are mentioned by Wigen and White (1964a, b). Some seismic seiches on rivers and lakes in central Canada are described by Strilaeff (1964).

Donn (1964) mentions reports of waves as much as 6 feet high on the gulf coast caused by the Alaska earthquake. He suggests that these waves and the fluctuations on the Freeport, Tex., tide gage are probably seismic seiches generated in resonance with the seismic waves.

McGarr (1965) formulated a theory to explain the generation, occurrence, and damping of seismic seiches. The theory was then applied to a marigram at Freeport, Tex., that showed the Alaska earthquake.

Thus, seismic seiches represent a type of hydroseism that has received little attention previously, especially such seiches recorded in rivers.

HYDROSEISMIC DATA

TYPES OF WATER-LEVEL RECORDERS AND CHARTS

The water-level recorders currently in operation on observation wells throughout the United States meet a wide variety of needs; they are adapted to fit all types of well construction, are selected to fit economical servicing schedules, and must respond to many different ranges in water-level fluctuation. As a result, each well tends to give a record characteristic of itself. Consequently, when charts from many wells are gathered as they have been in this study, they present an amazing variety of records.

The hydroseismic data on waterlevel recorder charts are affected by the time scale and the vertical or gage-height ratio. Those recorders equipped with 1:30, 1:24, and 1:12 gage-height gears are most likely to record the extreme upper and lower limits of the fluctuation. Those recorders with 1:1 and 1:2 ratios are most likely to record the aftershocks. Thus there is no one ratio that is best for recording hydroseisms. Those recorders with a time scale of 2.4 inches per day are more likely to show hydroseismic detail than are those with a time scale of 0.3 inch per day.

Of the 716 wells for which data have been tabulated, copies of charts from 433 wells showing the earthquake record have been received by the author. Of these, 12 were from pressure recorders and 421 were from float-type re-Table 1 (next page) corders. shows the various gage-height ratios and time scales at which these 421 records were made. The gageheight scale is chosen to give a record that minimizes "background noise" such as pumping effects and tides, and emphasizes water-level trends. The time scale is chosen primarily to get the

Gage-height ratio		Recorder time scale, in inches per day							
	0.3	1.0	1. 2	1.8	2.3	2.4	9. 6	576	
:30						3			
$egin{array}{c} : 24 & \dots & \dots & \dots \\ : 12 & \dots & \dots & \dots \\ \vdots & \vdots & \vdots & \vdots & \dots \\ \vdots $			19 19			9			
: 10 : 6	$\begin{bmatrix} -1 & 28 \\ 2 & 2 \end{bmatrix}$		$\frac{20}{39}$		3	58			
:5 :3	54 	2	$\begin{array}{c} 10 \\ 2 \end{array}$			1			
$egin{array}{c} : 2.4 _ _ _ _ \\ : 2 _ _ _ _ _ \end{bmatrix}$	50			4		2			
: 1.2 : 1	$\begin{bmatrix} -1 \\ -2 \end{bmatrix}$ $\begin{bmatrix} 1 \\ 73 \end{bmatrix}$	1	20		5	8 1	1		

Table 1.—Gage-height ratio and time scales of 421 float-type recorders that registered the Alaska earthquake

maximum of record with a minimum cost. From the recorders operating with a time scale of 2.4 inches per day, the arrival time of a seismic disturbance can be determined at best to the nearest 10 minutes. For recorders operating at 1.2 inches per day, time can be read only to the nearest 20 minutes. On recorders operating at 0.3 inch per day it is extremely difficult to determine a time closer than 1 to 2 hours.

The time-compressed record of water-level recorders in normal use has prompted the development and operation of several recorders operating at a greatly expanded time scale. Blanchard and Byerly (1935, 1936; Byerly and Blanchard, 1935) were the first to install such a recorder. By maintaining it on a well in California, they were the first to obtain a hydroseismic record similar to a seismogram. From it, they identified several types of waves (P, PP, PPP, S, L, and M).

Since 1947, an expanded-scale water-level recorder has been developed and maintained by E. E. Rexin on a well at Milwaukee, Wis. Details on this instrument have been published previously (Rexin, 1952, 1960; Rexin, Oliver, and Prentiss, 1962), and a few of the records have received detailed

seismological study. Through this, Rexin, Oliver, and Prentiss (1962) identified nine types of waves additional to those reported by Blanchard and Byerly. These waves were PKS, SKS, PS, PPS, SS, SSS, LQ, LR₁, and LR₂. This recorder was in operation at the time of the Alaska earthquake, and its hydroseismogram is the only one known from this earthquake. The record is discussed in "Hydroseismograms from Nunn-Bush Shoe Co. well," (p. C10).

In recent years the Rensselaer Polytechnic Institute has built and operated expanded-scale recorders on several shallow and deep wells in New York State as part of a research program in explosion detection (Katz, 1961, 1962, 1963; Katz, Carragan, and Michalko, 1962a, b; Carragan, Katz, and Michalko, 1963; and Carragan, Michalko, and Katz, 1964). Their work is of great interest, and the equipment they assembled is the most advanced of any used at present. Lack of funds caused operation of these recorders to be discontinued shortly before the Alaska earthquake.

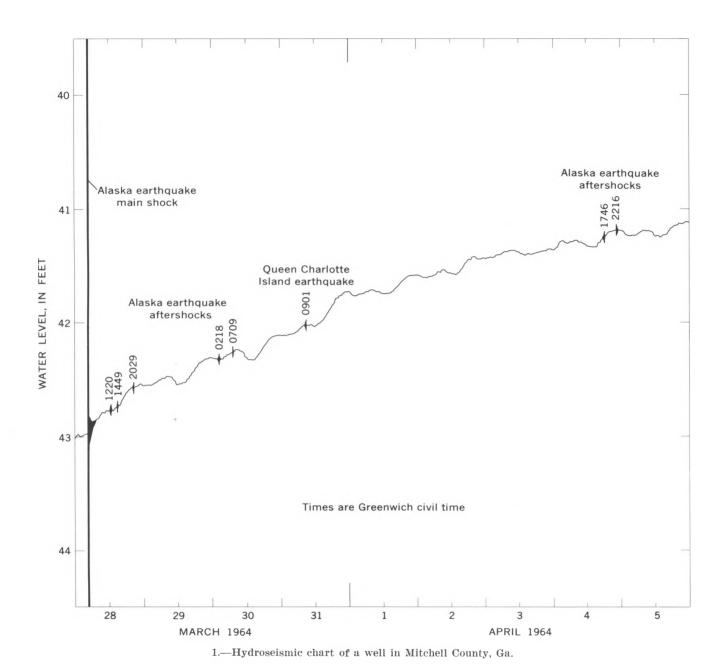
The U.S. Geological Survey operated sensitive expanded-scale recorders on a well in Arizona and a well near St. Augustine, Fla., but lack of funds caused operation to be discontinued about a month before the Alaska earthquake. No data obtained from these recorders have been published.

RECORDABLE HYDROSEISMIC DATA

HYDROSEISMS IN WELLS

The water-level recorders currently maintained on observation wells can at best record only limited data from any seismic event. Recordable hydroseismic data include the following:

- 1. The depth to water in the well at the start of the seismic record.
- 2. The maximum seismically caused water-level rise.
- 3. The maximum seismically caused water-level decline.
- 4. For the largest seismic events, some wells record a coda portion during a period of 1-2 hours following the maximum fluctuation during which the fluctuations decrease to static level.
- 5. In especially sensitive wells, the surface wave that took the long way around the world (the W₂ wave) may record as a distinct fluctuation, and, likewise, the wave traveling in the opposite direction (the W₃ wave) may record as a still smaller but distinct fluctuation. These waves have been identified for the first time on records from the Alaska earthquake.
- 6. A change in water-level trend due presumably to seismically caused changes in the aquifer such as increase or decrease in transmissibility and enlargement or contraction of fractures.
- 7. The water level at the end of the seismic disturbances,



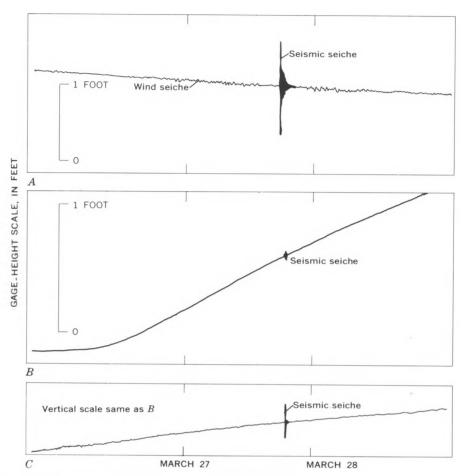
which may be quite different from the level at the start for any of a number of reasons: Seismically induced change in aquifer stress, change in barometric pressure, change in tidal cycle, pumping effects,

8. The approximate time of the disturbance.

or recharge.

One of the more detailed hydroseismic charts (fig. 1) shows the very large fluctuation of water level that occurred due to the Alaska earthquake, seven aftershocks, and the Queen Charlotte Island earthquake of March 31. On the type of recorder in operation, the pen reverses when it reaches the edge of the chart, so the

size of the fluctuation is known to have been in excess of 5 feet with at least 3.5 feet presumably accounted for in the rise. Because many wells tend to have a water-level rise equal to the decline, it is reasonable to assume that water in this well fluctuated more than 7 feet. From examination of this chart, one can see how relatively



2.—Typical records of hydroseisms caused by the Alaska earthquake. *A*, Record at Blakely Mountain Dam on Lake Ouachita, Ark. Record furnished by U.S. Corps of Engineers. *B*, Record at Castor Creek near Grayson, La. *C*, Record at St. Francis River at Marked Tree, Ark. Central standard time.

few data can be taken from any one chart. From the above list of possible data that can be recorded, only items 1, 4, 7, and 8 can be read from the record shown in figure 1.

The most detailed of all the hydroseismic records are from the Nunn-Bush Shoe Co. well in Milwaukee, Wis. From March 27 to 30, the water levels fluctuated in response to the arrival of P, S, and L waves from the Alaska earthquake and aftershocks. The records (figs. 8, 9) are discussed in detail on page C10.

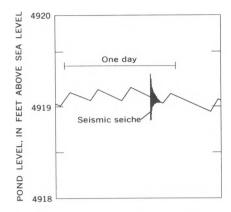
HYDROSEISMS FROM SURFACE-WATER GAGES

SEISMIC SEICHES

Most of the 753 surface-water hydroseisms were recorded as seismic seiches. Included in this category are the fluctuations starting with a maximum oscillation that gradually diminished to a steady water level (fig. 24). A variant type includes some seiches that began with a small oscillatory rise and fall, increased with time, and then died back to normal (fig. 2B). Another variant is a sudden

rise and fall of water level consisting of only one or a few oscillations. That the number of oscillations is small is known from the narrow width of the pen line on the chart (fig. 2C).

One of the largest of the seismic seiches was recorded by the U.S. Corps of Engineers on Lake Ouachita in Arkansas at the Blakely Dam Headwater Gorge (fig. 2A). This record shows 1.45 feet of fluctuation gradually diminishing to "normal" over about $2\frac{1}{2}$ hours.



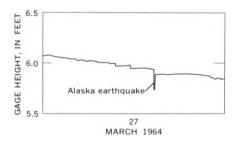
3.—Hydroseism caused by the Alaska earthquake, recorded at the disposal pond of the National Reactor Testing Station, Idaho.

One of the best recorded seiches was from a pond at the National Reactor Testing Station in Idaho (fig. 3). Although the seiche had a maximum rise and fall in water level of only 0.56 foot, the oscillations continued over a period of 2 hours. The pond bottom is in alluvial sand and gravel of the Big Lost River. The sand and gravel overlie basalt. This geologic setting seemingly is favorable for the generation of a seismic seiche in the pond.

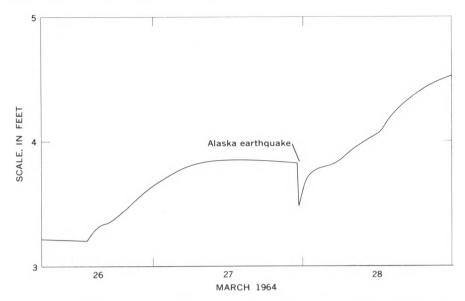
SEISMICALLY INDUCED CHANGE IN STAGE

An unusual record is the one from Little Haw Creek near Seville, Fla. (fig. 4). At the time the seismic waves reached the gage, the water stage began to decline and dropped 0.33 foot in about 15 minutes. Then the trend reversed and the water level began to rise.

Another unusual record was obtained on Tensas River at Tendal, La. (fig. 5). The water level sud-



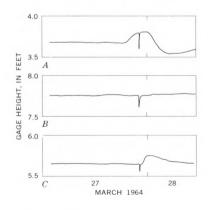
5.—Bubbler-gage record of Alaska earthquake for Tensas River at Tendal, La. Central standard time.



4.—Seismically induced decline in stage at Little Haw Creek near Seville, Fla. Eastern standard time.

denly declined 0.20 foot, then rose only 0.15 foot and remained level for 7 hours even though the trend before and after the quake was a slow steady rise. Because the bubbler-type gage on which this was registered has a built-in delay, rapid fluctuations do not record. The first motion detected by the instrument was a water-level decline.

Three surface-water recorders in Kansas also registered the quake as a temporary decline in stage. A small sharp decline in water level was followed by a slightly less rapid rise to the preearth-

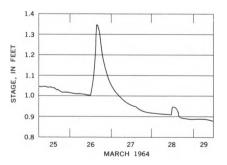


6.—Bubbler-gage records of the Alaska earthquake from Kansas. A, Big Blue River near Manhattan. B, Neosho River near Chanute. C, Neosho River near Burlington. Central standard time.

quake level (fig. 6). Inasmuch as all three charts were recorded by bubbler gages, the traces may not represent a true decline in stage. The author has not yet been able to establish whether rise or fall is related either to the location of the gage on the cross section of the stream or to the seismic waves.

SEISMICALLY INDUCED FLOW(?)

Among the 755 charts examined, only one appears to show what may be seismically induced flow. The hydrograph from March 25–29, 1964, of Paxton Creek in Pennsylvania shows (fig. 7) the effect of a rain in the basin on March 26. On March 28 a smaller rise



7.—Stage of Paxton Creek, Pa., showing increased flow on March 26 from 0.36-inch rainfall and possible seismically induced discharge from an aquifer on March 28, 1964. Eastern standard time.

is shown, but the rise seems quite rapid and is followed by a gradual decline to or below the preearthquake trend. There was no known rainfall that could have caused this rise. The time is inconclusive, for the rise occurred 13 hours after the earthquake. This increased flow has been interpreted by Louis Carswell (oral commun., September, 1964) as reflection of a slug of ground water squeezed out of an aquifer into the creek some miles upstream. If this supposition is true, the slug then retained its identity for 13 hours during its travel to and past the gage.

The hydrologic effects recorded

at surface-water gages, although far smaller in size than many of the hydroseisms in wells, are of interest because they are so unusual. Never before, to the author's knowledge, have seismic seiches been reported in flowing streams at great (teleseismic) distances from an epicenter.

HYDROSEISMOGRAMS FROM THE NUNN-BUSH SHOE CO. WELL. WISCONSIN

By Elmer E. Rexin and Robert C. Vorhis

The most detailed of all hydroseismic records of the Alaska earthquake are from the Nunn-Bush Shoe Co. well in Milwaukee, Wis. These were obtained from a recorder built and maintained on the well by Rexin. The recorder operates with a chart speed of 576 inches per day and magnifies the fluctuations to five times their natural size.

The well, which is at Fifth and Hadley Streets in Milwaukee, was drilled in 1925. It has a 10-inch casing with welded joints to a depth of 107 feet, an 8-inch casing from 104 to 215 feet, and an 8-inch open hole from 215 feet to the bottom at 400 feet.

The artesian aquifer penetrated by this well is formed by the Waubakee and Niagara Dolomites of Silurian age. This aquifer characteristically is not uniformly permeable, and water occurs chiefly in joints and along bedding planes.

On the night of March 27, 1965, as the watchman marked 22:00 hours on the chart of the water-level recorder, he found that something quite violent was being recorded. He immediately called Rexin to report that the float was banging down in the well, the water was gurgling, and the pen was flying back and forth from end to end of the recording drum.

The senior author arrived 20 minutes later and verified that the chart (fig. 8) was recording a major earthquake. The preliminary movement was recorded at 21:43: 20 c.s.t., March 27 (03:43:20 G.c.t., March 28) with a clear and distinct initial drop in water level of 0.005 foot. Movement continued small and somewhat indecisive for 2½ minutes, then the water level quickly rose 0.034 foot. This was followed by a decline, and rhythmical movements were recorded for the next 4½ minutes. Then movement became more violent (apparently owing to arrival of the S wave) for a period of less than a minute. A lesser motion (in the sense that the motion was barely within the recording limits of the instrument) followed for a 3-minute period. Immediately thereafter the water level began to fluctuate so violently that the range of movement exceeded the limits of the recorder. The period of these violent fluctuations was about 15 seconds each. The maximum movement during this phase could not be measured but was estimated to have been about 12-14 feet.

Large waves with periods measured in minutes, in addition to the 15-second waves, are suggested if one sketches in a line that connects the midpoints of fluctuations.

The more significant details shown on these hydroseismograms (figs. 8, 9) are tabulated below (table 2) along with the epicentral times of the quake and the aftershocks that were recorded.

The hydroseismograms from this well are truly unique in that they are the only expanded-scale records showing in detail the effect of the Alaska earthquake on water levels. As such, they will undoubtedly be subjected to much detailed study in the years ahead. Rexin's observations of the many earthquakes recorded in this well have shown that the long-period waves such as followed the Alaska quake are invariably associated with major earthquakes that also generate tsunamis. He believes that this aspect may have an importance in itself that will make further study worthwhile.

Table 2.—Chronological list of hydroseismic data from the Nunn-Bush Shoe Co. well at Milwaukee, Wis., March 28-30, 1964

[Greenwich civil time]

8).
03:49_____ Arrival of S wave (fig.

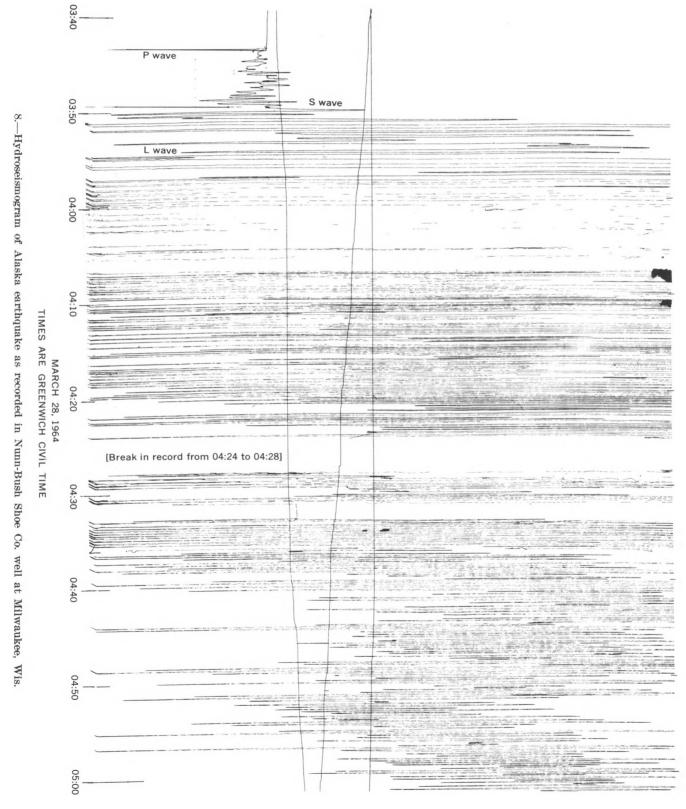


Table 2.—Chronological list of hydroseismic data from the Nunn-Bush Shoe Co. well at Milwaukee, Wis., March 28-30, 1964—Continued

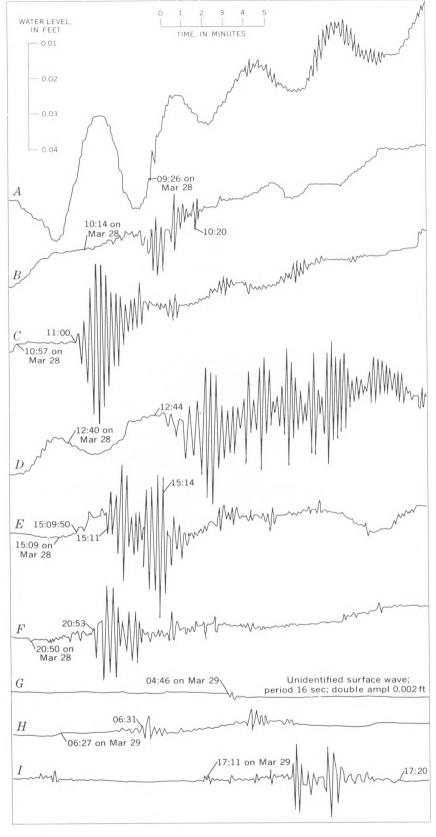
March 28, 1964-Continued

- 03:52_____ Start of L(?) wave (fig. 8).
- 03:55_____ Start of major water-level oscillations.
- 04:24_____ Pen ran out of ink and had to be refilled.
- 04:28...... As recording resumed,
 the oscillations began to decrease in
 amplitude. The
 record suggests that
 there was a "superwave" of about 24
 minutes in period
 and, as the record
 continued, the period
 gradually lessened to
 about 4 minutes.
- 05:12_____ Water level began a slow steady rise that continued to 11:18 G.c.t. and, as measured from the charts, represents a minimum rise of 1.64 feet.
- 06:17_____ A distinct water-level rise and decline that occurred over a $2\frac{1}{2}$ -minute interval.

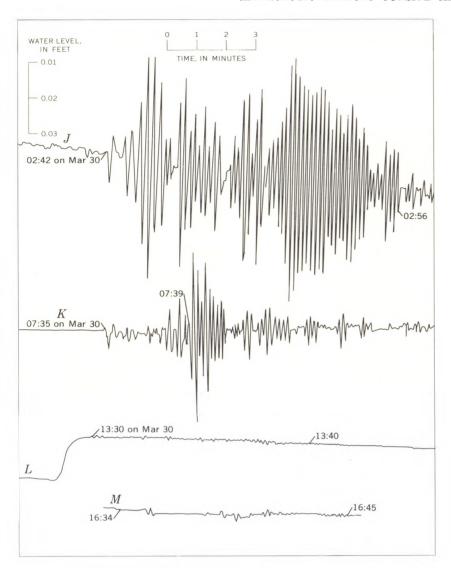
 End of chart 1.
- 09:26_____ Start of distinct train of 16-second waves with maximum double amplitude of 0.008 foot superimposed on long continuing waves of 4-minute period (fig.
- 09:52:54____ Aftershock of magnitude 6.2 (as determined at Pasadena, Calif.).

9A).

- $10:14\frac{1}{2}-10:20$. Record of aftershock, with wave of 16-second period and maximum double amplitude of 0.016 foot (fig. 9B).
- 10:35:39____ Aftershock of magnitude 6.3 (Pasadena).
- 11:00_____ Aftershock recorded as a distinct train of 18-second waves



9.—(above and on p. C13).—Hydroseismograms of aftershocks of the Alaska earthquake recorded in the Nunn-Bush Shoe Co. well. Greenwich civil time.



March 28, 1964-Continued

with maximum dou-

ble amplitude of 0.046 foot (fig. 9 C).

12:20:48.8.... Aftershock of magnitude 6.5 (Pasadena).

12:44-13:02... Aftershock recorded as a distinct train of 18-second waves with maximum double amplitude of 0.037 foot (fig. 9 D).

End of chart 2.

14:47:38.7 and Two aftershocks of magnitudes 6.3 and 6.5, respectively (Pasadena).

15:02:40_____ P? wave. 15:09:50_____ S? wave. 15:11_____ L wave.

March 28, 1964—Continued

15:14_____ L maximum with

period of 13 seconds and double amplitude of 0.04 foot (fig. 9E).

20:29:05.9 Aftershock of magnitude 6.6 (Pasadena).

20:50 S? wave.

20:54 L wave.

20:54 L maximum with period of 16 seconds and double amplitude of 0.26 foot (fig. 9F).

March 29, 1964

04:46..... Small surface wave of unidentified origin

March 29, 1964—Continued

with period of 16-seconds and double amplitude of 0.002 foot (fig. 9G).

06:04:43.4____ Aftershock of magnitude 5.8 (Pasadena).

06:31...... L maximum with period of 18 seconds and double amplitude of 0.0068 foot (fig. 9H).

22:10...... Measured depth to
water was 87.07
feet giving a waterlevel rise of 12.13
feet since preearthquake measurement on March
27, 1964.

End of chart 3.

16:40:59.3 Aftershock of magnitude 5.8 (Pasadena).

17:11-17:20___ Distinct wave train with waves of 16-second period and double amplitude of 0.02 foot (fig. 9I).

March 30, 1964

02:18:05.6 Aftershock of magnitude 6.6 (Pasadena).

07:09:34____ Aftershock of magnitude 6.2 (Pasadena).

07:35_____ S? wave.

07:39_____ Surface waves with 12-second period and maximum double amplitude of 0.048 foot (fig. 9K).

13:03:34.7____ Aftershock of magnitude 5.3 (Pasadena).

13:30-13:40___ Very faint waves with period of 14 seconds and maximum double amplitude of 0.002 foot (fig. 9L).

16:09:27.2____ Aftershock of magnitude 5.5 (Pasadena).

16:34-16:45... Waves with maximum period of 25 seconds and double amplitude of 0.0022 foot (fig. 9M).

End of chart 4.

GEOGRAPHIC DISTRIBUTION OF HYDROLOGIC EFFECTS

AFRICA LIBYA

A good record of the Alaska earthquake was made on a recorder in Wādī Labdah near Homes, Libya (Fituri Deghaies, Libyan Ministry of Agriculture, written commun., Feb. 15, 1965). The well (3236–1417–B) is about 4 kilometers from the Mediterranean Sea, has a depth of 77 meters and a diameter of 8 inches. It yields 40 cubic meters per hour. The fluctuation had a double amplitude of 0.24 feet, and the rise was equal to the decline. Three other wells at Bir al Ghanam, Wādī al Magānin, and Qasr Khiar recorded the quake, but no other details have been furnished (Hadi Ali Tarhuni, Libyan Ministry of Agriculture, written commun., Dec. 25, 1964).

REPUBLIC OF SOUTH AFRICA

About 100 charts from observation-well recorders in the Republic of South Africa were examined, but only three showed a fluctuation caused by the Alaska earthquake (O. R. Van Eeden, Di rector, Geol. Survey, written commun., Sept. 8, 1965). Two of the wells are on Robben Island (lat 33°49′ S. and long 18°22′ E.). They penetrate Malmesbury Hornfels of Precambrian age and are 135 and 74 feet deep. The Alaska earthquake caused a fluctuation at about 04:00 G.c.t., March 28 of 0.23 foot in the shallower well and 0.20 foot in the deeper well.

The third well, at Fauresmith (lat 29°45′ S. and long 25°20′ E.),

penetrates shale and sandstone of the Beaufort Series of the Karroo System (Permian-Triassic in age) and is 130 feet deep, and the depth to water is about 13 feet. The Alaska earthquake caused a fluctuation of 0.60 foot.

SOUTH-WEST AFRICA

The Alaska earthquake was registered in an observation well at Windhoek (Dr. W. L. Van Wyk, Assistant Director, Geol. Survey of South-West Africa, written commun., Aug. 25, 1965). The recorder chart, with a time scale of 12 mm per day and a gage-height ratio of 1:5, had a fluctuation of 0.50 foot at about 05:00 G.c.t. The well is 600 feet deep, in quartzite and mica schist. The water was struck in a fault zone, and the water level is 100 feet below land surface.

UNITED ARAB REPUBLIC (EGYPT)

The Alaska earthquake was recorded in an artesian well in Kharga Oasis in the Western Desert of Egypt. The initial response was a fluctuation with a double amplitude of 0.079 meter (0.24 ft) followed about 13/4 hours later by a fluctuation of 0.030 meter (0.09 ft), and 5 hours after the initial response by a fluctuation of 0.007 meter (0.02 ft). (R. L. Cushman, U.S. AID-USGS engineer, written commun., January 1966.)

ASIA

ISRAEL

The Alaska earthquake was recorded in eight observation wells, of which three were in the mountains and five in the coastal plain of Israel (M. Jacobs, Director, Water Comm. Israel, written commun., May 19, 1965). The double amplitudes of fluctuations ranged from 0.003 to 0.075 meters (0.01–0.25 ft).

REPUBLIC OF THE PHILIPPINES

Hydroseisms from the Alaska earthquake were recorded in 17 of 25 instrumentally equipped wells on the Island of Luzon, Republic of the Philippines. The hydroseisms ranged from 1 to 15 centimeters (0.03–0.46 ft) in double amplitude. (A. B. Delena, Bureau of Public Works, written commun., April 13, 1966.)

AUSTRALIA

In the Northern Territory, two observation well recorders, 16 miles southeast of Darwin, were in operation at the time of the Alaska earthquake and both registered hydroseisms (R. N. Eden, Director of Water Resources, written commun., Aug. 11, 1965). These two are 170 feet apart, are at lat 12°-30'35" S. and long 131°04'50" E. Well M1, with a depth of 114 feet, had a fluctuation at about 04:00 G.c.t. of 0.10 foot. Well M2, with a depth of 227 feet, had fluctuations of 2.25 feet at 04:20 G.c.t., 0.98 foot at 06:00 G.c.t., and 0.08 foot at 07:20 G.c.t.

A seismic seiche was recorded at gaging station 113A on the Victoria River at lat 16°22′ S. and long 131°06′ E. This station was the only one of a large number in operation in the Northern Territory of Australia that registered any effect of the Alaska earthquake. The seiche had a double

amplitude of 0.033 foot and was recorded at 04:20 G.c.t.

A recorder on the Tatangara Reservoir in New South Wales, at lat 35°47′53″ S. and long 148°-39′44″ E., recorded a seiche at 04:20 G.c.t. on March 28, 1964, that also was caused by the Alaska earthquake.

Twelve charts from water-level recorders of the Victoria State Electricity Commission were examined closely for unusual movements on March 28, 1964, by G. Patterson, Engineer for Design and Construction (written commun., Oct. 19, 1965). On only one was there any discernible fluctuation. This gage on the Melicke Munjie River recorded a seismic seiche at about the time of the Alaska earthquake.

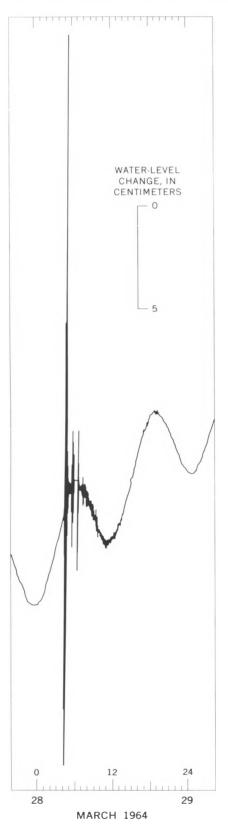
EUROPE

BELGIUM

An extremely interesting and unusual hydroseism was recorded at Heibaart, Belgium (fig. 10). A large fluctuation preceded the maximum one, whereas all the hydroseisms recorded in the United States seemingly showed the maximum fluctuation at the start of the record. Another unusual feature of this hydroseism is that the waves that went the long way around the world (W2, W_3, \ldots were recorded distinctly. A copy of the hydroseismogram was received from A. Sterling, Director, Hydraulic Research Laboratory, Belgian Ministry of Public Works, and is reproduced as figure 10.

DENMARK

The Alaska earthquake was recorded in 7 of 14 observation wells in Denmark (Andersen, 1965, p. 40). The largest double amplitude was 0.12 meter (0.40 ft) in a well that is 66 meters deep, cased to 62 meters, and 10 inches in diameter. The well produces from



10.—Hydroseism of the Alaska earthquake recorded at Heibaart, Belgium. Greenwich civil time.

limestone in which the casing is seated. The water level in the well at the time of the quake was 14 meters below land surface.

UNITED KINGDOM

The effect of the Alaska earthquake was recorded on 34 wells in England (J. Ineson, Chief Geologist, Water Dept., Geol. Survey and Museum, written commun., Feb. 12, 1965). Of these, 25 are in Cretaceous chalk, 7 in Jurassic limestone, 1 in Permo-Triassic sandstone, and 1 in the Lower Greensand. The maximum fluctuation was 1.08 feet in the Lincolnshire Limestone of Jurassic age. The maximum fluctuation in chalk wells was nearly as large, being 1.05 feet. These wells range in depth from 200 to 1,271 feet.

CANADA ALBERTA

Hydroseisms of the Alaska quake were recorded in 24 of 48 observation wells in Alberta. Fluctuations ranged from greater than 5 feet to less than 0.02 foot. Three records showed a permanent change in water levels after the quake. One of these is interpreted by Gabert (1965) to result from stress induced in the aguifer by the quake and which was dissipated gradually with time. None of the records showed any of the aftershocks even though Alberta is closer to the epicenter than any other geographic area from which hydroseisms were reported.

Thirty seismic seiches from surface-water gages in Alberta were reported (R. H. Clark, Secretary, Canadian Natl. Comm. Internat. Hydrologic Decade, written commun., Sept. 21, 1965). These ranged from 0.01 to 0.32 foot.

BRITISH COLUMBIA

Three observation wells in British Columbia that penetrate unconsolidated Pleistocene clay, till, and

sand failed to record any effect of the Alaska quake (E. Carl Halstead, Canada Geol. Survey, written commun., Oct. 30, 1964).

The Alaska earthquake was registered on many surface-water recorders in operation in British Columbia. On most of these the quake was recorded as a small jog on the chart. A total of 13 seismic seiches ranging in size from 0.05 to 1.25 feet are reported from British Columbia by R. H. Clark (written commun., Sept. 21, 1965). In addition, 10 others, of which 7 are illustrated, are given by Wigen and White (1964b, p. 6, figs. 2, 4). A peculiarity in the distribution of the seiches was noted by H. T. Samsden (District Engineer, Canada Dept. of Natural Resource, written commun., Dec. 1, 1964). None of the recorders on Vancouver Island or in the Koanagan River and Lake system in British Columbia registered any effect of the earthquake.

MANITOBA

Wells in the Red River Valley near Winnipeg, Manitoba, showed fluctuations greater than 1 foot. These wells penetrate an artesian aquifer in the Red River Formation of Ordovician age. The aquifer is in fractured carbonate rocks and is confined by till and glacial-lake clay (Scott and Render, 1965, p. 264).

In a tabulation of seiches in Canada (R. H. Clark, written commun., Sept. 21, 1965), seven are listed from surface-water gages in Manitoba. The largest fluctuation, 0.39 foot, was at a gage on Nelson River; the smallest, 0.03 foot, was at a gage on Lake Manitoba.

NORTHWEST TERRITORIES

Wigen and White (1964b, p. 6) list a seiche of 0.30 foot at Cambridge Bay. R. H. Clark (written commun., Sept. 21, 1965) lists four

other seiches in the Northwest Territories: Talston River (0.15 ft), Willowlake River (0.03 ft), Great Bear Lake (0.22 ft), and Lockhart River (0.08 ft).

ONTARIO

Three out of 20 instrumentally equipped wells of the Ontario Water Resources Commission recorded hydroseisms of the Alaska earthquake (B. A. Singh, Division of Water Resources, written commun., Jan. 3, 1966). Near Toronto, two wells in a gravel aquifer recorded hydroseisms with double amplitudes of 0.14 and 0.08 foot. The third, a well in a sand and gravel aquifer in the County of Perth, recorded a double amplitude of 0.08 foot.

A well record in the Ottawa area also showed the Alaska earth-quake. This well, which penetrates an unconfined aquifer, showed an initial increase of 0.20 foot in water level followed by a decline of 1.1 feet and a recovery to the original level after several days (Scott and Render, 1965, p. 267).

Four small seismic seiches at stream gages in Ontario are reported by R. H. Clark (written commun., Sept. 21, 1965): Gull River (0.03 ft), Skootamata River (0.04 ft), Mississagi River (0.07 ft), and French River (0.03 ft).

SASKATCHEWAN

R. H. Clark (written commun., Sept. 21, 1965) reported five seismic seiches from surface-water gages in Saskatchewan: Buffalo Pound (0.06 ft), Fond-du-Lac River (0.07 ft), Weyburn Reservoir (0.05 ft), Deloraine Reservoir (0.45 ft), and Long Creek (0.32 ft).

A farmer in Saskatchewan reports that on the day following the Alaska earthquake his well water had a distinctive purple color. Believing that the ejector was re-

sponsible for the discoloration, he opened the well and pulled the casing to check on the ejector. The farm well is 6 inches in diameter, 111 feet deep, and had the ejector set at 90 feet. The static level normally was 44 feet but when the well was opened the level was at 69 feet, about 25 feet lower than normal. Prior to the Alaska quake, the purple coloration had never appeared. When sampled on April 24, 1964, the well water still had a purple color. The purple color faded gradually and by midsummer had disappeared. (W. Nemanishen, Saskatchewan Dept. of Agriculture, written commun., June 4, 1965).

UNITED STATES

Hydroseismic effects were reported virtually throughout the United States, although New England and the States east of the Appalachians did not register many hydroseisms. New Jersey, however, reported 40 hydroseisms in wells but only 1 from surfacewater gages. Vermont reported none from a well but two from gages. Hydroseisms were most numerous and of largest size in the southeastern States, the ones, surprisingly, that are most distant from the epicenter.

Hydroseisms in the United States are listed by State in table 3, and are broken down into ground-water observation wells and surface-water gages. Listed also are the maximum well and gage fluctuations recorded in each State.

Data on individual hydroseisms in wells caused by the Alaska earthquake are given in table 7 (p. C39).

ALABAMA

Hydroseisms from the Alaska earthquake were recorded in 20 observation wells scattered through-

Table 3.—Number and maximum hydroseisms recorded in the United States from the Alaska earthquake of March 27, 1964

Aska	I I		ion wells	Surface-water gages		
Alaska			double amplitude		double amplitude	
Visconsin 17 3. 5 6 .02	labama_ laska	3 12 5 42 1 0 No report 92 24 18 24 21 1 20 37 1 48 15 11 31 3 9 5 0 40 12 9 3 32 6 1 19 4 0 8 4 21 28 14 0 1 7	$\begin{array}{c} >10\\ >24(?)\\ 1.1\\ 3.3\\ 2.4\\ 3\\ 2.4\\ 3\\ 2.4\\ 3\\ 3\\ 2.4\\ 3\\ 3\\ 3.9\\ 3.9\\ 3.9\\ 3.9\\ 3.9\\ 3.9\\ 3.$	32 9 41 27 14 0 No report 93 26 5 5 8 16 3 12 4 69 0 16 1 1 22 18 16 14 0 1 1 (?) 27 4 1 3 25 3 7 8 8 9 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0. 22 1. 53 . 35 1. 45 . 42 . 30 66 . 22 . 17 . 56 . 10 . 39 . 02 . 34 . 57 . 68 1. 83 . 03 . 90 . 87 . 10 . 18 18 18 10 . 18 10 . 18 10 . 18 10 . 18 10 . 18 10 . 18 10 . 18 10 . 18 10 . 18 12 . 14 . 14 . 05 12 . 14 . 42 . 67 . 06 . 23 10 . 12 . 14 . 23 . 34 . 35 12 . 14 . 25 . 34 . 34 . 35 35 . 36 . 36 . 37 . 38 . 38 . 38 . 39 . 39 . 39 . 39 . 39 . 39 . 39 . 39	
v y 0111111g	Vyoming	2	2. 0	12	. 08	

out the Valley and Ridge, Piedmont, and Coastal Plain provinces of the State. A water-level fluctuation of more than 10 feet in one well in Jefferson County (Jef-1) that is equipped with 1:10 gage-scale gears was indicated by the

fact that the drum made more than one rotation. Drums in three other wells equipped with 1:2 and 1:1 gears also made complete rotations. In a well in Lawrence County (Law-2), the water motion was so severe that it caused

the beaded cable to jump off the grooved pulley of the drum.

Seismic seiches were recorded at 25 gaging stations on rivers in Alabama. The maximum double amplitude of 0.22 foot was recorded at Buttahatchee River below Hamilton, Ala. A double amplitude of 0.18 foot was recorded at two gaging stations, one on the Tennessee River at Triana, the other on Locust Fork at Sayre.

ALASKA

For hydrologic effects in Alaska, see Waller (1966a, b).

ARIZONA

Water levels in wells in several areas in Arizona fluctuated as a result of the Alaska earthquake. The water level in a well in Avra Valley near Tucson, on the fringe of a highly developed agricultural area where large amounts of ground water are withdrawn for irrigation, fluctuated about 8 inches, and lesser fluctuations continued for several hours after the initial shock. The water level in another well near Phoenix fluctuated about 6 inches as a result of the earthquake. Other measurable changes in water level occurred near Bowie where ground water is under artesian pressure. Hydroseisms were recorded in 10 wells in the Colorado River valley. The largest hydroseisms were in two of these wells; fluctuations exceeded 1 foot, but only one well recorded any aftershocks.

The largest seiche in the State was recorded at Coolidge Dam on the San Carlos Reservoir. The maximum double amplitude was 0.35 foot, and fluctuations continued for nearly 2 hours. Seiches were recorded at five other gages, a minor drop in stage was recorded for the earthquake at two other gages, and a slight trace was recorded at another.

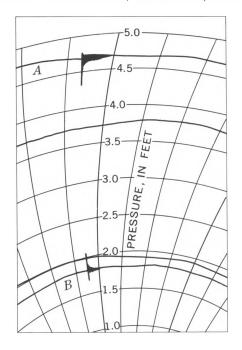
ARKANSAS

The hydroseisms reported from five wells in Arkansas were all rather large: Two that rotated the recorder drum showed movement in excess of 1 foot; the other three ranged from 1.49 to 3.30 feet in double amplitude. Even though the hydroseisms are large, none of the wells recorded any aftershocks.

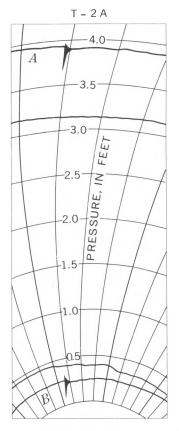
Nearly all the 41 hydroseisms from surface gages in Arkansas were recorded as seiches. largest was 1.45 feet on Lake Ouachita near Hot Springs (fig. 2A). The record from Piney Creek near Dover seems to show a second seiche recorded an hour later than the main shock. The amplitudes were 0.24 foot for the first seiche and 0.03 foot for the second. The gage on South Fork of Ouachita River near Mount Ida recorded a seiche of 0.11 foot followed by a drop in stage of 0.015 foot. At Six Mile Creek subwatershed near Chismville, the earthquake was recorded as a brief 0.03-foot decline in stage.

CALIFORNIA

The hydroseisms recorded in California were rather uniformly small. The maximum reported was 2.39 feet, and only 3 wells out of 42 had movement greater than 1 foot; however, some of the records are unusual. Two adjacent piezometers, 7E2 and 7E4, in T. 6 S., R. 10 W., Orange County, recorded hydroseisms, one from an aquifer (fig. 11A) at 90-120 feet, the other from an aquifer (fig. 11B) at 300-330 feet. The contrast between these two records is interesting because the upper aquifer registered a rise of only 0.03 foot but a fall of 0.17 foot. In the lower aguifer the relative movement was the reverse—a rise of 0.11 foot and a fall of 0.06 foot.



11.—Two dissimilar hydroseisms recorded from two aquifers tapped by adjacent piezometers in Orange County, Calif.



12.—Two similar hydroseisms recorded from two aquifers tapped by adjacent piezometers in Orange County, Calif.

In two other similarly adjacent piezometers, 1Q4 and 1Q6, in T. 6 S., R. 11 W., the hydroseisms recorded almost identically in both the upper aquifer (fig. 12A) at 70-170 feet and the lower one (fig. 12B) at 300-360 feet. These two wells are the only ones known to the author in which hydroseisms from the same quake have been recorded for two different aquifers. The second deepest well known from which a hydroseism has been recorded is in Fresno County (well 19S/17E-35N1). It is 2,030 feet deep (measured depth, 1,955 ft), and the casing is perforated from 608 feet to bottom.

In California, 27 hydroseisms were recorded at gaging stations. On the chart from Lower Twin Lake near Bridgeport, the seiche was recorded during a 4-hour period. The gage at Lake Success near Success, Calif., registered a 0.02-foot rise over about a 10-minute period at the time when the quake was recorded at the other gages in the State.

The gage on LaFayette Reservoir east of Berkeley showed fluctuations above the normal water level but none below. The earthquake-induced, water-level movement continued for possibly as long as 4 hours, but the maximum rise was only 0.02 foot. The Yuba River at Englebright Dam recorded a seiche that seemingly lasted about 8 hours. The gage at Merced River diversion showed a 0.01-foot permanent drop in water level at the time of the earthquake. The largest seiche, of 0.42 foot, was recorded on the gage at Chabot Reservoir, and fluctuations died down in about 3 hours.

COLORADO

Three recorders were in operation in Colorado wells, but only one recorded a hydroseism. It is on the flood plain of the Arkansas River in southeastern Colorado.

The distribution of hydroseisms at Colorado gaging stations was unusual. Fourteen were recorded on the western slope of Colorado, but not one was recorded in the entire eastern half of the State. About 40 stations were out of operation due to ice conditions at the time of the quake. No doubt some of these stations would have recorded the earthquake if they had been operating. The largest seiche of 0.30 foot recorded at White River near Meeker was unusual; for so large a fluctuation, there was no coda portion. The water level returned instantaneously to normal level.

CONNECTICUT

Neither wells nor stream gages in Connecticut recorded the earthquake.

DELAWARE

No report received from Delaware.

FLORIDA

The Alaska earthquake gave Florida two distinctions. Even though it is the State farthest from the epicenter, more wells and more streams in Florida recorded the earthquake than in any other State. It furnished 92 well records compared with 48 from Michigan, which had the second largest number. Likewise, the fluctuation of 17 feet in a well (Taylor 35) at Perry, Fla., is the second largest recorded fluctuation for any well outside Alaska and is the largest reported fluctuation in an openhole well. The earthquake evidently caused violent water movement in some wells, especially in the Tampa area, for there one recorded pen was dislodged, and at six wells the beaded cable was thrown off the recorder pulley. Of the 92 wells in which the earthquake was recorded, 49 had fluctuations with a double amplitude greater than 1 foot. Aftershocks were recorded in only one well, that at Perry, Fla.

A few wells in Florida were residually affected by the earthquake. A well in Clay County on the crest of a water-table high recorded a "normal" hydroseism of large size, but immediately afterward the water level began a slow decline (fig. 13). This decline continued for several weeks until the water level finally stabilized about 4 feet below the preearthquake level. This prolonged decline may indicate that the earthquake caused water-table highs to be lowered slightly by somehow facilitating drainage. Water levels in other wells seemed to show a change in trend coincident in time with the earthquake. A well in Hardee County (731-145-1) showed a sudden drop of 0.4 foot coincident with the initial phase of the earthquake, and a "coda" portion then registered at the lower level. From the chart it would appear that the drop in level must have some physical significance.

Practically every surface-water gage in Florida recorded the Alaska earthquake. The records were so numerous that copies of only the 93 best records were submitted. The maximum hydroseism at a gage was 0.66 foot.

The gages on some tidal streams and canals of coastal Florida are equipped to record both stage and deflections of velocity vanes. From the record of the deflections, the changes in velocity and in direction of flow can be calculated. Two examples of such records are shown in figure 14. The records of these gages promise some interesting interpretations if studied further.

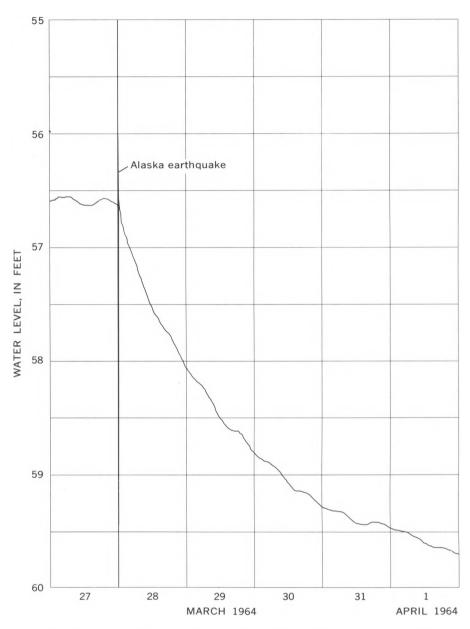
GEORGIA

The hydroseism recorded in a Piedmont well at the Georgia Nuclear Laboratory, Dawson County, enabled the author to score a scientific "first." examining the recorder chart on the day after the earthquake he realized from study of previous hydroseisms from this well that the earthquake was of great magnitude. He telephoned his findings to the Atlanta Journal-Constitution, and the Sunday paper reported that the quake was greater than an 8.3 magnitude and "may be bigger than any quake yet recorded instrumentally." This is the first and only instance known where a hydroseism has provided an estimate of earthquake magnitude as quickly as one furnished by seismologists.

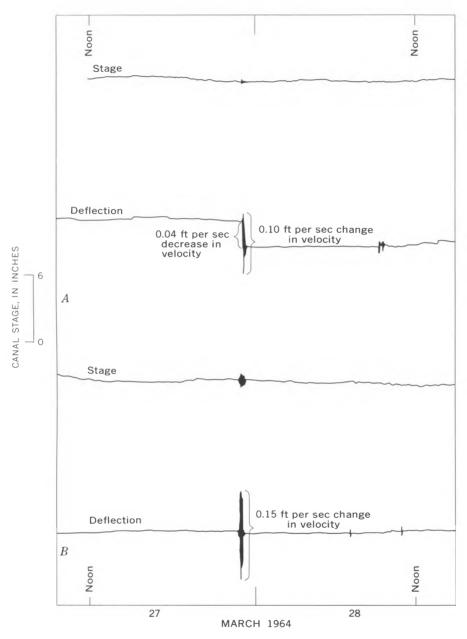
Of the 24 hydroseisms recorded in Georgia wells, 20 were larger than 1 foot. Seismic seiches were also recorded at 26 gaging stations scattered throughout the Valley and Ridge, Piedmont, and Coastal Plain Provinces (fig. 15). Most of these stations are on fairly deeply entrenched streams. The maximum double amplitude was 0.18 foot.

No seiches were recorded or reported from Brunswick or Savannah on the Atlantic coast. This absence was unexpected because seiches were so large and numerous on the gulf coast.

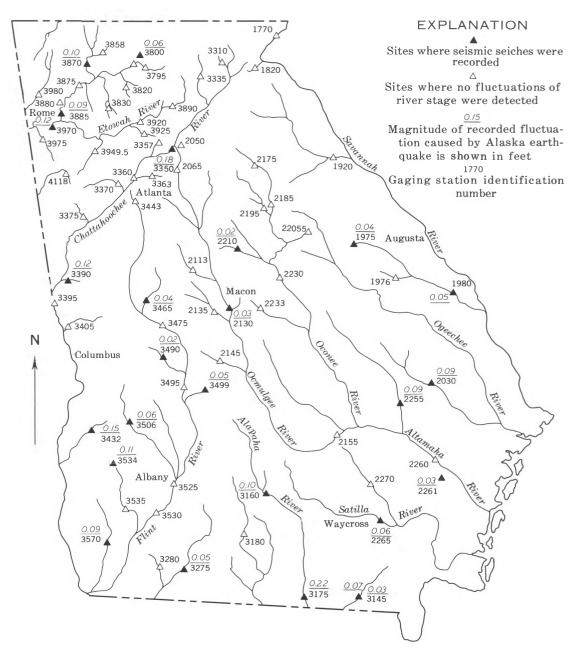
At Brunswick, the water levels began to rise in all the wells immediately after the quake, and a slow steady rise continued for about 15 days. The measured rise was 3.3 feet in well E-143, 2.9 feet in well J-35, 3.0 feet in well J-36, and 2.6 feet in well J-67. The water level in well J-67 continued to rise after the others leveled off. Residents of the area reported that after the earthquake their wells yielded water containing black sooty material. The earthquake seemingly produced a surge so violent that it loosened black iron sulfide that had gradually coated well casings,



13.—Part of a 4-foot decline caused by the Alaska earthquake in water level in a well in Clay County, Fla. Eastern standard time.



14.—Alaska earthquake shown on stage and deflection records. A, Biscayne canal near Miami, Fla. B, Snake Creek Canal at North Miami Beach, Fla. Eastern standard time.



15.—Map of Georgia showing locations of gaging stations and size of seiches recorded from the Alaska earthquake.

pipes, and iron fixtures of water systems. Others reported that former flowing wells began again to flow. Most of these were old wells that had been drilled to a water-bearing unit of sand and calcareous sand that lies above the principal artesian aquifer. These old wells are from 450 to 500 feet deep and obtain water from sand at depths of 350–450 feet.

The rise in ground-water level due to the earthquake was seemingly a permanent change in the Brunswick area. When piezometric maps for the end of 1962 and 1964 were drawn and compared, the seismic boost in water level made the two maps look similar despite increased consumption that normally would have caused a decline in regional water levels.

The spring flow used as a public supply at Cave Spring and water from the city supply wells at Cedartown became turbid at the time of the earthquake and remained so for several days. The earthquake coincided with extremely heavy rainfall, so it is not certain whether one or the other or both were the cause of this temporary deterioration in water quality.

HAWAII

Sixteen hydroseisms reported from wells in Hawaii had recorded double amplitudes ranging from 0.05 to 1.85 feet. A comparison of seven of these with the tidal efficiencies of the wells suggests that tidal efficiency of a well has no relation to the amplitude of a hydroseism.

The largest fluctuation in Hawaii was not recorded in a well but in a horizontal tunnel 1,614 feet long. The tunnel, driven into a mountain for water, has the innermost 24 feet shut off by a 10-foot bulkhead that holds water at 160–180 feet of pressure. At the time of the Alaska earthquake the water was discharging and pressure

was only 126 feet, or 55 p.s.i. (pounds per square inch). It was this pressure that fluctuated 4.60 feet owing to the earthquake.

Five seiches were recorded at gages on the Islands of Kauai and of Hawaii, the largest having a double amplitude of 0.17 foot. The others were all small and hardly noticeable on the charts. No seiches were recorded at gaging stations on the Islands of Oahu, Maui, or Molokai.

IDAHO

A total of 24 hydroseisms was reported from Idaho. The most outstanding is from a well in Latah County where the double ampliture was more than 5 feet and where nine aftershocks were registered. No other well in the State is known to have recorded more than one of the aftershocks.

The largest seiche reported from Idaho was recorded in a pond in Butte County (fig. 3). The depth of water in the pond at the time was 14.10 feet. The maximum double amplitude was 0.56 foot, and the fluctuations continued for about 2 hours and diminished slowly to a static level. The pond bottom is in alluvial sand and gravel of Big Lost River. Only four other seiches were reported from gages in Idaho. These were 0.04 foot or less in size and were all in the Idaho Falls area.

The earthquake worsened the pumping problem at the Clayton Silver Mines in Custer County, Idaho. After the earthquake the flow of water in the mine increased from 750 to 1,150 gallons per minute.

ILLINOIS

A total of 28 hydroseisms was reported from Illinois wells. Only two aftershocks were recorded, and both were registered in well DuPage ANL-10.

Seismic seiches were recorded in Illinois at two lake stations: Wolf Lake at Chicago and Money Creek at Lake Bloomington. The se seiches were both recorded at 10:00 p.m. c.s.t. on March 27, 1964, and had a double amplitude of 0.08 foot at the former and 0.05 foot at the latter. The Illinois Water Survey reports that two river gages, one on the West Branch, the other on the East Branch of DuPage River, recorded seiches from the Alaska earthquake. The fluctuations were 0.04 foot and 0.03 foot, respectively.

A well in Cook County (37N 14E-22.1b), which taps a Cambro-Ordovician sandstone and has a depth of 1,648 feet, reportedly pumped sand following the quake. Two wells in Union County reportedly yielded muddy water after the quake.

INDIANA

Twenty-one hydroseisms having double amplitudes ranging from 0.08 to 8.25 feet were reported from Indiana wells. Well Marion Ma-32 was exceptional in that it recorded 12 aftershocks, one of the most complete records in any well in the United States. This well is equipped with a recorder operating with 1:1 vertical gears, so the Alaska earthquake itself was shown only as greater than 1 foot, but oscillations of this amount or more continued throughout at least a 2-hour period.

Seismic seiches were recorded at 16 stations in Indiana. Of these, the maximum fluctuation was 0.39 foot at an auxiliary gage on the White River near Nora where fluctuations were recorded over a period of 55 minutes. Four of the seiches were on lakes and one was on a reservoir.

IOWA—ALASKA EARTHQUAKE EFFECTS ON GROUND WATER

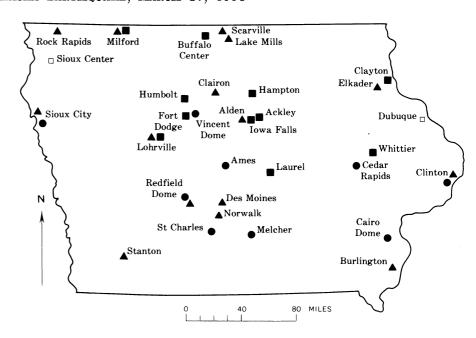
By R. W. COBLE

The Alaska earthquake caused the water levels to fluctuate in many wells in Iowa. The earthquake occurred at 9:36 p.m. c.s.t., and the L wave arrived in Iowa about 9:50 p.m. The L wave was calculated to have arrived at Loras College Seismograph Station, Dubuque, Iowa, at 9:52 p.m., c.s.t., or 03:52 G.c.t. (Dr. William Stauder, St. Louis Univ., written commun., Feb. 26, 1965). The timing mechanisms on the water-level recorders on wells in Iowa are not precise enough to determine the exact minute that the earthquake affected the aquifers in the State, but there were many indications that something happened just before 10 p.m.

Aquifers in Iowa responded to the earthquake waves as shown by (1) the seismic fluctuations on some recorder charts; (2) turbid water in some wells and springs, probably caused by the disturbance and movement of silt, clay, and colloidal particles within the aquifers; and (3) in some wells a permanent change, either a rise or fall, of the water level. These effects are summarized in figure 16 and in table 4.

At Redfield Dome, the water levels in several different aquifers showed various types and amounts of seismic fluctuations as is shown in figure 17. In this same area, two of the four observation wells drilled to the St. Peter Sandstone of Ordovician age showed a seismic fluctuation; the other two showed no effect. Why some wells are affected and others are not is yet to be determined.

The best record of a seismic fluctuation is shown on a recorder chart from an observation well in the Franconia Sandstone of Cambrian age at Vincent Dome (fig. 18). A seismic fluctuation of 0.23 foot occurred just before 10 p.m. A series of smaller fluctuations were recorded after the main one. Many of these can be matched with some aftershocks; however, many aftershocks were not recorded.



16.—Location of reported ground-water disturbances in Iowa caused by the Alaska •, seismic fluctuation; A, permanent change; I, turbid water.

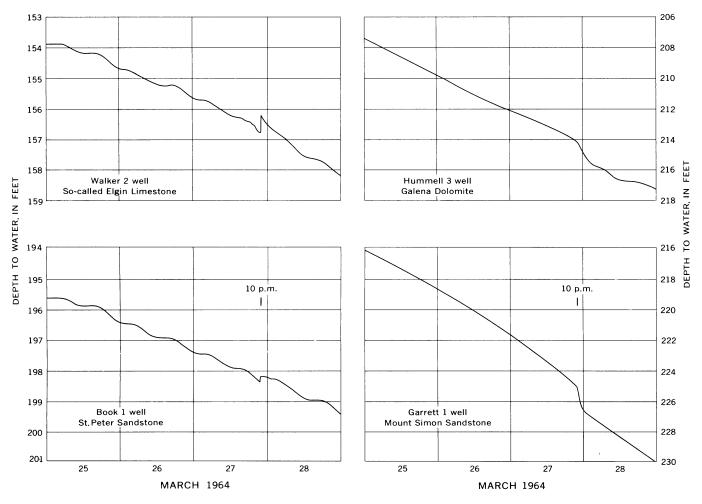
Table 4.—Summary of ground-water disturbances in Iowa caused by the Alaska earthquake

	Aqu	ifer	Effect			
Locality	System	Lithology	Seismic fluctua- tion (ft)	Turbid water	Water-level chang lasting more than I week	
Ackley				×		
Alden Ames					Lowered.	
Ames	Quaternary	Sand and gravel	0.15			
Buffalo Center				×		
Burlington	Mississippian	Carbonate			Lowered 2 ft.	
0	and Devonian.					
	Ordovician	Sandstone			Lowered 1.8 ft.	
Cairo Dome 1		Carbonate	1.2			
	Ordovician	do	4.7(+?)			
Cedar Rapids	Silurian	do	.4`			
Clarion					Raised.	
Clayton				×		
Clinton	Cambrian and	Sandstone	''		Raised 15 ft.3	
Des Moines	do	do			Raised 18 ft.3	
Elkader	do	do			Raised 40-50 ft.3	
Fort Dodge				×		
Hampton						
Humboldt	Mississinnian	Carbonate		🗘		
Hampton Humboldt Iowa Falls		our bond to : : : : :		×××		
Lake Mills				l ^ i	Several reported	
·					lowered. Sev- eral reported raised.	
Laurel	Mississippian	Carbonate		×		
Lohrville	do	do		×	Lowered 10 ft.3	
Melcher Milford	Quarternary	Sand and gravel_	. 05			
Milford				X	Lowered.	
Norwalk					Raised.	
Redfield Dome 4	Ordovician	Carbonate	. 5			
	do	do			Lowered 1 ft.	
	do	Sandstone	1.25			
	do	do	. 25			
	Cambrian	do			Do.	
Rock Rapids	Cretaceous	do			Raised 8 ft.3	
St. Charles	Mississippian	Carbonate	1.1			
Scarville	Mississippian Quaternary	Sand and gravel.			Lowered 5 or 6 ft.	
Sioux City	Cretaceous	Sandstone	² 2. 5+		Raised 5 ft.3	
Stanton					Lowered 30 ft in	
					2 days. Raised	
					40 ft after 7 days	
Vincent Dome 4	Cambrian	Sandstone	. 23			
Whittier						

¹ Data from Natural Gas Pipeline Co. of

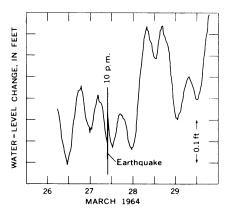
² Pumping rate fluctuated.

Known to have lasted more than 7 months.
 Data from Northern Natural Gas Co.



17.—Water-level fluctuations at Redfield Dome, Iowa. Central standard time.

The seismic fluctuation was so rapid in observation wells drilled to the Dakota Sandstone of Cretaceous age at Sioux City and the Ordovician dolomite at Cairo



18.—Seismic fluctuations in the Peterson 1 well at Vincent Dome, Iowa. Central standard time.

Dome that the recorder pens became disengaged from the float-pulley mechanisms. At Sioux City, the water-plant operator could feel air moving in and out of the well casing as the water level fell and rose. He noted that this "sucking and blowing" of air, which gradually increased in intensity and then slowly diminished, lasted from 5 to 10 minutes.

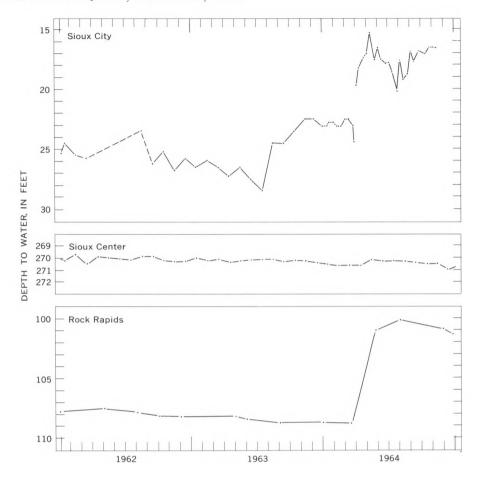
Several wells produced turbid water after the earthquake. The water generally became clear after a few hours or a few days of normal pumping. Similarly, water from several springs, the water supply for Humboldt, also became turbid. These springs, on the bank of the West Fork of the Des Moines River, flow from limestone

of Mississippian age. This water had always contained less than 5 ppm (parts per million) of suspended matter. On the morning of March 28, the turbidity ranged from 70 to 80 ppm. Nearby, water from several small springs that discharge through the river bed was observed to be red, brown, or blackish brown. The turbidity diminished on March 30, but increased again after a few rainy days during the first part of April. It did not completely disappear for another $2\frac{1}{2}$ weeks.

In several localities the groundwater levels seem to have been permanently changed. At Sioux City, the water level in an observation well, tapping the Dakota Sandstone, rose 6 feet and remained high for at least the rest of the year (fig. 19). At Rock Rapids, 75 miles north of Sioux City, the water level in another well drilled to the Dakota Sandstone rose 8 feet. A third well, bottoming in the same aquifer at Sioux Center, which is almost midway between Sioux City and Rock Rapids, showed no seismic fluctuation or permanent change whatsoever.

Limestone of Mississippian age yields water to the municipal well at Lohrville. The nonpumping water level had been 97–98 feet below the land surface for more than 1 year before the earthquake (fig. 20). On March 28, the water level had dropped 3 feet, and after 1½ months the total drop was 10 feet below the original level. This diminished level persisted through the rest of the year.

The water level in a well in the Jordan (Cambrian) and St. Peter (Ordovician) Sandstones at the Ford Motor Co. plant in Des Moines rose 18 feet (from 101 to 83 ft) after the earthquake. The level was still high in June 1965. The town of Elkader has several wells that produce water from the Jordan-St. Peter interval. Shock waves from the Alaska earthquake affected all of them in the same manner—the water level rose 40-50 feet and has remained high (fig. 21). In the city of Clinton, adequate records are available for two of the wells which produce water from the Cambrian and Ordovician interval (Mount Simon Sandstone and several overlying sandstone formations through the Prairie du Chien Group). Immediately following the initial shock, the water level rose more than 20 feet in city well 7 (fig. 21). Seismic fluctuation was inferred in city wells 3 and 7 in that the pumping-rate recorders show a total fluctuation of more than 1

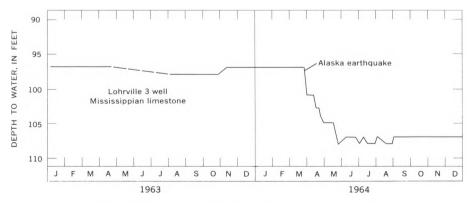


19.—Water levels in the Dakota Sandstone in northwest Iowa.

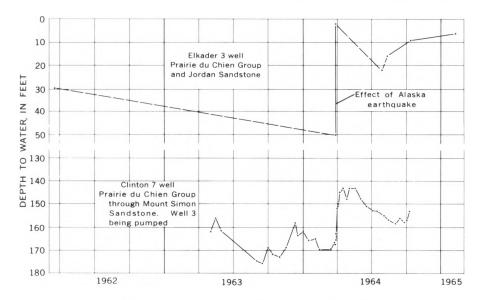
percent just at 9:55 p.m. c.s.t. (03:55 G.c.t.).

A permanent change in water levels implies a change in the physical properties of the aquifers. A logical assumption is that the porosity and thickness of the aquifers have decreased where the water levels rose and increased where they fell. This change need

not be large, even where the level increased as much as 50 feet as at Elkader. Such a change would require only a change of 22 psi in the hydrostatic pressure in the aquifer. Considering only the Jordan Sandstone, which is 100 feet thick near Elkader, and assuming that it has a porosity of 15 percent, the compression of the



20.—Water levels in the city well at Lohrville, Iowa.



21.—Water levels in wells at Elkader and Clinton, Iowa.

water would have to be only about 1.5×10^{-5} feet to raise the pressure 20 psi. This amount of compression would decrease the thickness of the aquifer from 100.000000 feet to 99.999985 feet and the porosity from 15.000000 percent to 14.999985 percent. These computations take into account only the compression of the water. If the sandstone itself were compressed, as it probably would be, the thickness of the aquifer would be decreased somewhat more than $1.5 \times$ 10⁻⁵ feet. This decrease in porosity is extremely small and can be considered insignificant with respect to the productivity of the aquifers.

The earthquake was recorded at two surface-water gages in Iowa. Shell Rock River at Northwood declined 0.02 foot in stage between 03:00 and 04:00 G.c.t. on March 28. The stage was steady then until 05:30 G.c.t. and rose 0.01 foot by 06:00 G.c.t. (midnight). A seiche of 0.02 foot was recorded on Lake Ahquabi.

KANSAS

In Kansas, in only 1 out of 12 observation wells is a hydroseism from the Alaska earthquake known to have been recorded, and

it had a double amplitude of 0.37 foot. Only 7 surface-water recorders out of 150 or so in operation gave noticeable evidence of the Alaska earthquake. Gaging stations close to those that were affected went through the period without recording the slightest change. Three gages that responded noticeably to the earthquake each showed (fig. 6) a sudden drop in stage with complete recovery in 15-30 minutes. There was no rise above normal level at any of the three stations. These three hydroseisms were recorded at stations equipped with bubbler gages, so the response probably reflects instrumental failure to record rapid fluctuations. some interesting-looking records of the earthquake may be worthless as far as indicating true water-level response to the seismic waves.

KENTUCKY

Twenty hydroseisms were recorded in a total of 60 observation wells in Kentucky, but only one well had a fluctuation greater than 1 foot. This well is at Mammoth Cave National Park.

Four seismic seiches were recorded by gages. The largest, 0.57

foot, was recorded on Buckhorn Reservoir at Buckhorn. The next largest, 0.40 foot, was recorded on Nolin River Reservoir near Kyrock.

The Louisville Courier-Journal (Mar. 31, 1964) carried an article describing the effect of seismic seiches on two other Kentucky lakes:

Witnesses said water about 4 miles from Dix Dam at Lake Herrington slopped around like it does in a dishpan, but people at either end of the lake reported nothing unusual.

The superintendent of Lake Cumberland State Park confirmed * * * reports by fishermen of a series of mysterious waves that swept across Lake Cumberland at about the time of the Alaska quake. Superintendent John Flanagan said the waves were a foot to 18 inches high, and snapped two cables on the Jamestown boat dock. Other reports told of the lake falling and rising from 3 to 4 feet several times. The boat-dock operator called up and said the lake was acting funny—calm in the middle but whirling in circles near the shore.

LOUISIANA

A total of 37 hydroseisms was reported from Louisiana wells with double amplitudes ranging from 0.04 foot to greater than 5 feet. One record (EB-90) is from a well 2,120 feet deep cased to 2,025 feet; this is the deepest well in the Nation from which a hydroseism was reported. Eight identifiable aftershocks were recorded in 1 well (SJB-17). In no other well in the State are aftershocks known to have been recorded.

Seismic seiches along the gulf coast and in the bayous were large enough to cause destruction. The New Orleans States-Item (Mar. 28, 1964) reported:

Boats were sunk and some roads in communities close to the Gulf of Mexico were flooded by a wave that rolled in, then subsided. Other boats were torn from their moorings. At Golden Meadow on Bayou Lafourche a big oyster vessel was thrown against a store building on the bayou. For a brief instant a

foot and a half of water covered roads on both sides of the bayou at Golden Meadow and neighboring Galiano. Grand Isle, located right on the Gulf, apparently didn't get a ripple from the strange wave action. At Delacroix Island, where several boats were washed from their moorings, one man crawled along his dock to land on hands and knees to keep from getting washed away by the tide.

Between midnight and 01:00 c.s.t. on March 28, 1964, in the midst of a TV bulletin giving the first news of the Alaska earthquake, a special bulletin announced that a tidal wave had struck the Louisiana coast, sunk small boats in Chef Menteur pass, and is now entering Lake Pontchartrain.

A large barge-mounted drilling rig on location in Lake Ponchartrain at lat 30°09.6′ N. and long 89°56.8′ W. experienced a seismic wave also. The barge was lifted approximately 2 feet as tanks were being flooded to sink it to the bottom. Only one wave was noted. Tugs in the Industrial Canal, which with the Gulf Seaway connects Lakes Pontchartrain and Borgne, reported they experienced momentary tides of 6 feet or more (Rex Meyer, written commun., June 12, 1964).

Almost all the gages throughout Louisiana recorded seismic seiches from the Alaska earthquake. The largest was 0.90 foot in double amplitude. Many others were of large size, but none showed a coda portion lasting more than an hour.

MAINE

One hydroseism of 0.19 foot was recorded in a well at Brunswick, Maine. This is one of the few hydroseisms recorded in the New England States. No surface-water gages in the State showed any effect of the quake.

MARYLAND

The earthquake was recorded in four wells in Maryland. Although the four all bottom in sand aquifers, the response of one was markedly different from the other three which showed normal hydroseisms. In the fourth well (Dor-Cd 40) the water-level dropped 0.20 foot during a 25-minute period, but this drop is coincident with a decline due to earth tide, and it is impossible to separate visually the effect of each.

The earthquake was also recorded at three gages on streams, but the maximum fluctuation was only 0.04 foot.

MASSACHUSETTS

The earthquake was recorded in one well in western Massachusetts. The well, which penetrates Stockbridge Limestone, registered a fluctuation of 0.62 foot.

MICHIGAN

The State of Michigan reported 48 hydroseisms in wells—second only to Florida in the number reported; however, only two wells recorded aftershocks. A well in Genessee County, which recorded three aftershocks, is unusual in that it bottoms in an old waterfilled coal mine. This type of construction may possibly be favorable for recording hydroseisms because many have been reported from this well over the years.

Two partly buried reservoirs owned by the city of Lansing and equipped with recording gages showed seismic seiches. In one reservoir, having a capacity of 7 million gallons, a fluctuation of 1.84 feet was recorded; in the other, having a capacity of 10 million gallons, a 1.25-foot fluctuation was recorded. The time of occurrence at Lansing was 03.55 G.c.t. on March 28.

The 16 seiches recorded at gaging stations were all small; the largest was 0.06 foot in double amplitude. Lasting changes in stage were recorded at two gages: a 0.01-foot drop on the Cedar River at East Lansing and a 0.01-foot rise

on the Cass River on the northern peninsula of Michigan.

MINNESOTA

Although some of the 15 hydroseisms reported from Minnesota were rather large, no aftershocks were recorded. In each of two wells in Hennepin County, the water level declined 2 feet in 40 hours following the earthquake. This decline may have been caused by local pumping, but the similarity of record and the timing suggest that the response may represent some local effect induced by seismic waves from the earthquake. In another well, the water level changed so rapidly that the ink of the pen did not flow fast enough to give a complete record.

Only one seiche was reported from surface-water gages in Minnesota. It had a double amplitude of 0.03 foot and was recorded on the Roseau River at Ross.

MISSISSIPPI

The instruments on observation wells in Mississippi recorded 11 hydroseisms. Of these, five were on or near the Tatum salt dome. The response was somewhat anomalous in that the water level declined instantaneously and then took several hours to rise to the preearthquake level.

Of the 22 hydroseisms reported from surface-water gages in Mississippi, all were recorded as seismic seiches, and 10 had double amplitudes of 0.10 foot or greater The largest seiche, 0.90 foot, was recorded on the Pearl River gage at Monticello.

MISSOURI

Of the 28 hydroseisms in wells reported from Missouri, some were quite large and others were quite unusual. One well in Greene County had a fluctuation greater than 10 feet and, following the

earthquake, the water level rose 50 feet between March 28 and June 2, 1964. The same reaction but on a smaller scale occurred in Madison County where well 33N/7E–20bcd had a water-level rise of 5.55 feet in the first 40 hours after the earthquake and an additional 1.65 feet of rise in the next 98 hours. In Polk County well 33N/Z1W–5adc the water level fell 1.1 feet in 1½ hours after the quake.

The largest seiche of 0.87 foot was recorded by the Black River gage at Poplar Bluff, but no coda portion was recorded.

Several stations in Missouri equipped with bubbler gages recorded the earthquake, but water-level change was recorded either as an upward or downward motion, never as both up and down, evidently because of the relative unresponsiveness of this type of instrument.

According to Fellows (1965), many home wells and the municipal wells at Rogersfield and Mansville, in southwestern Missouri, yielded turbid water, some of it reportedly "blood red," for a few hours to a few days following the earthquake.

Fellows (1965, p. 3, 4) reports further that:

Within two months after the quake, static water levels in two deep wells in Springfield rose by several tens of feet. Unfortunately, these wells were not equipped with automatic depth recorders and static water levels were not determined at regular intervals.

Fishermen at Table Rock Lake in Taney County, Mo., observed mysterious waves on the lake the night of the quake.

MONTANA

Hydroseisms were recorded in only three wells in Montana. In one of these wells (Gallatin County well A1–4–25dc) the response was greater than 1 foot for the main quake, and seven of the major aftershocks were recorded. The recorder on this well in alluvium at

Bozeman has been maintained for many years by Dean C. C. Bradley of Montana State College and has registered many other earthquakes.

The Billings Gazette (Mar. 30, 1964) reported that a "small wave developed on Hauser Lake northwest of Helena a few minutes after the Alaska quake and tore a boat dock from its moorings."

Of the 21 hydroseisms from stream-gaging stations in Montana, the largest double amplitude was 0.16 foot. Practically all the records came from the mountainous part of the State, and none were recorded on the main stem of the Missouri River. Gages on the North Fork of Milk River and Sage Creek, both on the international boundary, showed a rise in water level of 0.01 or 0.02 foot at the time the earthquake was recorded at other gages.

NEBRASKA

The largest hydroseism in Nebraska for the seven wells reported was 4.10 feet, but no aftershocks were recorded. This well had a fluctuation 18 times as great as that which was recorded for the Hebgen Lake, Mont., earthquake of August 17, 1959. A well in Thayer County (4–1–9bac) had an unusual response: The water level rose 0.87 foot but at no time declined below the prequake level.

The 14 hydroseisms from Nebraska stream gages were all recorded as seismic seiches. This response was unusual in that seiches were sparsely recorded elsewhere in the northern Great Plains.

NEVADA

All five of the hydroseisms recorded in Nevada were in Clark County wells. The one in well S19/60-9bcc caused the recorder drum to make a complete rotation, and four of the major aftershocks

were recorded. This record is to be expected because hydroseisms have consistently been clearly recorded in this well.

NEW HAMPSHIRE

Hydroseisms were not recorded in New Hampshire wells, but one surface-water gage registered the earthquake.

NEW JERSEY

A total of 40 hydroseisms was reported from wells in New Jersey. Of these, only six had fluctuations greater than half a foot. The maximum fluctuation, 4.37 feet, occurred in Hillside well 4 in Union County. One distinct aftershock was recorded by this well, which is the one in New Jersey most sensitive to earthquake shocks.

Only one somewhat questionable hydroseism was recorded by a surface-water gage in New Jersey. The large number of wells that responded to the quake seem in odd contrast to the one questionable record from a stream gage.

NEW MEXICO

In New Mexico hydroseisms were recorded at 12 observation wells. Of these, two had fluctuations of more than 5 feet, but neither one showed any aftershocks. The Hot Springs well 6 in Sierra County had a fluctuation of more than 1 foot, and some aftershocks were registered. The main shock thus may have caused the recorder drum to rotate many times, so the actual fluctuation must have been considerably greater than 5 feet. At another well the motion caused the pen to pull the paper off the recorder, and in another the motion was so rapid that the ink could not flow fast enough to give a complete record.

C. V. Theis (written commun., Aug. 4, 1964) furnished the following comments:

In New Mexico, the Alaska earthquake produced a fluctuation of the Greenfield observation well in the Roswell artesian aquifer of about 13 feet. This well is comparatively shallow, is just below the lip of the confining beds, and is in a part of the aquifer with transmissibility in the millions. The Artesia recorder well, located where the aquifer is deeper, with a transmissibility of only 100,000 or so, and where the strata are becoming more calcareous near the old reef, had a fluctuation of only a small fraction of a foot. * * * Lea Lake, east of the Pecos and approximately east of Roswell, the largest of the Bottomless Lakes, about 300 feet deep, and an acre or so in surface area, produced a water spout said to be about 15 feet high. Maddox [U.S. Geol. Survey. Roswell] saw old tires and other objects floating on the surface of the lake the next day, these having been cast up from the bottom. The brine observation wells at Malaga Bend near Carlsbad, which fluctuated about a foot from the Turkish earthquake of about 1939, lost the record of the Alaskan quake because the pen of the Friez recorder was thrown over the cylinder.

Roy Foreman, who runs a concession at Lea Lake, N. Mex., observed the effects of the Alaska quake on the lake and the following is a summary of his observations:

About 9:40 p.m., March 27, 1964, waves about 10 feet high rose on Lea Lake. At this time my wife and I heard a loud noise, which sounded like a strong wind although it was a calm evening. The water flowed over a 3-foot high guardrail which is 54 feet from the normal port margin. A section of the guardrail was washed out by the flow of water.

Before the earthquake a small trickle of water flowed from the lake through a 12-inch culvert. The morning after the quake, the 12-inch culvert was carrying a full capacity of outflow. The discharge as of August 1964 was still more than the prequake discharge.

In the Carlsbad area, 6 of 21 gages on flowing streams recorded hydroseisms, but the 6 hydroseisms recorded were all of small double amplitude. Elsewhere in the

State, 21 hydroseisms were recorded of which 19 were seismic seiches and 2 were minor changes in stage.

NEW YORK

Nine hydroseisms were recorded in New York State wells. These did not include Saratoga 529 and Queens 64 which previously had shown rather outstanding responses to earthquakes. Recorders had been removed from both these wells before the Alaska earthquake. The Chautauqua 10 well, however, had a fluctuation of 2.10 feet and also recorded one aftershock. It is interesting to note that this well had a fluctuation of 0.22 foot for the Hebgen Lake earthquake of August 17, 1959. At another well, the recorder pen failed during the Alaska quake.

Only four seiches, each less than 0.01 foot, were recorded at stream gages in New York State. In addition, one record from the Mahwah River near Suffern showed a drop in stage of about 0.01 foot.

NORTH CAROLINA

Three hydroseisms were reported from North Carolina wells; the largest hydroseism had a double amplitude of 1.85 feet.

NORTH DAKOTA

North Dakota reported three hydroseisms from wells that penetrate glaciofluvial sand and gravel.

The earthquake was also recorded at two surface-water gages in the State. The record from one, on the Cheyenne River near Kindred, is questionable in that it seemingly was made an hour earlier than it should have although this may be due to clock error. The record looks like a seiche of 0.05 foot double amplitude followed immediately by a decline in stage of 0.02 foot. The other gage, on Jamestown Reservoir near

Jamestown, showed a rise in water level of 0.07 foot in about 10 minutes, declined 0.05 foot during an approximate 20-minute interval, and then remained steady for the next $9\frac{1}{2}$ hours.

OHIO

Observation wells in Ohio recorded 32 hydroseisms from the Alaska earthquake. The largest fluctuation was 5.8 feet in a well in Van Wert County. Even with so large a fluctuation, the aftershocks were not strong enough to be recorded, perhaps because the gagescale gears were 1:10, and the record was so compressed that the minor fluctuations were obscured.

A total of 188 analog recorders was in operation in Ohio at the time of the Alaska earthquake, but only 25 showed any noticeable effects. The seismic seiches probably were obliterated at a few of the gaging stations because of normal river surging or fluctuation caused by wind action. The maximum double amplitude was 0.25 foot on a lake gage near Jefferson, Ohio. Another gage, 800 feet away on Mill Creek, recorded a sharp drop in stage of 0.04 foot followed by a rapid rise of 0.03 foot. The water level then remained steady at this slightly lower level. The Mahoning River gage at Alliance, Ohio, showed a drop of 0.01 foot at the time of the earthquake. One other gage, at Atwood Reservoir near New Cumberland, also showed a similar reaction.

OKLAHOMA

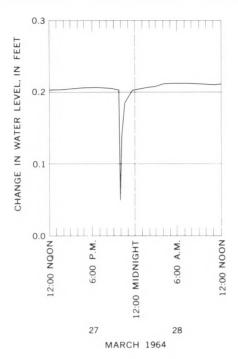
Hydroseisms were noted in wells equipped with water-level recorders in the Oklahoma Panhandle, central, and eastern parts of the State. The fluctuations in water levels were more than 1 foot in wells in the panhandle, in Grady County, and in the Arbuckle Mountains; about 0.4 foot in a well

in Washita County; and about 0.1 foot in one in the Arkansas Valley.

Hydroseisms are tabulated for six observation wells (table 7). In one well the water movement was so rapid that the beaded cable slipped on the pulley of the recorder. Two wells in sec. 25, T. 6 N., R. 18 E. provide an interesting contrast. Both penetrate the Ogallala Formation to a depth of 99 feet, and both are equipped with recorders which show similar responses to barometric changes and rainfall. However, showed rapid water-level fluctuations due to the earthquake. whereas the other showed no response.

At Byrd Mill Spring south of Ada, a surface-water pool showed a drop of 0.15 foot in water level at the time of the quake, and it took about 1½ hours to recover to the preearthquake level (fig. 22). The spring originates along a faulted limestone section in the Arbuckle Group and seemingly the shock wave for a time partly closed the opening along which water flows to the spring.

A total of 45 seismic seiches caused by the Alaska earthquake was recorded in Oklahoma. Of these, the largest had a 0.44-foot double amplitude recorded on Lake of the Cherokees at Langley. Minor decline in stage seemingly caused by the earthquake occurred at gages on Little River near Wright City, Muddy Boggy Creek near Farris, and Verdigris River near Inola. A slight rise in stage seemingly caused by the earthquake occurred at the gage on Sallisaw Creek near Sallisaw. The decline and recovery of water level as described above for Byrd Mill Spring Pond was also recorded at four other gages, but the maximum decline was only 0.04 foot. These four gages are on Lake Texoma near Denison.

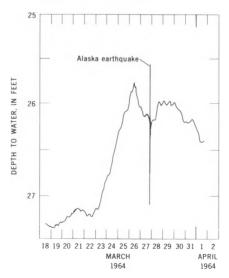


22.—Effect of the Alaska earthquake on the surface-water pool formed by Byrd Mill Spring, Okla. Central standard time

Glover Creek near Glover, Okla., Sand Creek at Okesa, and Verdigris River near Claremore.

OREGON

Only one hydroseism of 0.055 foot was recorded in Oregon wells; however, only three well re-



23.—Hydroseism from well Yo 180 in York County, Pa. Eastern standard time.

corders were in operation at the time of the quake. Inasmuch as Oregon is fairly close to Alaska, unrecorded water-level fluctuations probably occurred in many wells.

Seismic seiches were recorded by eight gages in Oregon. The largest had a double amplitude of 0.14 foot.

PENNSYLVANIA

Among the 19 hydroseisms reported from Pennsylvania wells are some that are unusual. The earthquake records in seven wells in Dauphin, Luzerne, and York Counties all were at the bottom of a "low" superimposed on a water-level "high." This type of response may have been caused by local barometric changes rather than by the earthquake (fig. 23).

A stream gage on Paxton Creek did not record a seiche but it registered a sudden increase in stage 13 hours after the quake. No rain was reported in the basin. This rise in stage may represent water squeezed from an aquifer cropping out upstream (fig. 7).

Only two seismic seiches were recorded in Pennsylvania although 102 gaging stations equipped with analog recorders were in operation at the time. The double amplitudes recorded were 0.05 and 0.04 foot.

PUERTO RICO

Of the four hydroseisms reported from wells in Puerto Rico, one was surprisingly large. The fluctuation measured 3.40 feet and was recorded so fast that the beaded cable slipped on the pulley. This well (Jauca 2) is in a graben near a fault zone.

RHODE ISLAND

No trace of the Alaska earthquake was recorded either at observation wells or surface-water gages in Rhode Island.

SOUTH CAROLINA

The hydroseisms recorded from South Carolina were large. In Beaufort County well 304, the recorded fluctuation was 8.98 feet, but the beaded cable was thrown off the pulley, so the fluctuation may have been even larger. Because of this disruption, no aftershocks were recorded; however, Jasper County well 46, which had a fluctuation of only 4.72 feet, recorded five aftershocks.

Seismic seiches were recorded by eight gages in South Carolina. The largest double amplitude was 0.12 foot.

SOUTH DAKOTA

Only two hydroseisms were reported from South Dakota, but one is the largest recorded for this quake in any well outside of Alaska. This one was recorded on a pressure recorder in a test well drilled to an artesian aquifer, which had an original pressure head of 266 feet above land surface. At the time of the quake the pressure head was 121 feet above land surface, so a pressure recorder of 200 feet capacity was mounted on the well. This unusual situation permitted the full range of the Alaska earthquake-pressure effect of 23 feet to be recorded in the well. The well produces from sandstone of the Opeche (Permian) and Minnelusa (Pennsylvanian and Permian) Formations and is at the northwest edge of the Black Hills. Because $4\frac{1}{2}$ inches on the chart represents 200 feet of pressure, it is not surprising that no aftershocks appear on the chart.

Six seismic seiches were recorded by gages in South Dakota. The maximum double amplitude was 0.14 foot.

TENNESSEE

The following information concerning Tennessee wells is extracted with slight modification from a paper by Hassler (1965): Hydroseisms recorded at Geological Survey wells ranged from a trace to 3.90 feet. The shock was so violent that recorder pens were flipped off the charts at Jellico and New Johnsonville. Two major aftershocks on March 29 and 30 were also recorded at the Capleville (J-1) well. Only two wells (Sloanville, U-1 and U-2) were equipped with Stevens A-35 recorders with large time scales (2.4) in, per day). Records from these instruments indicate the first waves arrived at approximately 10 p.m. c.s.t., and the major fluctuations occurred about 20 minutes later. Water levels in both wells fluctuated for about 3 hours after the major shock.

Hassler (1965) also reported that seismic seiches were recorded at approximately 20 percent (22 gages) of the surface-water gaging stations equipped with analog recorders. This figure would probably be much higher had not many streams been receding from fairly high stages at the time the quake occurred. The duration of the oscillations ranged from about 5 minutes to 35 minutes.

The three largest seiches in the State were all recorded at gages on the Cumberland River: 0.36 foot at Carthage, 0.42 foot below Old Hickory, and 0.42 foot at Rome. The other seiches recorded were all much smaller; ranged from 0.10 to 0.14 foot, and 13 were less than 0.10 foot.

In Dickson, Knox, and Montgomery Counties, muddy water was reported in many wells tapping the Fort Payne Chert. The quake coincided with a period of heavy precipitation, so it is not known whether the heavy rains, the earthquake, or a combination of both produced the muddy water in wells.

TEXAS

Data on some of the hydroseisms from the Alaska earthquake as recorded in Texas wells have already been published. Miller and Reddell (1964) list five wells in the High Plains area that recorded such fluctuations. Montgomery (1964) describes three from wells in Bexar County. Mills (1964) lists 28 recorded in U.S. Geological Survey observation wells in Texas: the data from these 28 records are included in table 7. Aftershocks are known to have been recorded in only two Bexar County wells—five in one and four in the other. Most of the largest hydroseisms occurred in wells penetrating the Edwards Limestone.

A list of seismic seiches in Texas compiled by W. B. Mills (written commun., November 1964) shows that 69 gages were affected by the earthquake. The up and down amplitudes were all equal except for Lake Winnsboro near Winnsboro where the entire motion was down (0.03 ft). The largest double amplitude was 0.68 foot recorded at Sabine River near Ruliff. The next largest was 0.64 foot recorded on Angelina River near Zavalla. The duration of the disturbance at both of these stations was 60 minutes. At Lake Houston near Sheldon, the double amplitude was only 0.13 foot, but the disturbance continued for 90 minutes.

The earthquake was recorded at two stations equipped with bubbler gages: one on the Guadalupe River at Cuero and the other on the Nueces River at Mathis. Both gages recorded only a downward motion of the water level, 0.29 foot at the former and 0.06 foot at the latter. The bubbler gages have a built-in delay of a few seconds. The gages probably responded to the first motion to reach them and

were unable to respond quickly enough to the upward surges of the seiche. Consequently, the records of these gages may be proof that the first motion to affect them was a decline in water level.

UTAH

Fourteen hydroseisms were recorded on observation wells in Utah. Of these, two were on pressure recorders, but the 2-foot and 1.2-foot fluctuations recorded are small compared to the 23 feet for the pressure recorder in South Dakota. Two wells recorded aftershocks: seven in a Tooele County well and one in a Weber County well. Both wells were equipped with recorders that would measure a maximum fluctuation of 1 foot—less than the fluctuation for the main quake. A well that had a measurable fluctuation of 2.50 feet seemingly recorded one aftershock, so the fluctuation in the others probably was more than 2.5 feet.

Eight seismic seiches were recorded at gages in Utah. The largest was 0.06 foot in double amplitude. The others were minor, ranging from 0.01 to 0.03 foot.

VERMONT

Two seismic seiches were recorded in Vermont. The larger seiche, at the Wrightsville Detention Reservoir gage, had a fluctuation of 0.23 foot. The smaller, at the East Barre Detention Reservoir gage, had a fluctuation of 0.06 foot. The general lack of hydroseisms at both wells and stream gages in New England makes it surprising that these were recorded. The gages were both on reservoirs; no seiches were recorded on streams anywhere in New England.

VIRGINIA

One well in Virginia recorded the Alaska earthquake. The well is at Shenandoah National Park, is 280 feet deep, and penetrates metabasalt of the Catoctin Formation of Precambrian (?) age. The total fluctuation was 1.60 feet, which is divisible into a rise of 0.45 foot, an upward displacement of the water level of 0.55 foot that apparently occurred at the time of maximum fluctuation, and a decline of 0.60 foot.

No surface-water gages in the State showed a trace of the earth-quake.

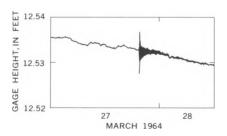
VIRGIN ISLANDS

The most distant hydroseism to be recorded in United States territory was a 0.05-foot fluctuation in a well in the Virgin Islands.

WASHINGTON

Seven hydroseisms were reported from Washington, of which one had a fluctuation of 3.92 feet and was followed by six aftershocks. The other hydroseisms were all relatively small in amplitude and none was followed by aftershocks. In one well, which had an instantaneous fluctuation of only 0.16 foot, the water level rose 1.20 feet during a 4-hour period immediately after the quake and then stayed at this higher level.

The gage on Snohomish River at Snohomish showed a strong fluctuation of about 0.45 foot superimposed on a much larger tidal cycle. This fluctuation occurred at 03:50 G.c.t. on March 27, so it undoubt-



24.—Seismic seiche and wind seiches at Franklin D. Roosevelt Lake at Grand Coulee Dam, Wash. Pacific standard time.

edly was caused by waves from the earthquake.

The gage on Franklin D. Roosevelt Lake at Grand Coulee Dam recorded an interesting seiche (fig. 24). Prior to the seismic seiche a long train of wind seiches had been recorded. Four other smaller but typical seismic seiches were recorded elsewhere in the State.

Several atypical seismically caused water changes were recorded at other gages in Washington. At Whitestone Lake near Tonasket, the gage recorded a sudden 0.03-foot rise of water level followed by the recording of a seiche. Slight residual upward changes in water level were recorded at two gages, and slight temporary changes were recorded at three other gages. The gage at Lenore Lake near Soap Lake recorded seiches from wind all day on March 27, but beginning at 8 p.m. P.s.t. the Alaska earthquake surface waves increased the amplitude of the seiches.

WEST VIRGINIA

Only one hydroseism was found in West Virginia, and it was recorded in a well at the extreme east tip of the State (see table 7). This well in Berkeley County penetrates Beekmantown Limestone of Ordovician age and had a 0.30foot fluctuation.

WISCONSIN

Of the 17 hydroseisms recorded in Wisconsin, the most detailed is the partial record obtained by E. E. Rexin at the well of the Nunn-Bush Shoe Co. in Milwaukee, discussed on page C10.

In three wells, the water level rose and stayed at the higher level. In the Nunn-Bush Shoe Co. well there was an apparent rise of about 12 feet. In another Milwaukee County well the water level rose 7.3 feet after the quake. In a Monroe County well the

water level rose 1.43 feet and remained at this level.

Six surface-water gages in Wisconsin recorded small seiches caused by the Alaska earthquake. The largest seiche was 0.02 foot, and the others were barely visible on the charts. Both the small size and the small number of seiches

that were recorded at surfacewater gages contrast markedly with the large size and the large number that were recorded at wells.

WYOMING

The Alaska earthquake was recorded in two observation wells in Wyoming. In one the motion

was about 2 feet, but no aftershocks were recorded.

The earthquake was recorded at nine stream gages, all in western Wyoming. Thus the distribution of the records corresponds to the distribution in Colorado where effects were recorded only in the western part of the State.

HYDROSEISMS FROM AFTERSHOCKS

The Alaska earthquake generated literally thousands of aftershocks, but few of the major aftershocks occurred near the epicenter of the main shock. Instead, many occurred 2°-8° southwest of the epicenter, and practically none occurred to the northeast. A magnitude-8.4 earthquake normally would have generated a few shocks of magnitude 7, but the largest magnitude of any of the Alaska aftershocks was Those aftershocks that generated a hydroseism in one or more wells are listed in table 5. Two other

earthquakes that occurred during the period from March 28 to April 4, 1964, and that generated hydroseisms are also listed. One occurred on March 31 in the Queen Charlotte Islands of British Columbia. The other occurred on April 3 off the northwest coast of Sumatra.

The aftershocks were not recorded consistently in wells. For example, of the two aftershocks of magnitude 6.6, the one on March 28 at 20:29 G.c.t. was recorded in 14 wells, but the one on March 30 at 02:18 G.c.t. was recorded in 30

wells. Some of the aftershocks reported as having been recorded in one, two, or three wells may have been misidentified.

The aftershocks were recorded in only about 4 percent of the wells in which the main shock was recorded. In general, those wells that recorded the aftershocks are those in which earthquakes are best and most frequently recorded. No aftershocks were recorded at any surface-water gages outside of Alaska. The aftershock records from the more seismically sensitive wells are listed in table 6.

		Epic	enter	Magnitude	Number of wells in
Date	Epicentral time (G.c.t.)	North latitude (°)	West longitude (°)	measured at Pasadena	which recorded
March 28	03:36:12.7	61.05	147. 5	8.4	71:
	09:01:00	56. 5	152.0	6. 2	
	09:52:54	59. 7	144.6	6. 2	
	10:35:39	57. 2	152. 4	6.3	
	11:08:26	60.1	148. 5	6. 2	
	12:20:49	56, 5 60, 4	154. 1 146. 5–147. 1	6.5	2
	14:47-14:49	59. 8	140. 5-147. 1	6. 3-6. 5 6. 6	1
March 29	20:29:06	56. 2	154. 2	5. 8	1
viaicii 20	16:40:59	59. 8	146. 9	5. 8	
March 30	02:18:06	56. 6	153. 0	6.6	3
	07:09:34	59. 8	145. 9	6. 2	1
	13:03:35	56. 5	152.7	5. 3	-
	16:09:27	56, 6	152. 2	5, 5	
March 31	09:01:33	50, 8	130.1	6	1 2
April 2	01:11:56	6.1	2 95. 4	7	3
April 3	22:33:39	61.7	147. 7	6	
April 4	17:46:08	56. 3	154. 5	61/2	1
	22:16:57	59. 5	145. 0	Not reported	

¹ Earthquake in Queen Charlotte Islands, British Columbia.

³ Earthquake off northwest coast of Sumatra.

Table 6.—Seismic fluctuations in sensitive wells, March 28-April 4, 1964
[Greenwich civil time]

									Fluctu	ations,	in feet								
State, county, and well				Marc	eh 28				Mar	ch 29		Mar	ch 30		Mar 31	Apr 2	Apr 3	Ap	ril 4
,	03:36	09:01	09:52	10:35	11:08	12:20	14:47 and 14:49	20:29	06:04	16:40	02:18	07:09	13:03	16:09	09:01	01:11	22:33	17:46	22:10
rizona:						0.005					0.014	0.005			0.000			0.010	
Yuma, (c-11-24)23bcb	>1					0. 025					0. 014	0.005			0, 003			0. 01?	
Taylor, 35	1 17					. 15	0, 20	0.05			.10	. 04			. 06			. 12	
eorgia:																			
Dawson, 12-3	>1	0.042	0.018			. 020	. 014	Tr.?			. 02	. 005			Tr.?			. 012	
Dougherty, 13L4 Mitchell, 10G313	>5 >5					. 065	. 035	. 054		0, 02	. 054	. 016			. 03	0.02		. 048	0, 0
awaii:	>0					. 05	. 046	. 034			. 072	. 042			. 036			. 05	0.0
Oahu, 83	1.85					. 05	. 005		0.008		. 006				. 04				
laho:	1.00								0.000										
Cassia, 35-21E-18bb1	1.44					. 04					Tr.								
Latah. 39N-4W-7	>5		. 05			. 10	. 05	. 14	Tr.	. 05	.10	. 06			. 14	. 05	0.04	. 05	
linois:																			
DuPage, ANL-10	7.70					Tr.?	. 025				. 07				. 03				
ndiana: Marion, Ma-32	>1	. 021	. 020	. 021	0.027	. 052	. 034	. 040	. 008	. 005	. 21	. 044			. 011		. 018	occ	
Pulaski, Pu-6	1 1	. 021	. 020	. 021	0.027	. 002	. 004	. 040	. 008	. 000	. 03	. 02			. 011		. 018	. 066	
ouisiana:	1										. 00	. 02							
St. John The Baptist.																			1
SBJ-17	. >1										. 058	. 012			. 008			. 010	
lichigan:																			
Genesee, 7N-7E-17-1	>2					. 076	. 032				. 08				. 02			. 072	
Kent 6N-12W-34-1 Iissouri:	. >5										. 02	. 012							
Barton, 32N/30W-30cd	. >5					. 07	. 05	. 04		(2)									
Franklin, 44N/1W-27	3 90					. 04	.00	.01		(-)	. 04								
Iontana:	0.00					.01					. 01								
Gallatin, A1-4-25dc	. >1			. 009?		. 002	. 02				. 026	. 024			. 05			. 006	
levada:														100					
Clark, S19/60-9bcc1	. >1					. 03	(3)	. 014			. 022				. 016				
lew Jersey: Union, Hillside 4	4 97					Tr.					. 06								
lew Mexico:	4. 37					11.					. 00								
Sierra, Hot Springs 6	>1	. 012		. 005		. 04	(3)	. 024			. 166	. 02			. 06	22.2.63			
Jew York:		. 012					()	. 0				. 02							
Chatauqua, Cu-10	2.10										. 04								
ennsylvania:																			
Luzerne, Dennison St.	> 0				3.54.34	0.0	0.4				. 076	000			. 022				
outh Carolina:	- > 2				******	. 06	. 04				. 076	. 020			. 022				
Jasper, 46	4.72					. 03	. 035	. 015			. 03	. 025			. 01				
'ennessee:	1, 12					. 00	. 000	. 010				. 020							
Shelby, Sh:J-1	2. 27?										. 034								
'exas:																			
Bexar, F-172	>5						. 01	Tr.?			. 05							. 04	
J-17 Jackson, PP-80-03-101	>3.80										. 02								
Jackson, PP-80-03-101	1 5.8										. 04	. 01			. 02	. 02			
Tooele, (C-3-2)14bad-1_	51		. 007			. 014	Tr.	Tr.			. 032	Tr.			. 025				
Weber, (B-6-1) 30cca-1	\leftrightarrow{1}{1}		. 007			. 012	. 004	. 005	Tr.	Tr.	. 032	. 007		?	. 018				
Vashington:	1 1							. 000		***	. 022	. 001			. 010				
Pierce, 20/3-18c1	3.92					. 03		. 03			. 08	. 03			. 07				
Visconsin:																			1
Milwaukee, M1-120	1		. 016	. 045		. 037				. 02	. 062	. 048	0.002	0.0022					

 $^{^{1}}$ Estimated. 2 Pen ran out of ink? $^{-3}$ Masked by water-level change.

CONCLUSION

This report is the first in which hydrologic effects of a major earthquake have been gathered from throughout the world. Little attempt has been made at interpretation because the major effort so far has been concentrated on assembly of the data. Interpretive studies will possibly show that the areal distribution of effects has some significance. Several interpretive studies are in progress but are not ready.

The theory of "The response of well-aquifer systems to seismic waves" as developed by Cooper and others (1965) goes far toward explaining seismic fluctuations in wells, but no theory as yet explains adequately the lasting change in water level that was observed in many wells. Similarly no theory accounts for the asymmetry of water-level response in wells. The theory of Cooper and others assumes seismic waves to be sinusoi-

dal, whereas some well records show only a brief rise or fall from static level.

Theory as yet explains adequately neither the "draining" of a piezometric high nor the observed "recharging" of a cone of depression, phenomena that were both observed in the weeks immediately following the Alaska earthquake. Thus there is still a gulf between theory and observation.

Similarly, the rigorous mathematical interpretation of seismic seiches by McGarr (1965) goes far toward explaining the many seiches that were recorded. However, one of the two major factors which he states (p. 853) "help to convert the energy of large-magnitude earthquakes efficiently to produce seiches at large distances from the epicenter" is "a very thick layer of soft sediments." According to McGarr (1965), this layer serves

to amplify the seismic motion and especially the horizontal ground acceleration. This factor seemingly is minimal in importance or nonoperative in Colorado, Wyoming, and Vermont. In the western half of both Colorado and Wyoming many seiches were recorded, and two were recorded in Vermont, places where sedimentary deposits are thin to absent, but in eastern Colorado, eastern Wyoming, and through the northern Great Plains, no seiches were recorded although this area is underlain by a sizable thickness of such deposits. Thus, there is also a gulf between theory and observations concerning seismic effects on surface-water bodies.

It is hoped that the data presented in this report will encourage further studies so that the discrepancies that exist between theory and observation can be narrowed and ultimately bridged.

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Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake

Depth to water: Depth in feet below land surface unless otherwise indicated; a preceding plus sign indicates height of the piezometric level above land surface; a preceding plus or minus sign in parentheses indicates depth to water above (+) or below (-) sea level . LSD, land surface datum. Figure in parentheses following depth to water is rise (if plus) or fall (if negative) from the preearthquake to postearthquake water level; this change where reported is not due to water-level trend but is either change in stress on the aquifer caused by the earthquake waves or excessive friction in the recorder installation.

Water-level fluctuation: E, estimated.
Remarks: Because aftershocks can be expected to record only in wells with the largest hydroseisms, failure to record aftershocks is mentioned only for those wells in which aftershocks might have been recorded. Recorder character as to inches of chart per day and gage-height ratio are given only for wells whose charts were examined by the compiler. T, transmissibility; S, coefficient of storage. Other data are as reported.

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water (feet)	From pree	earthquake vel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
				Alabama					
Baldwin, Bal-1 Calhoun, Cal-1 Coffee, Cof-1 Colbert, Col-1	30°24′ 33°42′ 31°19′ 34°46′	87°42′ 85°49′ 85°51′ 87°38′	Sand Conasauga Formation Sand of Lisbon Formation Fort Payne Chert	134/134 213/124 265	24. 6 9. 2 179 7. 5	0. 03 1. 5 . 47 2. 9	0. 02 1. 6 . 28 3. 1	0. 05 3. 1 . 75 6. 0	No aftershocks recorded;
Col-2 Franklin, Fra-1	34°41′ 34°31′	87°41′ 87°44′	Tuscumbia Limestone Tuscaloosa Group and Bangor Limestone.	171 210/146	10. 35(30) 27. 87(06)	1. 60	1, 85	3. 45 >1	0.3 in. per day, 1:12. 0.3 in. per day, 1:1. Drum rotated many
Jefferson, Jef-1	33°26′	86°53′	Bangor Limestone.	140/68	27. 1			>10	times. No aftershocks recorded 0.3 in. per day, 1:1.
Henry, OW-5		85°04′	Clayton Limestone		18. 85	. 55	. 41	. 96	First quake ever recorded in well.
Lawrence, Law-2		87°21′	Fort Payne Chert		15. 0	2. 4	1, 5	3. 9	Cable thrown off pulley; 0.3 in. per day, 1:12.
Limestone, Ct-2		86°58′	do		11. 1	4. 4	3. 1	7. 5	No aftershocks recorded; 1.2 in. per day, 1:12.
Madison, Ct-81 Mad-1 N-48 P-29	34°38′ 34°44′ 34°45′ 34°40′	86°34′ 86°35′ 86°35′ 86°43′	Tuscumbia Limestone Fort Payne Chert Tuscumbia Limestone Fort Payne Chert	188/36 140/69 104/61 119/21	22. 30(10) 51. 65 25. 7 22. 53(04)	. 10 . 07 1. 32	. 17 . 05 1. 31	. 27 . 12 2. 63 >1	Another quake(?) on Mar 29, but no aftershocks
Q-174 Marshall, Mal-2 Monroe, Mon-3	34°42′ 34°20′ 31°31′	86°35′ 86°19′ 87°20′	Tuscumbia Limestone Fort Payne Chert Sand, gravel, and lime- stone.		16. 3 14. 1 62. 28	. 11 1. 6 >. 50	. 15 1. 5 >. 50	. 26 3. 1 >1	recorded; 1.2 in. per day, 1:1. No aftershocks recorded 0.3 in. per day, 1:1.
St. Clair, St. C-1	. 33°35′	86°16′	Floyd Shale, Fort Payne Chert, Maury Forma- tion.	209	. 2	1.0	1. 6	2. 6	o.s m. per day, 1:1.
Talladega, Tal-2 Tuscaloosa, Tus-2	33°10′ 33°11′	86°15′ 87°36′	Marble Gravel	202/68 90/72	15. 1(2) 17. 93	41.7	. 9 . 27	1. 6 . 68	
				Arizona					
Maricopa, (C-1-4) 6bba	32°09′ 33°31′ 33°32′ 33°50′ 33°03′ 33°04′	112°42′ 111°15′ 114°44′ 113°13′ 114°26′ 112°53′ 113°17′ 114°30′	Alluvium do Sand, gravel, and silt Alluvium Sand, silt, and gravel Alluvium do Sand, gravel, and silt	712 560 1, 692 584/120 600 274	83. 86 (-, 24) 324. 65 88. 48 (-, 12) 297. 35 (-, 51) 11. 70 35. 68 (+, 08) 131. 45 (-, 035)	0. 20 . 36 . 07 . 00 . 28	0. 22 . 36 . 25 . 51 . 25	0. 42 . 72 . 32 . 51 . 53 >1 . 237 1. 14	0.3 in. per day, 1:2. No rise; drop in level only 0.3 in. per day, 1:1. No aftershocks recorded. 0.3 in. per day, 1:1. Clock stopped 3 days be- fore quake; 0.3 in. per
(C-9-25) 35bab (C-10-24) 15cdd	32°36′ 32°33′	114°48′ 114°43′	do	1, 190 202	17. 14	. 13	. 17	. 30 . 73	day, 1:2. 0.3 in. per day, 1:2. Clock stopped 5 days be- fore quake: 0.3 in. per
(C-10-25) 31bbb (C-11-24) 23bcb	32°31′ 32°27′	114°40′ 114°42′	Sand and gravel Silt, sand, and gravel	286 1, 038	80. 31 77. 98	. 10	. 13	>1 23	day, 1:2. 0.3 in. per day, 1:1. Aftershocks recorded (see table 6); 0.3 in. per day 1:1.
				Arkansa	3				
Craighead, 13N-2E-35daa1	35°42′	90°50′	Quaternary sand	120	55. 97	0.72	0. 77	1. 49	No aftershocks recorded; 2.4 in, per day, 1:6.
Dallas, 10S-13W-34aca Desha, 11S-2W-3cca1	33°48′ 33°46′	92°25′ 91°17′	Sparta Sand Cockfield Formation	888/836 754	114. 10 21. 86	1. 70	1. 60	3. 30 >1	Do. No aftershocks recorded; 2.4 in. per day, 1:1.2.
Drew, 158-4W-12dda1 Lincoln, 7S-5W-17ccc1	33°24′ 34°06′	91°28′ 91°37′	Sparta SandQuaternary sand	760 120/110	32. 32 17. 56	. 95	1, 15	>1 2. 10	Do. No aftershocks recorded; 2.4 in. per day, 1:6.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water (feet)	From pree	earthquake vel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
				California	a.				
Fresno, 15S/16E-20R1	36°36′	120°14′	Alluvium: upper aquifer	1, 250/490	71. 10	0. 71	0. 60	1. 31	
17S/17E-21N2	36°26′	120°08′	zone. Alluvium: upper and lower aquifer zones.	1, 005/404	313. 75	. 07	. 07	. 14	
198/17E-35N1 [mperial, 8 (14S/11E-32R)	36°13½′ 32°54′	120°26′ 115°52′	AlluviumAlluvial silt, sand, and	2, 030/608	494. 08(09) 120. 02	. 07	. 24 . 12	. 31 . 24	0.3 in. per day, 1:1.
11 (15S/18E-15M)	32°35′	115°05′	gravel. do		25.30(21)	. 03	. 27	. 30	Took 2 weeks for water level to recover to stay indicated by prequake
12 (16S/19E-11D) 2 (16S/19E-32G1)	32°30′ 32°28′	114°59′ 115°01½′	Deltaic alluvial deposits Alluvium	252	13. 04 34. 00	. 46	. 44 . 12	. 90 . 20	trend. 0.3 in. per day, 1:2. Water-level trend reverse coincidentally at time of quake; 0.3 in. per day, 1:2.
Kern, 25S/26E-1A2 32S/28E-20Q1	35°47′ 35°07′	119°07′ 118°59′	Tertiary basalt	875/200 950	352, 56 212, 10(-, 50)	. 07	. 04 . 64	. 11	, , , ,
8N/10W-8R3	35°03′ 34°48′	117°47′ 117°57′ 117°58′	Alluvium	_ 230	161. 5	. 38	. 37	. 67 . 75 . 15	0.3 in. per day, 1:1.
1S/9W-3B1	34°35′ 34°07′ 34°01′	117°49′ 118°03′	do do	408	232. 57 103. 92 20. 50(10)	. 09 . 06 . 20	. 08 . 02 . 22	. 17 . 08 . 42	2.4 in. per day, 1:6. 1.0 in. per day, 1:1. 1.0 in. per day, 1:5.
Kern, 25S/26E-1A2	34°00′ 33°52′ 33°52′	118°07′ 117°50′ 117°51′	Alluvium Alluvium (La Habra	552	109. 30(06) 218. 48 158. 06	. 50 . 50 . 08 . 21	. 42 . 07 . 17	. 92	Do. 2.3 in. per day, 1:10.
4S/10W-1C2		117°53′	Formation). Pleistocene sand and	400 514	102. 05	. 31	. 24	. 38	Do.
4S/9W-30E1 5S/11W-16D2	33°48′ 33°45′	117°52′ 118°02′	gravel.	235	64. 35 (-) 2. 60	. 53	. 36 . 20	. 89	Do. 1.2 in. per day, 1:5.
6S/11W-1Q9	33°40′	117°59′	Alluvium Pleistocene sand and gravel.		?	. 02	. 38	. 40	D
6S/11W-12G1 6S/10W-7E2		117°59′ 117°58′	do		2.73	. 05	. 63	. 68	Pressure recorder; 51° per day, 1:7.
6S/10W-7E4	33°40′	117°58′	do	120/90	4.70	. 05	. 37	. 42	Pressure records from tw depths in adjacent piezometers; 51° per day, 1:12.
6S/11W-1Q4 6S/11W-1Q6 6S/11W-1Q3	33°40′ 33°40′ 33°40′	117°59′ 117°59′ 117°59′	dodododo	360/300		. 02 . 02 . 10	. 25 . 21 . 35	. 27 . 23 . 45	Do. Pressure recorder; 51° per
6S/11W-1Q9		117°59′	do		(+)3.8	. 02	. 38	. 40	day, 1:28. Pressure recorder; 51° per
Riverside, 7S/22E-17P San Bernardino,11/26E-31C1.	33°34′	114°43′	Gravel.	265	6.00(05)	. 07	. 12	. 19	day, 1:14. 0.3 in. per day, 1:1. 0.3 in. per day, 1:2.
San Bernardino,11/26E-31C1. Santa Barbara, 6/32-11G3 7/35-22N2	35°00′ 34°37′ 34°40′	114°38′ 120°14′ 120°34′	Alluviumdo	143 28/25 194	6. 50 7. 349 7. 24(. 60 . 002 . 18	. 64 . 010 . 31	1. 24 . 012 . 49	0.3 in. per day, 1:2. 0.3 in. per day, 1:1. 0.3 in. per day, 1:5.
7/35-28 R1 7/35-33 R1	3403016/	120°33½′ 120°34′ 120°34′	Sandy gravel Careaga Sand Gravel of Careaga Sand	551 420	7. 24(02) 60. 328(005) 112. 157(004) 113. 21	.017	. 018	. 035	0.3 in. per day, 1:1. Do.
10/33-7R1	34°38½′ 34°57′	120°23′	Alluvium	_ 210			. 07	. 13	$T=2\times10^5$; $S=0.15-0.30$; 0.3 in. per day, 1:5.
Santa Clara, 7S/1E-9D2 7S/1E-16C5	37°20½′ 37°19½′	121°52½′ 121°52′	do	917	135, 47 164, 36(-, 15)	1. 07	1. 32 1. 05	2. 39 1. 68	2.4 in. per day, 1:12.
Solano, 8N/1W-33A1 Tulare, 23S/25E-16N3 23S/25E-16N4	35°55½′ 35°55¼′	119°17′ 119°17′	do do	200/20 430 250	94. 2 173. 99 95. 88	. 20 . 23 . 027	. 24 . 18 . 022	. 44 . 41 . 049	2.4 m. per day, 1.12.
23S/25E-17Q3 Yolo, 8N/1E-17F1	35°55½′ 35°55½′	119°17½′	do	355 200/20	100. 19 67. 23	. 32	. 36	. 68	2.4 in. per day, 1:6.
				Colorado		, ,			
Prowers, C23-42-13cdab	38°03′	102°06′	Alluvium	48	7. 45	0. 12	0, 15	0, 27	2.4 in. per day, 1:6.
				Connecticu	ut	'			
			No w	ells recorded	the quake.				
				Delaware	•				
			1	No report rece	eived.				
				Florida					
Bay, 006-536-423	30°06′	85°36′	Limestone of Floridan aquifer.		40. 10	0.40	0.38	0. 78	1.2 in. per day, 1:6.
012-550-331 012-541-213	30°12′ 30°12′	85°50′ 85°41′	do		25. 60 6. 13	. 75 . 85	. 82 . 73	1. 57 1. 58	1.2 in. per day, 1:12. 0.3 in. per day, 1:5.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water (feet)	From pre-	earthquake vel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
			F	lorida—Cont	inued				
Broward, F291	26°00′	80°08′	Limestone of Biscayne	107	(+)1.17	2.04	2. 47	4. 51	0.3 in. per day, 1:2; float hung after quake.
G561	26°05′ 26°11′ 26°18′ 26°06′	80°08′ 80°20′ 80°09′ 80°05′ 80°12′ 82°02′	aquifer. do. do. do. do. do. do.	215	(+). 54 (+)3.6 (-). 88 (+)3.34(01) (+). 17 56. 61	. 03 . 12 2. 87 . 03 1. 91	. 03 . 11 2. 89 . 01 1. 91	. 06 . 23 5. 76 . 04 3. 82 >5	0.3 in. per day, 1:5. Do. Tape thrown off pulley. 1.2 in. per day, 1:6. 1.2 in. per day, 1:6. No aftershocks; water level declined at time of quake and continued as
Collier, C131	26°06′ 26°10′	81°16′ 81°15′ 81°41′ 81°42′ 82°	Tamiami FormationNonartesiandodotosionstone of Floridan	60/9 60/12 60/12 836/680	23. 98 (+)4. 30 (+)5. 00(39) (+)7. 96 91. 82	1. 30 . 50 . 19 . 82 . 42	1. 28 . 56 . 44 . 93 . 32	2. 58 1. 06 . 63 1. 75 . 74	fi aquifer was being drained to a level 4 ft lower. 1.2 in. per day, 1:6. Do. Do. Do. 1.2 in. per day, 1:12.
Dade, F45	. 25°50′	80°12′ 80°20′ 80°15′ 80°28′	aquiferdododoLimestone of Biscayne aquifer.	53	(+)1, 60 (+)1, 35 (+)1, 51 (+), 93	1. 24 1. 16 1. 51 0. 17	1. 47 1. 32 1. 66 0. 15	2. 71 2. 48 3. 17 0. 32	2.4 in. per day, 1:6. 0.3 in. per day, 1:6. 0.3 in. per day, 1:5.
G72	_ 25°57′	80°25′	aquifer. Oolitic limestone of Biscayne aquifer. Limestone of Biscayne aquifer.	4. 6	(+)4.25	. 33	. 36	. 69	0.3 in. per day, 1:5.
G476		80°18′ 80°20′	Limestone of Biscayne aquifer.	24/19 91/36	(+)1.13	. 48 1. 10	. 39 1, 19	. 87 2, 29	1.2 in. per day, 1:6. Do.
G695. G613. G614. G617. G618. G619. G620. G858. G860. G861. G863. G864. G968. G973. G974. G975.	25°24′ 25°32′ 26°05′ 26°05′ 25°45′ 25°46′ 25°39′ 25°39′ 25°33′ 25°33′ 25°52′ 25°52′ 25°52′ 25°52′	80°02' 80°32' 80°36' 80°26' 80°26' 80°46' 80°46' 80°24' 80°34' 80°34' 80°26' 80°24' 80°24' 80°27' 80°27'	do	14/11 21/18 20/18 20/11 13/7 16/6 50 15/10	(+)2. 21 (-). 60 (+). 33 (+)2. 02(01) (+)3. 60 (+)6. 12(01) (+)6. 64 (+)4. 80 (+)3. 08(02) (+)1. 38 (+)2. 45 (+). 74(02) (+)4. 70 (+)2. 60 (+)3. 85(01) (+)5. 60(02) (+)4. 80	. 02 . 08 . 04 . 63 . 31 . 36 . 38 . 05 . 09 . 10 . 06 . 22 . 19 . 06 . 20 . 06	. 02 . 15 . 07 . 54 . 35 . 42 . 50 . 13 . 10 . 21 . 08 . 24 . 18 . 07 . 00 . 08	. 04 . 23 . 11 1. 17 . 66 . 78 . 88 . 18 . 19 . 31 . 14 . 46 . 37 . 13 . 20 . 14	0.3 in. per day, 1:5. Do. Do. 1.2 in. per day, 1:2. 1.2 in. per day, 1:6. 0.3 in. per day, 1:6. 0.3 in. per day, 1:6. 0.3 in. per day, 1:6. Do. Do. Do. 1.2 in. per day, 1:6. Do. 1.2 in. per day, 1:6. Do. 1.2 in. per day, 1:12. Do. Do. Do.
G1183 NP44 NP46 NP57 NP62 NP67 NP62	25°29′ 25°24′ 25°19′ 25°19′ 25°24′ 25°19′	80°23′ 80°42′ 80°47′ 80°31′ 80°45′ 80°39′	do	33 25 54/8 20 20	(+). 75 (+). 66 (+. 04) (-). 47 (+). 07 (02) 6. 13 (+)0. 68 4. 56 (01)	. 35 1, 22 . 40 . 51 . 13	. 33 . 07 . 47 . 43 . 14	. 68 1. 29 . 87 . 94 . 27	0.3 in. per day, 1:5. 1.2 in. per day, 1:3. Do. 0.3 in. per day, 1:2. 0.3 in. per day, 1:5. 0.3 in. per day, 1:2; tape thrown off pulley. Do.
S18 S19	25°55′ 25°48′ 25°48′	80°46′ 80°17′ 80°17′	do	52 95/91 64/51	(+)1.98 (-).20 (-)1.93(04)	. 35 2. 40 1. 47	. 23 2. 89 1. 07	. 16 . 58 5. 29 2. 54	No aftershocks recorded; 0.3 in. per day, 1:5. 1.2 in. per day, 1:6. Do.
S182 S196 Desoto, 704-147-332		80°21′ 80°29′ 81°47′	Limestone of Hawthorn		(+)1,71 (+)1,23(-,02) (+)4,91(+,33)	. 06 . 04 . 90	. 04 . 10 . 14	. 10 . 14 1. 04	1.2 in. per day, 1:12.
Duval, 206 Gulf, 30 (948-518-1) Hardee, 731-145-1 734-202-332	. 29°48′	81°45′ 85°18′ 81°45′ 82°02′	Formation. Floridan aquiferdodo Floridan aquifer (lime-	_ 563/300	14, 85 7, 18(-, 02) 25, 40(-, 40) 65, 38(-, 29)	. 15 1. 50 . 18 . 81	. 12 1. 50 . 52 1. 19	. 27 3. 00 . 70 2. 00	1.2 in. per day, 1:6. 1.2 in. per day, 1:12.
738–151–223 Hillsborough, 801–213–213a_	27°38′ 28°01′	81°51′ 82°13′	stone and dolomite)do Limestone of Floridan aquifer.	737/50 417/67	43 . 02 1. 77	1. 60 ≥1. 63	1. 86 ≥1. 73	3. 46 ≥3. 36	Pen dislodged by quake. Cable thrown off; lost float and counterweight 1.2 in per day, 1:6.
803-234-313 803-238-212 805-235-113 805-236-333 805-238-100 807-230-133 807-230-421 13 (807-230-433).	28°05′ 28°05′ 28°05′ 28°07′ 28°07′ 28°07′	82°34′ 82°38′ 82°35′ 82°36′ 82°38′ 82°30′ 82°30′ 82°30′	do	807/710 1, 200/656 1, 200/697 1, 117/605 300/141 1, 250/720 362/70	(+)1.73 (+)2.70 12.53 (+)4.89 (+).32 ? 17.87	?	. 91	>5 	Cable thrown off pulley. Do. Do. Do. Do. No aftershocks recorded. Cable thrown off pulley.
809-232-414	_ 28°09′	82°32′	do		14. 85	2. 62	2.6E	5. 2E	Water level declined 0.2 ft in 3 hrs after quake recorded, then rose 2.15 ft during next 63 hrs. 1. in. per day, 1:6.
Lake, 832-154-334 Lee, L-246 L-414	_ 26°38′	81°54′ 81°49′ 81°49′	Tamiami Formation	_ 27/19	2. 32 (+)16. 43 (+)15. 23	1. 17 . 18 . 60	1. 13 . 20 . 60	2. 30 . 38 1. 20	1.2 in. per day, 1:6. 1.2 in. per day, 1:12.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to screen, per-	Depth to water (feet)		earthquake vel	Double	Remarks
				forated casing, or open hole (feet)		Upward	Down- ward	amplitude	
			F	orida—Conti	nued				
Leon, 7(027–416–1)	30°27′	84°16′	Limestone of Floridan aquifer.	314/165	159, 31	4.6E	4, 57	9. 2E	1.2 in. per day, 1:12.
Madison, 18(028-325-1) Pasco, 13 (815-226-1) 821-217-221	30°28′ 28°15′ 28°21′	83°25′ 82°26′ 82°17′	do	322/307 49/43 699/205	22. 25 5. 76 108. 52(10)	2. 10	2. 58	4. 68 1. 53 >5	1.2 in. per day, 1:6. 6 hours after quake re- corded, water level began to rise and rose 0.94 ft in 12 hrs.
826-211-214 Pinellas, 561 (750-240-1) 665 (758-244-4) 246 (758-247-1)	28°26′ 27°50′ 27°58′ 27°58′	82°11′ 82°40′ 82°44′ 82°47′	do dodo	227/49 188 299/81 208	16. 29 (04) 3. 40 21. 10 26. 10	1. 26 . 73 1. 18 1. 80	1. 02 . 85 2. 01 1. 70	2. 28 1. 58 3. 19 3. 50	1.2 in, per day, 1:12. 1.2 in, per day, 1:6.
246 (758-247-1) 667 (759-243-313) 77 (804-245-1)	27°59′ 28°04′	82°43′ 82°45′	do	845 282	54. 32 65. 58	1. 09	1. 7E . 26	2. 8E . 53	1.2 in. per day, 1:12.
13 (808-245-1) Polk, 753-158-311 810-144-1	28°08′ 27°53′ 28°10′	82°45′ 81°58′ 81°44′	do do do	141/33 710/237 425/102	9. 48 26. 41(+. 06) 8. 72(08)	1. 90 3. 88 2. 15	1. 91 3. 70 1. 48	3. 81 7. 58 3. 63	Do. Do. Water level began to de- cline after quake recorder and fell 1.7 ft over 17
Sarasota, 9(719–225–1)	27°19′	82°25′	Limestone of Floridan aquifer.	730/101	3. 92	. 51	2. 30	2. 81	days. 1.2 in. per day,1:6 Upward motion blocked (?). Water level rose 1.7 ft during 5 days after quake recorded: 1.2 in.
Seminole, 125(841-122-1)	28°41′	81°22′	do	158/74	38. 50	. 80	. 71	1, 51	per day, 1:12. Water level rose 0.6 ft in 48 hrs after quake; 1.2
Sumter, 821–202–3	28°21′	82°02′	do	143/20	4, 59(-, 04)	. 33	. 16	. 49	in. per day, 1:21. Water level began to rise 6 hrs after quake and 10se 0.65 ft in 20 hrs: 1.2
Γaylor, 35 (003-330-1)	30°03′	83°30′	do	245/189	20. 2(+. 8)	8. 5E	8. 5	1.7E	in. per day, 1:6. Only well in Florida to record aftershocks; for list, see table 6; 1.2 in. per day, 1:24.
36 (003-331-1) Volusia, 31 (856-105-1) 905-113-3 909-106-4	30°03′ 28°56′ 29°05′ 29°06′	83°31′ 81°05′ 81°13′ 81°06′	Sand Floridan aquifer do do	35 113 351/93 220/152	6. 52 4. 96 . 16 5. 74(04)	1. 13 . 78 2. 7E 2. 7E	1. 41 . 65 2. 74 2. 66	2. 54 1. 43 5. 4E 5. 4E	Water level rose 0.25 ft in 18 hrs after quake; 1.2
910-105-1	29°05′	81°05′	do	234/102	14. 90	4.86	4. 36	9. 22	in, per day, 1:6. Water level rose 0.5 ft in 6 hrs after quake; 1.2 in per day, 1:12.
				Georgia				,	
Chatham, 63 (37Q7)	32°05′ 32°04′	81°06′ 81°04′	Ocala Limestone	525/120	112, 50	2. 88	3. 30	6. 18 >10	No aftershocks recorded.
317 (38Q2)	32°00′ 32°02′	81°50′ 80°54′	do do	500/260 386 354/110	79. 25(05) 24. 10	5, 30	4, 45	$>_{9.75}^{10}$	Do. Do.
382 (36Q8) 429 (37Q34) Dawson, 12-3	32°05′ 32°00′ 32°20′	81°08′ 81°05′ 84′05′	do	413/250 327 400/79	101. 34(07) 74. 72(47) 22. 04			>5 3, 50 >1	Do. Do. Aftershocks recorded (see table 6); 2.4 in. per day 1:12.
Dougherty, 133-400-4	31°33′	84°00′	Ocala Limestone	243/206	27. 95	3. 88	?	7. 76E	Aftershocks recorded (see table 6); 2.4 in. per day
135-406-3	31°36′	84°06′	Clayton Formation	760/713	61. 25	. 20	. 15	. 35	1:6. First quake ever recorded in well.
Effingham, 7 (34R36)Fulton, 26	32°09′ 33°42′	81°23′ 84°26′	Ocala Limestone Injection complex	431/273 350	17. 00(-, 02) 13. 42	. 002	. 021	>2 . 023	First quake ever recorded in this watertable well.
Glynn, E143	31°10′	81°30′	Ocala Limestone and upper part of Claiborne Group,	950/823	(+)3.0		*********	>5	T=1.4×10°; S=2×10-4; no aftershocks recorded 1.2 in. per day, 1:5.
J35	31°08′	81°29′	Ocala Limestone	710/611	(+)12.7	3. 3	2. 7	6. 0	$T=1.6\times10^6$; $S=3\times10^{-4}$. Pressure recorder.
J36	31°07½′	81°29′	Ocala Limestone and upper part of Clairbone Group.	1, 007/589	(+)15.3	3. 6	3. 3	6. 9	T=2.7×10 ⁶ ; S=4×10 ⁻⁴ ; no aftershocks recorded Pressure recorder.
J67 Laurens, 21T1		81°25′	Ocala Limestone		(+)23.7	3. 0	2, 6	5. 6 1. 97	T=106; S=3×10-4; Pressure recorder.
Laurens, 2111 Lowndes, 19E2	32°27′ 30°50′	83°04′ 83°17′	Suwannee Limestone	342/200	25. 04(06) 115. 96(+. 02)	. 98	. 99 . 16	. 52	First quake ever recorded in well.
Miller, 8H2	31°10′	84°44′	Clayton Formation	1, 040/776	31.10(+.30)	1. 48	. 00	1. 48	Water level rose 1.48 ft in 20 min, then declined 1.18 ft in 1 hr; 1.2 in. pe day, 1:6.
Mitchell, 10G313	31°05′	84°26′	Ocala Limestone		43.00			>5	Aftershocks recorded (see table 6); 1.2 in. per day

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	evel fluctuat	tion (feet)	
County, well	Lati- tude; N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to screen, per-	Depth to water (feet)	From pre	earthquake evel	Double	Remarks
				forated casing, or open hole (feet)		Upward	Down- ward	amplitude	
			G	eorgia—Cont	inued				
Γhomas, 14Ε15	30°50′	83°58′	Ocala Limestone	548	195. 20			>5	Cable thrown off pulley
14E20	30°51′	80°57′	Suwannee Limestone	299/147	223.19(+.04)	1. 29	1. 26	2. 55	and float stuck. First quake ever recorded in well.
15E12 Fift, 17K1 Vayne, 30L3	30°47′ 31°27′ 31°37′	83°51′ 83°31′ 81°55′	Ocala Limestone_ Tampa and Suwannee Limestones.	183/136 312 594	(+. 084) 123. 50(40) 65. 18	. 20 1. 15	. 52 1. 22	>1 . 72 2. 37	Do. 1.2 in. per day, 1:6. T=2.5×10 ⁶ ; 2.4 in. per day, 1:6.
				Hawaii		1			
Oahu, 1A	21°16′	157°46′	Basalt of Kolau Volcanic	131/100	8. 27			>1	
2 83	21°17′ 21°18′	157°48′ 157°51′	Series. dodo	?/100 474/458	25. 15 (+)25. 84(+. 06)			>1	Aftershocks recorded (see
				11 1, 100	(1,20.01(1.00)			1100	table 6); 1.2 in. per day, 1:1. Tidal efficiency=10
286	21°35′	158°11′	Basalt of Wainae Volcanic Series.	447/447	+6.87	0. 03	0. 05	. 08	percent.
332	21°35′	158°07′	Basalt of Koolau Volcanic Series.	225/205	+2.98	. 28	. 32	. 60	Do.
333	21°35′	158°06′	do	163/68	9, 14	. 09	. 05	. 14	Tidal efficiency=15 percent,
T-24. T-28. T-45. T-52.	21°21′ 21°35′ 21°22′ 21°24′	157°53′ 158°06′ 157°55′ 157°56′	do do do	115/66 60/39 85/59 321/170	21. 95 (+. 10) 23. 28 19. 12 (+. 07) 19. 35	. 21 . 10 ?	. 12 . 17 ?	. 33 . 27 . 16	2.3 in. per day, 1:1. Tidal efficiency = 4 percen 2.3 in. per day, 1:1. Do.
T-57	21°36′	158°06′	Basalt of Koolau Vol-	33/11	2. 03	. 75	. 83	1. 58	Tidal efficiency=15 per-
T-67	21°23′	157°57′	canic Series.	1, 308/91	4. 17	. 2	. 03	. 05	cent. Tidal efficiency=2 per-
T-69	21°20′	157°52′	do	283/233	24. 55(?)	. 31	. 33	. 64	cent. Chart pulled from drum by pen; 1.2 in. per day,
T-75	21°23′	157°56′	do	250/75	18. 16(?)			>1	1:1. Rising float stuck betwee counterweight and cas-
T-96	21°35′	158°09′	Reef limestone	60/16	7. 00	. 35	. 60	. 95	ing; 1.2 in per day, 1:1. Tidal efficiency=30 per- cent.
Shaft 4	21°29′	158°01′	Basalt of Koolau Vol- canic Series.					Trace	001101
Shaft 17	21°35′	158°05′	do		237, 60	. 05	. 05	. 10	Maui-type shaft, 608 ft. long.
Waibee Tunnel (1,624 ft long; 24 ft satu- rated and held by dike and bulkhead)	21°27′	157°51′	do					4. 60	Pressure change on horizontal discharge line near the bulkhead.
				Idaho					
Bingham, 5S-31E-23ab1	42°58′	112°49′	Basalt of Snake River	46	24, 70	0.04	0. 05	0, 09	0.3 in. per day, 1:2.
Blaine; 1S-19E-3cc2	43°21′	114°12′	Group. Clay, silt, sand, and	51/51	17, 97	. 81	. 75	1. 56	Do.
2S-20E-1ac2	43°17′	114°01′	gravel. Basalt of Snake River	208/208	151. 24	. 56	. 47	1.03	Do.
Butte, 3N-29E-14ad1	43°35′	112°58′	Group and alluvium. Basalt of Snake River	588	459, 02	2, 27	1. 71	3. 98	2.4 in per day, 1:6.
7N-31E-34bd1 Canyon, 2N-1W-7bb4	43°55′	112°43′	Group. Basalt of Idaho Group. Balagraia limestone	320	269. 44	. 04	. 05	. 09	
Canyon, 2N-1W-7bb4 Cassia, 13S-21E-18bb1	43°32′ 42°18′	116°31′ 114°03′	Basalt of Idaho Group Paleozoic limestone	103/96 850/20	11. 28 430. 92	. 05	. 05	. 10 1. 42	Aftershocks recorded (see table 6); 0.3 in. per day 1:2.
Elmore, 1S-4E-10da1	43°21′	115°57′	Sand and gravel of Idaho	525/485	341. 68	. 04	. 06	. 10	0.3 in. per day, 1:1.
Gooding, 8S-14E-16bc1	42°44′	114°50′	Group. Basalt of Snake River Group.	53/50	39. 55	. 02	. 02	. 04	2.4 in. per day, 1:6.
Jefferson, 5N-32E-36ad1 7N-36E-22ab4	43°43′ 43°54′	112°38′ 112°07′	do	406/361 35/18	330. 07 7. 09	. 71	. 50	1, 21 , 73	Do.
Jerome, 7S-17E-6ac1 Latah, 39N-4W-7	42°51′ 46°45′	114°30′ 117°	Columbia River Basalt	345/322	314. 53(- 02) 255. 1	. 02	. 02	>5	0.3 in. per day, 1:1. 0.3 in. per day, 1:2. Aftershocks recorded (see table 6); 1.2 in. per day
Lincoln, 5S-17E-26ac1	42°58′	114°24′	Basalt of Snake River	255	202. 20	. 44	. 42	. 86	1:5. 0.3 in. per day, 1:2.
Minidoka, 7S-25E-19ba1	42°48′	113°35′	Group.	284/284	244. 38(02)	. 82	. 96	1. 78	Do.
8S-23E-2ba1 8S-24E-20db1	42°46′ 42°43′	113°44′ 113°40′	do	254/80 367/25	208. 65 154. 10	. 62	1. 14	1. 76 . 30	Do. 2.4 in. per day, 1:6. 0.3 in. per day, 1:2.
8S-24E-31dc1 8S-25E-24bd1	42°41′ 42°43′	113°42′ 113°29′	do	213/175 180/160	153. 53(04) 145. 13(02)	. 09	. 16	. 25 . 20	Do.
Power, 5S-33E-35cc1	42°56′	112°34′	Gravel	60	25. 29	. 04			Float hung on down movement; 0.3 in. per

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to screen, per-	Depth to water (feet)	From pred le	earthquake vel	Double	Remarks
				forated casing, or open hole (feet)		Upward	Down- ward	amplitude	
				Idaho-Conti	inued				
Power, 7S-30E-28bb1	42°47′	112°58′	Basalt of Snake River Group.	288	198. 29	0.00	0.12	0. 12	0.3 in. per day, 1:2.
Teton, 4N-45E-13ad1 Twin Falls, 11S-19E-17aa1	43°40′ 42°29′	113°05′ 114°15′	Alluvium Basalt of Snake River Group,	304 860	201. 46 321. 92	. 22	. 28	>5	1.2 in. per day, 1:6. 0.3 in. per day, 1:2.
11S-20E-21dc1	42°27′	114°07′	do	280	69. 85	. 24	. 32	. 56	Do.
		[Da	ata furnished by the Illinois	Illinois Water Surve	v and the Northern	ı Illinois G	as Co.l		
Champaign, CHM 19N9E-	40°07′	88°12′	Glacial sand and gravel	1	113, 77	0.03	0.05	0, 08	
8.7h. Cook, COK 39N12E-11.7f	41°53′	87°50′	Cambrian and Ordovi-	1,640	553. 0	0.03	0.05	>10	
DeKalb, DEK 40N3E-	41°56′	88°51′	cian sandstone.		130.87(+.51)	. 51	. 30	.81	
23.8e1. DuPage, DUP 39N11E-	41°51′	87°55′	Silurian dolomite	1	40. 90	. 82	1. 32	2. 14	
24.2g. DUP 38N10E-	41°47′	88°05′	do		13. 08	. 01	.00	. 01	
10.7a1. DUP 38N10E-	41°45′	88°05′	do	-	47. 20	. 40	. 33	. 73	
27.6h. ANL-9		88°	Niagara Dolomite			. 40	. 00	>1	1.2 in. per day, 1:1.
ANL-10		88°	do	199	96. 09 (+. 03) 81. 12 (+. 21)	4. 30	3. 40	7.70	Aftershocks recorded (see table 6); 1.2 in. per day, 1:10.
LaSalle, Wealdon 9	. 89°	41°	Troy Grove Gas Storage Field.		89. 4	. 45	. 45	. 9	Quake recorded at bottom of water level "low": 1.2
Weldon 15	. 89°	41°	do		90. 36	. 00	. 00	. 00	in, per day, 1:10. Water level changed trend and rose 1.0 ft in 13 hrs after quake; 1.2 in, per
Amfahr 3 Roulston 3		41° 41°	do		90. 37 80. 4	.8	.8	1. 6	day, 1:10. 1.2 in. per day, 1:10. Water level began drop-
					00. 1				ping when quake hit, fe 1.4 ft in 2 hrs, then re- versed and rose 3.6 ft in 60 hrs; 1.2 in. per day,
Fordyce 2 (or 3?)	. 89°	41°	Ancona Gas Storage		144. 4	. 3	. 3	. 6	1:10. 1.2 in. per day, 1:10.
Scheuer 2	. 89°	41°	Field. Garfield Gas Storage Field.		164. 7	. 7?	2, 1?	2. 8?	Float hung during quake; 1.2 in. per day, 1:10.
Fehr 2 Peoria, PEO 8N8E-6.1e	89° 40°42′	41° 89°37′	Glacial and and gravels.	163	142. 4(+. 2) 69. 0	. 2	. 2	. 4 1. 50	1.2 in. per day, 1:10.
PEO 8N8E-16.7g PEO 8N8E-17.2e2	40°40′ 40°40′	89°36′ 89°36′	do	_ 53	28. 82(+. 04) 33. 13(+. 02)	. 09	. 05	. 14	
Tazewell, TAZ 24N5W-3.8a. TAZ 26N4W-31.2g	40°33′	89°39′ 89°36′	do	- 80	37. 25(03) 3. 28(02)	. 02	. 13	. 15	
Will, WIL 37N10E-10.6f1	41°43′	88°04′	Silurian dolomite	93	73. 12	. 30	. 20	. 50	
				Indiana					
Allen, Al-4		84°53′ 84°50′	Limestone	44	30. 48 23. 71 (+. 01)	0. 32	0. 39	0.71	1.2 in. per day, 1:5. No aftershocks recorded;
Benton, Be-2	40°31′	87°23′	Gravel	. 37	12. 67	. 12	. 07	. 19	1.2 in. per day, 1:1. 0.3 in. per day, 1:1.
Clinton, Cl-4 Jasper, Jp-4	41°03′	86°30′ 87°01′	Limestone Ordovician rock	230	17. 58(+. 03) 4. 53(10)	. 09	. 02	. 11	1.2 in. per day, 1:1. 1.2 in. per day, 1:5. 0.3 in. per day, 1:5. 1.2 in. per day, 1:5. 1.2 in. per day, 1:10.
Jefferson, Jf-4 Madison, Md-8	. 40°16′	85°26′ 85°50′	Limestone	415	4. 53(10) 25. 71(08) 28. 90(10) 101. 70(+. 7)	. 06	. 08	. 14	1.2 in. per day, 1:5. 1.2 in. per day, 1:10.
Marion, Ma-31 Ma-32	39°51′ 39°52′	86°01′ 86°08′	Niagara Dolomitedo	347/210	101. 70(+. 7) 9. 92(14)	3, 65	4. 60	8. 25 >1	Aftershocks recorded (see table 6); 1.2 in. per day,
Marshall, Ml–2Ml–4.	41°21′ 41°27′	86°19′ 86°19′	Gravel Sand		22, 03	. 29	. 25	. 54	1:1. 0.3 in. per day, 1:5.
Miami, Mi-2 Newton, Ne-3	40°40′	86°08′ 87°27′	Limestone Sand and gravel	133 165/66	46. 24 44. 28(12)	. 03 1. 78	1.38	. 06 3. 16	0.3 in. per day, 1:10. 1.2 in. per day, 1:10.
Parke, Pa-3 Porter, Pt-9	39°48′	87°22′	Limestone	103	36. 57 48. 55	1. 14	. 80	1. 94	1.2 in. per day, 1:5. 0.3 in. per day, 1:5.
		87°13′			23, 14(-1, 92)		*********	>1	No aftershocks recorded; 0.3 in. per day, 1:1.
Posey, Py-2 Pulaski, Pu-6(29/4W-4L1)	40°59′	87°47′ 86°53′	Pennsylvanian rock Niagara Dolomite		11. 60 15. 93(?)	. 04	. 06	>1 10	1.2 in. per day, 1:5. Aftershocks recorded (see table 6); 1.2 in. per day,
Ripley, Ri-4 Spencer, 14	37°58′	85°06′ 87°08′	Sand and gravel	56/53	3. 59 8. 46	. 03	. 05	. 08	1:1. 1.2 in. per day, 1:1. 0.3 in. per day, 1:1.
Starke, Sk-2 Tippecanoe, Tc-7 Vanderburgh, Van-3	41°14′ 40°26′	86°37′ 86°55′	Gravel Sand and gravel	83 207	4, 04 168, 99	. 21	. 19	. 40	Do. 1.2 in. per day, 1:1.
vanderburgh, Van-3	37°59′	87°31′	Sand	90	24. 00	. 035	. 04	. 075	0.3 in. per day, 1:1.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati-	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water	From pree	earthquake vel		Remarks
		, , , , , ,		screen, per- forated casing, or open hole (feet)	(,	Upward	Down- ward	Double amplitude	
				Iowa				,	
Des Moines, 69-3-6A1			St. Peter Sandstone	1, 205/854	182.75(-1.75)	0.00	1. 75	1.75	Water level dropped at time of quake. No other water-level move ment; 2.4 in. per day,
Lee, 67–4–3J1 Linn, 83–7–28H1			Sand and gravel Silurian limestone and dolomite.	156 420/75	12.00 67.7(1)	. 1 . 15	. 1	. 2	4:5. 1.2 in. per day, 1:10. 2.4 in. per day, 1:12.
Cairo Dome of Natural Gas Pipeline Co. of America: Louisa, Jones G-1			Ordovocian shale and Silurian rock.	?/376	149. 9	. 00	. 16	. 16	Water-level decline only; 1.2 in. per day, 1:10.
Hutchinson 0-1			Galena Dolomite	?/571	133. 1	2.1(+?)	2. 6	4.7(+?)	Cable slipped on pulley at time of quake; 1.2 in. per day, 1:10.
Madison, 75-26-23A1 Marion, 74-21-11K1			Mississippian limestone	1, 058/657 113/76	262. 45 44. 30(-, 05)	. 55 . 025?	. 55 . 025?	1. 10 . 05?	2.4 in. per day, 1:3. Recorder made a jerky- type record; 2.4 in. per
Story, 83–24–2Q1 Washington, 77–924A1 Keota Reservoir of Natural Gas Pipeline Co. of			Sand and gravel	110 110/47	53. 85 3. 57	. 075	. 075 . 10	. 15 . 21	day, 1:6. 2.4 in. per day, 1:6. 0.3 in. per day, 1:6.
America: Washington, Anderson 1_			St. Peter Sandstone	?/1, 189	100. 23(+. 4)	. 75	. 41	1. 16	Float seemingly hung for 7½ hrs after quake; water level then began to rise (1.82 ft in 40 hrs.)
E.V.			do	?/1, 219	153. 9	. 18	. 18	. 36	1.2 in. per day, 1:10. 1.2 in. per day, 1:10.
Green 1. Flynn G-1. Woodbury, 89-47-22B2.			Galena Dolomite Dakota Sandstone	?/814 343	120. 0 22. 65(05)	. 25 . 61	. 46 . 55	. 71 1. 16	Do. 2.4 in. per day, 1:12.
				Kansas					
Kearny, 23-28-11db	38°	101°	Unconsolidated deposits	296	219. 53	0.18	0. 19	0. 37	2.4 in. per day, 1:6.
				Kentucky					
Christian, S-1,502.3-196.8 Edmondson, S-1,901.4-311.9	36°51′ 37°11′	87°27′ 86°05′	Ste. Genevieve Limestone	85 295	14. 09 129. 10(+. 13)	0. 06 1. 0	0. 07 . 8	0. 13 1. 8	No aftershocks recorded; 0.3 in. per day, 1:10.
Elliott, N-2,315.1-210.7	38°04′	87°09′	Rocks of Pennsylvanian age.	70	22. 60	. 06	. 06	. 12	0.3 in. per day, 1:5.
Graves, S-1,150.7-208.2 S-1,154.15-139.90 fefferson, N-1,540.0-258.6 N-1,566.00-275.35	36°52′ 36°40′ 38°11′ 38°14′	88°39′ 88°38′ 85°51′ 85°45′	Sanddo Glacial sand and gravel Limestone	106 183 112 190	16. 34 88. 17 60. 37 49. 46	. 06 . 23 . 09 . 141	. 04 . 18 . 09 . 165	. 10 . 41 . 18 . 306	Do. Do. 0.3 in. per day, 1:1. Do.
N-1,544.7-264.0 ohnson, S-2,864.6-536.6	38°12′ 37°46′	85°50′ 82°45′	Glacial sand and gravel Sandstone and shale of Breathitt Formation.	117 115	77. 39 26. 96	. 12	. 10	>1 22	No aftershocks recorded; 0.3 in. per day, 1:1.
Letcher, S-2,851.7-329.6	37°12′	82°49′	do	180	17. 0	. 53	. 60	1. 13	No aftershocks recorded; 0.3 in. per day, 1:10.
S-2,909.1-321.7 S-2,858.0-299.9 Livingston, S-1,276.1-259.3 S-1,276.6-347.9	37°10′ 37°06′ 37°01′ 37°15′	82°37′ 82°48′ .88°14′ 88°14′	dododo Warsaw Limestone. Fredonia Limestone, Member of Ste. Gene- vieve Limestone.	146 53 205 365	21. 87 11. 18 46. 77 22. 30	. 07 . 04 . 20 . 10	. 05 . 03 . 20 . 11	. 12 . 07 . 40 . 21	
yon, S-1,298.8-270.2 Marshall, S-1,246.2-272.0 AcCracken, S-1,119.0-310.2 Dhio, S-31,672.0-396.7 Pulaski, S-2,332.3-243,3	37°03′ 37°03′ 37°08′ 37°25′ 36°59′	88°09′ 88°20′ 88°46′ 86°52′ 84°36′	vieve Limestone. Warsaw Limestone. Gravel and sand. do. Tradewater Formation Limestone of Fort Payne	99 92 86 298 146	31. 84 26. 03 47. 65 152. 46 81. 08	. 10 . 01 . 30 . 01 . 09	. 10 . 01 . 26 . 02 . 04	. 20 . 02 . 56 . 03 . 13	0.3 in. per day, 1:5.
Warren, S-1,888.1-265.3	37°03′	86°08′	Formation. St. Louis Limestone	94	71. 51	. 09	. 09	. 18	Do.
				Louisiana					
cadia, AC-40	30°18′	92°25′	Chicot aquifer	303	53. 28(?)	1.07	0.94	2.01	Float hung after fluctua- tion; 2.4 in. per day,
ascension, An-2	30°14′	90°55′	Older alluvium	590/550	1.64			>2	1:12. No aftershocks recorded; pen hung; 0.3 in. per
Calcasieu, Cu-77_	30°14′	93°16′	"500-foot" sand, Chicot aquifer.	512/450	125. 55	. 20	. 20	. 40	day, 1:2. 2.4 in. per day, 1:6.
Cu-445 Cu-446	30°11′	93°19′ 93°19′	"700-foot" sand, Chicot	540/460 738/658	114. 40	. 78	. 80	1. 58 2. 00	0.3 in. per day, 1:5. Do.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	vel fluctuat	ion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water		earthquake evel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
			Lou	iisiana—Cont	inued				
Calcasieu, Cu-583	30°13′	93°17′	"700-foot" sand, Chicot	670/570	170.70(05)	0.24	0. 24	0.48	2.4 in. per day, 1:6.
East Baton Rouge, EB-78_ EB-90_ EB-123 EB-127 EB-128 .	30°30′ 30°28′ 30°26′ 30°26′ 30°26′	91°11' 91°09' 91°10' 91°11' 91°10'	aquifer. "400-foot" sand. "2,000-foot" sand. "600-foot" sand. "400-foot" sand. "800-foot" sand. "800-foot" sand.		115. 34(+.05) 145. 22(27) 90. 26(+.03) 29. 16 102. 00(14)	>. 33 >. 23 . 03 . 04	. 29 . 34 . 35 . 00	>. 62 >. 57 . 38 . 04 >5	2.4 in. per day, 1:6. 2.4 in. per day, 1:6. Do. Do. No aftershocks recorded; 0.3 in. per day, 1:5.
EB-155 _ EB-293 _ EB-652 _ East Feliciana, Ef-1	30°29′ 30°30′ 30°32′ 30°52′ 30°17′	91°10′ 91°11′ 91°08′ 91°01′ 91°14′	"400-foot" sand "600-foot" sand "1,500-foot" sand Quaternary upland Alluvium.	412/311 600/540 1,345/1,264 143 280/260	103. 50(+.06) 129. 50 64. 83 7. 45(+.008) 5. 30	. 40 1. 07 2. 77 . 06	.50 1.10 >1.14 .06	. 90 2. 17 >3. 91 . 12 >5	2.4 in. per day, 1:6. 2.4 in. per day, 1:30. 2.4 in. per day, 1:12. 2.4 in. per day, 1:1.2. Reversed on both sides of 1:5 chart; 2.4 in. per day
IB-92Jackson, Ja-49	30°07′ 32°17′	91°15′ 92°46′	do Sparta Sand	200/160 570	+1.20 158.35	1. 10 . 62	>. 20 1. 41	>1.30 2.03	1:6. 2.4 in. per day, 1:6. Record unusual but diffi- cult to interpret becaus of reversal; 2.4 in. per
Jefferson, Jf-120 Jefferson Davis, JD-485 Morehouse, Mo-5 Orleans, Or-42 Or-47 St. Charles, SC-9 SC-14 St. John The Baptist, SJB-17.	29°59′ 30°13′ 32°46′ 29°57′ 30°02′ 30°00′ 30°00′ 30°03′	90°09′ 92°59′ 91°55′ 90°02′ 90°04′ 90°24′ 90°24′ 90°27′	"700-foot" sand Chicot aquifer Sparta Sand "700-foot" sand do do "400-foot" sand do	610/527 777 404/324	78. 70 45. 50 198 106. 27(08) 100. 10 32. 13 69. 00 31. 95(+.07)	2. 24 . 10 1. 15 3. 10 1. 53 . 14 . 88	>1.30 .75 1.70 >.20 1.52 .12 .77	>3. 54 . 85 2. 85 >3. 30 3. 05 . 26 1. 65 >1	day, 1:6. 2.4 in. per day, 1:6. 2.4 in. per day, 1:12. 2.4 in. per day, 1:32. 2.4 in. per day, 1:36. Do. 0.3 in. per day, 1:5. Do. Aftershocks recorded (see table 6); 0.3 in. per day.
SJB-86 SJB-145 Union, Un-26 Vermilion, Ve-6 Ve-601 Vernon, V-104 Washington, Wa-7 Webster, Wb-27 West Baton Rouge, WBR-5 WBR-43	30°04′ 30°02′ 32°44′ 29°58′ 29°46′ 31°04′ 30°47′ 32°58′ 30°28′ 30°25′	90°29′ 90°39′ 92°09′ 92°08′ 92°20′ 93°13′ 89°51′ 93°27′ 91°12′ 91°13′	do Pleistocene Sparta Sand Chicot aquiferdo Miocene Pliocene(?) Sparta Sand "1,200-foot" sand Alluvium	320/305 745/670 214/125 249/167	31. 38(+.04) 11. 74(005) 159. 07(04) 15. 50 4. 20(+.02) 206. 25 11. 47 111. 43 93. 30 18. 16	. 20 . 68 . 17 . 12 . 27 . 08 . 68	1. 43 .58 .21 .01 .25 .08 2. 20	1. 63 >1 >1 1. 26 . 38 . 13 . 52 . 61 2. 86 >5	1:1. 0.3 in. per day, 1:2. 0.3 in. per day, 1:1. Do. 2.4 in. per day, 1:6. Do. Do. Do. Do. Do. Do. No affershocks recorded;
West Feliciana. WF-57	30°47′	91°23′	Zone 1 Tertiary	351/311	93. 22	. 23	. 44	. 67	0.3 in. per day, 1:5. 0.3 in. per day, 1:2.
				Maine					
Cumberland, C-26	43°54′	70°01′	Glacial sand and gravel	101/81	32. 37	0.08	0.11	0.19	Well drilled to bedrock. Barometric efficiency = 20 percent; 0.3 in. per day, 1:1.
				Maryland					
Charles, Ch-Cb7	38°34′	77°12′	Sand of Patapsco Forma-	400/154	68. 54	0.07	0.07	0.14	0.3 in per day, 1:5.
Dorchester, Dor-Cd40	38°34′	76°06′	tion. Sand of Piney Point Formation.	401/369	(20)	. 00	. 20	. 20	2.4 in. per day, 1:6.
Prince Georges, PG-Cf6	38°57′	76°44′	Sand of Magothy Formation.	207/?	+53.49(03)	. 10	. 14	. 24	0.3 in. per day, 1:1.
PG-Fd39	38°44′	76°50′	do	456/436	+39. 25	. 13	.15	. 28	2.4 in. per day, 1:6.
	1	1	L	Massachuse	tts	1	1		1
Berkshire, Lee-44	42°19′	73°14′	Stockbridge Limestone	49	9. 13	0. 31	0.31	0. 62	0.3 in. per day, 1:2.
				Michigar	1				
Bay, 17N 4E 22-1. Calhoun, 18 7W 32-3. 28 8W 2-1. Clinton, 5N 2W 31-1. Delta, 39N 3W 28-3. Eaton, 3N 3W 27-1. 4N 4W 11-1. 4N 4W 2-1. 4N 3W 12-1.	43°51′ 42°20′ 42°19′ 42°46′ 45°45′ 42°40′ 42°45′	83°59′ 85°09′ 85°12′ 84°36′ 87°09′ 84°38′ 84°45′	Saginaw Formation Marshall Formation do. Saginaw Formation Munising Sandstone Glacial drift Saginaw Formation	95/40 92/45 195 530 66/66 350	5. 35 25. 00(+.1) 15. 25 61. 22 2. 94 4. 04 251	0. 18 1. 27 . 71 . 16 1. 12 . 21 . 16	0. 18 1. 01 . 56 . 19 1. 10 . 15 . 11	0. 36 2. 28 1. 27 . 35 2. 22 . 36 . 27	0.3 in. per day, 1:5.
4N 4W 2-1 4N 3W 12-1 Genesee, 7N 7E 17-1	42°45′ 42°44′ 43°00′	84°45′ 84°37′ 83°40′	do do Saginaw Formation, Well bottoms in old coal mine.	376/23 381/140 222	29. 77 81. 88 25. 68	. 005	. 01	>2	0.3 in. per day, 1:1. Aftershocks recorded (see table 6); 0.3 in. per day, 1:2.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	evel fluctuat	tion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	ber, depth of well; second, depth of casing to	Depth to water	From pre	eearthquake evel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
			Mi	chigan—Con	linued				
Gogebic, 48N 47W 34-2		90°09′ 90°09′ 90°13′ 84°29′ 84°25′ 84°20′	Glacial driftdo	35/35 22/22 115/115 453/80 278/77 265	0. 71 (+. 02) 3. 04 22. 40 65. 10 6. 85 21. 525	0.00 .88 .60 .92	1. 35 . 74 . 58 . 82	1. 35 1. 62 1. 18 1. 74 >1	0.3 in. per day, 1:10. 0.3 in. per day, 1:1.
3N 2W 23-2 2N 1W 5-2 Ionia, 7N 7W 25-1 Jackson, 3S 1W 2-1 3S 1W 10-1	42°38′ 42°35′ 42°58′ 42°14′ 42°13′	84°30′ 84°26′ 85°05′ 84°23′ 84°25′	dodoGlacial driftMarshall FormationSaginaw and Marshall Formations.	268/50 210/37 23 221 323/55	7. 33 22. 50 16. 92 38. 50 31. 65	. 19 . 015 1. 67 . 73	. 185 . 00 2. 06 . 84	>1 . 375 . 015 3. 73 1. 57	Do. 0.3 in. per day, 1:10.
3S 1W-11-3 Kalamazoo, 2S 11W 20-11 3S 12W 11-1 Kent, 6N 12W 27-1 6N 12W 34-1	42°14′ 42°17′ 42°13′ 42°53′ 42°52′	84°23′ 85°37′ 85°41′ 85°43′ 85°42′	Glacial driftdododoMarshall Formationdo	36/33 81 248 265/207 300/150	12. 00 17. 20 (+) 0. 365 52. 65 68. 21	. 04 . 03 . 12 . 40	. 04 . 03 . 08 . 40	. 08 . 06 . 20 . 80 >5	Do. 0.3 in. per day, 1:1. 0.3 in. per day, 1:2. Aftershocks recorded (see table 6); 0.3 in. per day,
6N 9W 3-1 5N 12W 4-7 5N 12W 4-3 Livingston, 2N 4E 3-1 Mackinac, 42N 2W 7-1 Manistee, 21W 17N 14-1 Marquette, 47N 28W 1-1 47N 28W 3-1	42°56′ 42°50′ 42°50′ 42°36′ 46°03′ 44°14′ 46°30′ 46°30′	85°22′ 85°44′ 85°44′ 83°58′ 84°36′ 86°20′ 87°45′ 87°47′	Glacial driftdodo do Saginaw Formation Manistique Dolomite Glacial drift	70 227/182 86 148 102 212 216 75	17. 55 8. 63 11. 77(04) 12. 75 25. 70 33. 06 19. 09(01)	. 02 . 19 . 43 2. 27 2. 20 . 13 . 30	. 02 . 22 . 35 2. 27 2. 60 . 175 . 45	. 04 . 41 . 78 4. 54 4. 80 . 305 . 75 . 70	1:5. 0.3 in. per day, 1:1. Do.
Oakland, 3N 9E 36-1 3N 10E 13-2 3N 10E 31-1 3N 10E 32-1 3N 11E 4-1 Presque Isle, 33N 6E 15-1 Schoolcraft, 47N 16W 30-1 Van Buren, 48 16W 22-1 Washtenaw, 38 6E 16-3 38 7E 5-1 38 7E 9-3	42°38′ 42°40′	87°45′ 87°47′ 83°20′ 83°13′ 83°11′ 83°11′ 83°11′ 86°21′ 86°21′ 86°09′ 83°44′ 83°38′ 83°38′	do d	134 183/173 173/153 160/7 73 31/22 57/40 134/119 55/36 69 94/90	96. 78 86. 80 78. 85 79. 90 28. 60 6. 90 15. 45 27. 46 12. 50 3. 34 66. 34	. 03 1. 05 . 35 . 80 . 25 . 45 . 07 . 005 . 35 . 06	. 03 1. 05 . 45 1. 10 . 25 . 45 . 08 . 005 . 35 . 05	. 06 2. 10 . 80 1. 90 . 50 . 90 . 15 . 01 . 70 . 11	0.3 in. per day, 1:10. Do.
38 7E 24-6 Wayne, 1S 8E 17-1 Wexford, 21N 9W 4-1	42°12′	83°34′ 83°31′ 85°24′	do	75/70 114 277	33. 51 53. 90 25. 995	. 31	. 86 . 59 . 105	1. 17 1. 17 1. 17 . 285	0.3 in. per day, 1:1.
				Minnesota					
Dodge, 107.17.34dcc1	44°01′	92°50′	St. Peter Sandstone	500/118	88. 95	>0.5	>0.5	>1	Drum rotated more than once but no aftershocks
Grant, 129.42.9ccc1 Tennepin, 29.23.19cdd1 117.21.16cca 117.22.5abd2	45°59′ 44°58′ 44°56′ 44°58′	95°58′ 93°13′ 93°21′ 93°29′	Glacial drift Hinckley Sandstone Jordan Sandstone. Sandstone and limestone.	214/200 1, 016/925 421/280 483/201	77. 02 180. 13 76. 5(2) 45. 7	. 06 . 39 . 5 >1	.05 .12 .5	. 11 . 51 > 2	Drum rotated more than once; water level declined 2 ft in 40 hrs after quake
117.22.8dbd2 117.23.11bbd1	44°57′ 44°57′	93°29′ 93°33′	Jordan Sandstone	503/228 437/270	22. 26 18. 48(+. 32)	>. 79 >1	>. 44 >1	$\geq 1.23(?)$ ≥ 2	Do. Drum rotated more than once.
117.23.34daa2 tasca, 55.25.17acd1 Mower, 102.18.2bdd1 Nobles, 102.40.2rced1 St. Louis, 57.20.5dad1 57.20.31dbc1	44°53′ 47°14′ 43°40′ 43°36′ 47°26′ 47°22′	93°33′ 93°32′ 92°58′ 95°37′ 92°53′ 92°55′	Sandstone and limestone Glacial sand and gravel Limestone Glacial sand and gravel Biwabik Iron-formation Glacial outwash sand and gravel.	468/199 147/143 244 34/18 430/315 92/82	58. 30 32. 80 19. 05(+2. 35) 11. 33 65. 68(+. 40) 11. 26	>1 . 52 2. 35 . 09 . 62 . 02	>1 . 39 2. 05 . 09 . 00 . 02	>2 . 91 4. 40 . 18 . 62 . 04	Do. No aftershocks recorded.
58.18.12ccc1 Yellow Medicine, 114.45.4dcd1.	47°31′ 44°42′	92°34′ 96°17′	dodo	97/76 62/44	17. 24 (05) 8. 71	>. 26	>. 36	>. 62	Ink flowed too slowly to record fluctuation.
				Mississipp	i				
Forrest	31°19′	89°15′	Terrace sand and gravel	108/88	12. 90	0.7	0.6	1. 3	On bank of Leaf River; 1.2 in. per day, 1:1.
GrenadaLamar, HT16	31°11′ 33°50′ 31°09′	89°11′ 89°47′ 89°33′	Hattiesburg Formation Tallahatta Formation Pascagoula and Hatties-	416/392 282/227 889/838	127. 89 (03) 8. 66 101. 50	. 16 . 20 . 5	. 13 . 20 . 5	. 29 . 40	Riovi syncline, Tatum sal
HT2A HT5	31°07′ 31°08′	89°34′ 89°34′	burg Formations. Hattiesburg Formation Catahoula(?) Sandstone	1, 080/935 680/579	132. 53 104. 24	. 5	. 5	1	dome. Do. On top of Tatum salt
E7	31°08′	89°34′	Limestone caprock of salt dome.	1, 386/945	94. 10(+. 25)	2. 25	. 09	2. 34	dome. Do.
E9 Lowndes	31°08′ 33°23′	89°34′ 88°26′	Sand of Gordo Formation	1, 450 500/400	84.96(+.1) .85(06)	. 60 1. 15	. 70 1. 15	1. 30 2. 30	Do.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	evel fluctua	tion (feet)	
County, well	Lati- tude, N.	Longi- tude, W.	Water-bearing formation	of well; second, depth of casing to	Depth to water (feet)		earthquake evel		Remarks
				screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
			Mis	sissippi—Con	tinued				1.
Rankin Washington	32°18′ 33°02′	89°47′ 90°59′	Cockfield Formation Alluvium of Mississippi Valley.	594/565 105/80	104. 07 11. 59	. 20	. 20	. 40	
		[]	Pata furnished by the Missou	Missouri ri Geological		Resources]		
Barton, 32N/30W-30ed	37°29′	94°16′	Dolomite	971/553	223.83(+.4)	>2.85	>2.15	>5	Aftershocks recorded (see table 6); 1.2 in. per day,
Bollinger, 28N/9E–32dca Butler, 26N/5E–34ca Callaway, 44N/11W-15bab Cape Girardeau, 29N/12E–	37°03′ 36°52′ 38°36′ 37°11′	90°05′ 90°31′ 92°10′ 89°45′	Quaternary alluvium Gasconade Dolomite Quaternary sand Quaternary alluvium	115/70 631 99/91 75/70	7. 15 139. 27 (+. 65) 32. 07 18. 1	. 02 2. 45 . 1 . 02	. 04 >. 23	. 06 > 2. 68 . 1 . 22	1:6. 1.2 in. per day, 1:6.
8db. Dunklin, 22N/10E-34cdc Franklin, 42N/1W-26dd 44N/1W-27cbbc	36°30′ 38°21′ 38°32′	89°58′ 90°59′ 91°01′	Wilcox(?) Group Dolomitedo	130/104 255 1, 360	14. 9 67. 8 75. 0	. 55 . 05 >1. 0	>1.1 .12 2.9	>1.65 .17 >3.9	Aftershocks recorded (see
Greene, 29N/22W-13bcc	37°13′	93°17′	do	1, 346	393. 9	>9.9	>.1	>10	table 6); 1.2 in. per day, 1.6 Between Mar. 28 and June 2, water level rose 50 ft;
Howell, 24N/8W-21ca	36°44′ 36°56′ 36°45′	91°51′ 92°03′ 91°45′	do do	1, 305/800 780/650	308, 45 212, 5 148, 4(+, 24)	6. 45 1. 55 6. 5	>1.55 .5 >1.6	>8 2.05 >8.1	1.2 in. per day, 1:12. 1.2 in. per day, 1:6. 1.2 in. per day, 1:12.
Jasper, 27N/32W-1bac McDonald, 21N/33W-22aa 23N/30W-18aad Madison, 33N/7E-20bed	36°32′ 36°40′ 37°32′	94°29′ 94°14′ 90°18′	Dolomite Limestone Dolomite, sandstone, and arkose,	1, 747/375 850/99 346/44 590/187	42. 7 86. 5 123. 35 103. 55	1. 95 2. 2 . 1	. 85 2. 9 . 05 1. 10	2.80 5.1 .15 >1.10	Water level rose 5.55 ft in 4 hrs after quake and 1.65
Marion, 58N/5W-10ab Mississippi, 25N/16E-29ccb Pemiscot, 17N/11E-36ab Perry, 34N/8E-34c Phelps, 34N/9W-18 Polk, 33N/21W-5adc	39°51′ 36°47′ 36°04′ 37°36′ 37°39′ 37°37′	91°26′ 89°21′ 89°49′ 90°08′ 91°58′ 93°15′	Alluviumdodo	129/81 130/113 195/126 	23. 98 8. 15 15. 9 177. 3(+. 60) 189. 0(3) 67. 66(-1.1)	. 05 . 06 . 15 1. 65 2. 3	. 12 . 06 . 15 1. 05 >1. 5	. 17 . 12 . 30 2. 70 >3. 8	ft more in next 98 hrs. 1.2 in. per day, 1:6. Water level fell 1.1 ft in 1½
Ripley, 22N/4E-3dd_ St. Clair, 38N/26W-22adc St. Louis, 44N/3E-34cdba 47N/8E-18 Scott, 26N/14E-21bab Shannon		90°37′ 93°46′ 90°40′ 90°09′ 89°33′	Alluvium and dolomite Dolomite St. Peter(?) Sandstone Quaternary alluvium do Upper aquifer Lower aquifer	65/61 875/20 150	13. 65 109. 4 77. 55 35. 18 8. 25 (-, 55)	. 04 . 15 . 05 . 05 . 25	. 04 . 15 . 05 . 10 . 25	. 08 . 30 . 10 . 15 . 50 > 5, 00	hrs after quake.
Taney, 24N/18W-13d Texas, 30N/11W-17dda	36°45′ 37°18′	92°53′ 92°10′	Lower aquifer Dolomitedo	598/206 481/50	(55) 250.67(+1.17) 272.57	1. 17 . 13	. 1	2. 95 1. 27 . 26	
				Montana					
Gallatin, A1-4-25de	45°48′	111°10′	Alluvium	101	16.49(+.02)			>1	Aftershocks recorded (see table 6); 0.3 in. per day,
Flathead, 29–20–29bd Missoula, 13–19–8cb	48°14′ 46°54′	114°11′ 114°03′	Valley filldo	152 112	26, 52 51, 47	0. 385 1. 43	0. 22 1. 45	0. 605 2. 88	1:1. 0.3 in. per day, 1:1. No aftershocks recorded; 2.4 in. per day, 1:6.
				Nebraska					
Adams, 7-10-23al)	40°34′	98°24′	Pleistocene sand and	155	109	0. 10	0. 10	0. 20	
Hamilton, 10–6–26bc Kearney, 5–15–3ba Lancaster, A10–6–36cdd _ Merrick, 12–8–36bc	40°48′ 40°26′ 40°47′ 40°58′	97°58′ 99°00′ 96°41′ 98°11′	graveldodo Dakota Sandstone Pleistocene sand and	130 122 170 8	87. 5(05) 96 62. 60 3. 45	. 13 . 09 2. 05 . 025	. 13 . 03 2. 05 . 025	. 26 . 12 4. 10 . 05	No aftershocks recorded.
Polk, 14-2-21db	41°10′	97°33′	gravel.	180	81. 45				Water level declined 0.20 ft 8 hrs after quake.
Saline, A8-3-19ad	40°38′ 40°20′	97°07′ 97°26′	Pleistocene sand and gravel.	151 95	97. 37(-0. 07) 89	. 06	. 14	. 20	No fall below prequake
York, 9-4-6dd	40°53′	97°35′	do	102	86	1. 18	1. 16	2, 34	level.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First number, depth of well; second, depth of casing to screen, perforated casing, or open hole (feet)	Depth to water (feet)	Water-le	vel fluctuat	Remarks	
	Lati- tude, N.	Longi- tude, W.				From preearthquake level			
						Upward	Down- ward	Double amplitude	
				Nevada					
Clark, S 19/53-32aaa1 S 19/60-9bcc1	36°15′ 36°19′	116°02′ 115°19′	Alluviumdo	300 830/140	27. 11(+. 004) 106. 02(14)	0. 015	0. 01	0. 025 >1	0.3 in. per day, 1:1. Aftershocks recorded (see table 6); 1.2 in. per day,
S 21/54-10aac1	36°21′	115°53′	do	800/472	71.69(02)	. 51	1. 22	1. 73	1:1. Pen moved faster than in could flow. 0.3 in. per
S 21/54-28bd1 S 22/61-4bcc1	36°06′ 36°04′	115°55′ 115°10′	do	140 355	22.72 109.85(-, 28)	. 03	. 03	. 06	day, 1:2. 0.3 in. per day, 1:1. 1.2 in. per day, 1:1.
				New Hamps	hire				
			No we	ells recorded	the quake.				
				New Jerse	еу				
Atlantic, Pleasantville	39°24′	74°30′	Kirkwood Formation (800-ft sand).	680	(-)34.39	0.06	0. 03	0. 09	
Amatol	39°35′ 39°40′ 39°18′ 39°27′ 39°55′	74°41′ 74°40′ 74°37′ 74°27′ 74°50′	Cohansey(?) Sand Cohansey Sand Kirkwood Formation do Englishtown Formation	680/670 570/560 265/253	$\begin{array}{c} 5.11\\ (+)92.95\\ (-)28.74(04)\\ (-)21.11(03)\\ (+)45.84(01)\\ (-)14.90(02)\\ (+)128.85(03)\\ (+)114.92(02) \end{array}$. 03 . 09 . 24 . 01 . 10	. 03 . 09 . 09 . 12 . 13	. 06 . 18 . 33 . 13 . 23	
Lebanon 18-V _ Sawmill 1 Sawmill 2 Camden, Egbert	39°55′ 39°54′ 39°52′ 39°52′ 39°52′	74°50′ 74°28′ 74°31′ 74°31′ 75°04′	Raritan Formation Cohansey Sand do do Raritan and Magothy Formations	410/400 99 79 81/76 454	(-) 14. 90 (02) (+) 128. 85 (03) (+) 114. 92 (02) 10. 80 (-) 39. 63	. 08 . 06 . 01 . 03 . 04	. 10 . 12 . 05 . 04 . 13	. 18 . 18 . 06 . 07 . 17	
Elm Tree 3 Esterbrook Oaklyn	39°49′ 39°56′ 39°53′	74°56′ 75°07′ 75°04′	Formations. Englishtown Formation - Raritan Formation - Raritan and Magothy	717/706 300 112	(-) 25. 92(07) (-) 5. 04(18) (-) 34. 20	. 01 . 26 . 11	. 14 . 53 . 11	. 15 . 79 . 22	
N.Y. Ship New Brooklyn 1 New Brooklyn 2	39°54′ 39°42′ 39°42′	75°07′ 74°56′ 74°56′	Formations. Raritan FormationdoRaritan and Magothy	104 1, 495 848	(-)21.47(+.02) (-)12.33(04) (-)24.23(01)	. 10 . 06 . 06	. 06 . 11 . 01	. 16 . 17 . 07	
Cape May, Canal	38°57′ 39°06′ 38°57′ 40°44′	74°55′ 74°48′ 74°57′ 74°20′	Formations. Cohansey Sand Cohansey Sand do Wisconsin terminal	252/242 232/217 252/242 64	(-)13.39 (+)6.25(01) (-)11.80 (+)131.31(10)	. 06 . 08 . 20 . 00	. 05 . 12 . 21 . 10	. 11 . 20 . 41 . 10	
BallantineGloucester, Shell 5	40°43′ 39°49′	74°08′ 75°13′	moraine. Brunswick ShaleRaritan and Magothy Formations.	875 327	(-)80.14 31.63	. 13 . 13	. 07 . 13	. 20 . 26	
Shell 7 Texaco 3 Hercules (Gibbstown)	39°49′ 39°52′ 39°49′	75°13′ 75°09′ 75°16′	do Raritan Formation	322 298/225 100	31. 18 -)48. 42 (-)2. 68	. 14 . 22 : 07	. 09 . 14 . 06	. 23 . 36 . 13	
Middlesex, Forsgate 3	40°20′	74°27′	Raritan and Magothy Formations.	138/128	(+)70.65	. 05	. 08	. 13	
Duhernal 1	40°24′	74°21′	Old Bridge Sand Member of Raritan Formation.	67	(+)4.57	. 015	. 005	. 02	
Morris, International Pipe Randolph Township Whippany Madison 4. Deean, Colliers Mills 1. Colliers Mills 3. Garden State 2. Union, Hillside 4.	40°52′ 39°40′ 40°49′ 40°45′ 40°04′ 40°04′ 39°47′ 40°41′	74°26′ 74°33′ 74°23′ 74°23′ 74°27′ 74°27′ 74°14′ 74°13′	Wisconsin drift. Byram Granite Gneiss. Wisconsin glacial outwash. Wisconsin drift. Englishtown Formation. Mount Laurel Sand. Kirkwood Formation Brunswick Shale.	155 218 170 100 427/417 267/257 317 400	(+)295. 42(+. 04) 5. 32 (+)174. 18 (+)173. 45(+. 10) 54. 50 18. 68(01) (+)37. 18 (+)24. 40	1. 30 . 18 . 11 . 24 . 02 . 01 . 02 2. 14	1. 57 . 23 . 14 . 33 . 03 . 04 . 00 2. 23	2. 87 . 41 . 25 . 57 . 05 . 05 . 02 4. 37	Only one distinct after-
White 2 White 4 County Park Hatfield	40°40′ 40°40′ 40°41′ 40°37′	74°16′ 74°16′ 74°17′ 74°16′	do	250 350 290 ?	(+)56.16 (+)51.18(+.05) (+)59.03 (+)17.37(+.08)	1. 13 . 15 . 24 . 40	1. 21 . 10 . 22 . 49	2. 34 . 25 . 46 . 89	shock recorded (see table 6); 1.2 in. per day 1:5.
			1	New Mexic	0		l .		1
Chaves, Berrendo	33°27′	104°31′	San Andres Limestone	258	56. 15	0. 95	0. 99	1. 94	2.4 in. per day, 1:6.
(10.24.9.333) Berrendo-Smith (10.24.21.212)	33°26′	104°31′	do	324	53. 55	>3.5	>1.5	>5	Pen moved faster than ink could flow. No aftershocks recorded;
11.23.3.342 Eddy, 18.26.6.442	33°22′ 32°46′	104°36′ 104°24′	do	595 1, 008	188 148. 03	. 38	>. 14 >2	. 76E	2.4 in. per day, 1:6. 2.4 in. per day, 1:1.2. Aftershocks(?).

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

Eddy, 22.26.36.111a	2°21' 2°26' 2°45' 2°49' 4°15'	Longitude, W. 104°15′ 104°14′ 108°22′ 103°37′ 103°25′ 103°19′ 107°15′	Alluvium Capitan Limestone Conglomerate Ogallala Formation do Valley fill	of well; second, depth of casing to screen, perforated casing, or open hole (feet) Mexico—Co 260 327 580 252 97 43	Depth to water (feet) ntinued ?(+. 07?) 21. 61 158. 87 58. 13 18. 52		Downward 0.048 or .07	Double amplitude 0.10(+?)	Entire fluctuation was above prequake water level; 1.2 in. per day, 1:1 Chart pulled from drum by pen. 1.2 in. per day
21.26.36.221	2°26′ 2°45′ 2°49′ 2°57′ 4°15′	104°14′ 108°22′ 103°37′ 103°22′ 103°25′ 103°19′	Alluvium Capitan Limestone Conglomerate Ogallala Formation do Valley fill	forated casing, or open hole (feet) Mexico—Co 260 327 580 252 97	?(+. 07?) 21. 61 158. 87 58. 13	0, 05	0. 048 or . 07	0. 10(+?)	Entire fluctuation was above prequake water level; 1.2 in. per day, 1:1 Chart pulled from drum by pen. 1.2 in. per day
21.26.36.221	2°26′ 2°45′ 2°49′ 2°57′ 4°15′	104°14′ 108°22′ 103°37′ 103°22′ 103°25′ 103°19′	Alluvium Capitan Limestone Conglomerate Ogallala Formation do Valley fill	260 327 580 252 97	?(+. 07?) 21. 61 158. 87 58. 13	. 02	. 07		above prequake water level; 1.2 in. per day, 1:1 Chart pulled from drum by pen. 1.2 in. per day
21.26.36.221	2°26′ 2°45′ 2°49′ 2°57′ 4°15′	104°14′ 108°22′ 103°37′ 103°22′ 103°25′ 103°19′	Capitan Limestone Conglomerate Ogallala Formation Tubel Capitan Limestone Valley fill	327 580 252 97	21. 61 158. 87 58. 13	. 02	. 07		above prequake water level; 1.2 in. per day, 1:1 Chart pulled from drum by pen. 1.2 in. per day
Grant, 18.15.11.323 32 Lea, 17.33.13.341 32 16.36.5 Lotz 32 Roosevelt, 1N.33.36.400C 34	2°45′ 2°49′ 2°57′ 4°15′	108°22′ 103°37′ 103°22′ 103°25′	ConglomerateOgallala Formation doValley fill	580 252 97	158. 87 58. 13			>1	level; 1.2 in. per day, 1:1 Chart pulled from drum by pen. 1.2 in. per day,
Lea, 17.33.13.341	2°49′ 2°57′ 4°15′ 4°12′	103°37′ 103°22′ 103°25′ 103°19′	Valley fill	252 97	58. 13		00		1:1.
Roosevelt, 1N.33.36.400C 34	4°15′ 4°12′	103°25′ 103°19′					. 04	. 04 . 073	2.4 in. per day, 1:12. A water-table well; 1.2 in. per day, 1:1.
	4°12′ 3°07′		do		18. 52	. 005	. 010	. 015	1.8 in. per day, 1:2.4. A water-table well in which water level rose 0.03 ft then declined to normal over 4 hrs time:
_ 10			Magdalena Group	101 105	73.0(+.01) .10(38)	. 02	. 02	>1.04	1.8 in. per day, 1:2.4. 1.8 in. per day, 1:2.4. Aftershocks recorded (see table 6); 1.8 in. per day, 1:10.
				New York			1	-	
Chautauqua, Cu-10(208- 912-16) 42	2°08′	79°12′	Sand and gravel	232	30.06	1. 20	0. 90	2. 10	One aftershock recorded (see table 6); 1.2 in. per
Erie, 255-812-2	2°	78°	Glacial sand and gravel	81/81	4. 93(03)	. 00	. 04	. 04	day, 1:10. Water level rose 0.23 ft in 22 hrs after quake; 0.3
Genesee, 259-809-3 43 Nassau, N-7161 40 N-386740	3° 0°39′ 0°3 0 ′	78° 73°39′ 73°43′	Magothy Formation	54/51 671/661	21. 04 (+)5. 75	. 30	. 38	. 68	in. per day, 1:1. 0.3 in. per day, 1:1.
Niagara, 306–902–1 43 Onondaga, 253–614–1 42	3°06′ 2°53′	79°02′ 76°14′	Lockport Dolomite Hamilton Group Coarse sand and gravel Beekmantown Dolomite	517/506 36 160/43	(+)3.35 18.35 41.5	. 48 . 02 ?	. 55 . 02 ?	1. 03 . 04 1. 70	0.3 in. per day, 1:5.
Rensselaer, 235–342–10	2°35′ 4°52′	73°42′ 74°59′	Coarse sand and gravel Beekmantown Dolomite	96 180/54	27. 79 13. 12	. 21	. 22	. 43	0.3 in. per day, 1:1. Pen quit recording during quake; 0.3 in. per day, 1:1.
				North Caroli	na		-		
Chowan, CHO-2	66°14′	76°39′		320	9. 21	0. 17	0. 00	0. 17	2.4 in. per day, 1:6. Water level rose quickly 0.17 ft then declined to
	4°00′ 4°45′	77°55′ 77°25′	Castle Hayne Limestone	158 240	17. 30 9. 07	. 92 1. 00	. 93 . 78	1. 85 1. 78	prequake level in 20 min 2.4 in. per day, 1:6. Do.
				North Dako	ta				
Burleigh, 138–77–22aad			Glaciofluvial sand and gravel.	126/118	12			1.9	? in. per day, 1:12.
138-8-15cdd Ward Test Hole 2216, 155-82-19dbd.	8°	101°30′	do	168/140 107	36 40, 80	0.60	0. 29	. 32	? in. per day, 1:6. 2.4 in. per day, 1:6.
				Ohio					
	0°32′ 0°02′	84°23′ 80°44′	GravelAlluvial sand and gravel	100 59	6. 68(08) 24. 37	0. 08	0. 13 . 13	0. 21 . 27	Coda lasted 40 min. First detectable motion 15 mir before L max; 9.6 in. per day, 1:1.
Champaign, Ch-2	0°37′ 0°06′ 9°58′	81°05′ 83°45′ 83°43′	Sandstone Gravel Glacial outwash gravel	60 29	23. 51(07) 19. 63	. 16	. 16	. 32 . 10 1. 20	0.3 in. per day, 1:10.
Cl-2	9°55′ 9°58′	83°51′ 83°48′	Gravel	57 74 75	4. 6 5. 80 21. 90(+. 06)	. 52 . 20 . 39	. 68 . 20 . 33	. 40 . 72	0.3 in. per day, 1:5.
Fulton, Fn-1 419	0°21′ 1°35′ 1°25′	83°04′ 84°00′ 81°22′	Columbus Limestone Gravel Sandstone of Cuyahoga Formation.	135 130 120	29. 92(+. 48) 61. 11 40. 28(+. 08)	. 79(?) . 08 . 50	. 77 . 04 . 20	1. 56? . 12 . 70	Do.
H-2 39°	9°11′ 9°17′ 9°13′	84°47′ 84°39′	Graveldo	124 89	25. 36 12. 50	. 06	. 06	.12	Do.
H-10 39°	9°12′ 9°35′	84°27′ 84°28′ 81°54′	do	168 170 43	104. 10 92. 35 3. 58	. 22 . 30 . 58	. 19 . 27 . 46	. 41 . 57 1. 04	Do. 0.3 in. per day, 1:10. 0.3 in. per day, 1:2.
Marion, Mn-1 40°	1°37′ 0°34′ 0°02′	83°36′ 83°23′ 84°12′	Formation. LimestonedoGravel.	250 100	95. 26 9. 05	. 10 . 25 . 13	. 10 . 28 . 12	. 20 . 53 . 25	0.3 in. per day, 1:5.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

			ongi- e, W.	First number, depth of well; second, depth of casing to screen, perforated casing, or open hole (feet)	Water-le	vel fluctuat			
County, well	Lati- tude, N.	Longi- tude, W.			Depth to water (feet)	From preearthquake level			Remarks
						Upward	Down- ward	Double amplitude	
				Ohio-Conti	nued				
Montgomery, Mt-6	39°45′	84°11′	Gravel	60	36. 05	0. 22	0. 23	0. 45	0.3 in. per day, 1:10.
Pickaway, Pk-2	39°40′ 39°42′ 41°10′	84 16' 82°57' 81°02'	Gravel Sandstone	220 87	18. 69 18. 2 25. 60	. 04	. 03	. 07	0.3 in. per day, 1:5.
Po-4	41°10′	81°06′ 83°09′	do	225	28.72	. 11	. 04	. 12	0.01
Ross, Ro-6. Seneca, Se-2.	39°15′ 41°08′	83°09′	Gravel Limestone	250	2.62 $20.3(15)$. 46	. 44	. 90	0.3 in. per day, 1:1. 0.3 in. per day, 1:5.
Stark, St-5a Trumbull, T-2 Tuscarawas, Tu-1	40°49′ 41°16′	81°20′ 80°51′	Gravel Sandstone	132 124	29. 25 48. 89(+. 09)	. 37	. 30	. 67 . 30	Do.
Tuscarawas, Tu-1 Tu-2	40°36′ 40°36′	81°32′ 81°32′	Gravel		10.50(05)	. 18	. 26	. 44	0.3 in. per day, 1:5.
Tu-3	40°32′	81°29′	do	63	48.89(+.09) 5.35	.30	.10 .30 .16 .28	. 26 . 58	0.3 in. per day, 1:10.
Van Wert, Vw-1	40°36′ 40°52′	81°32′ 84°33′	Limestone	43 340	7. 05 27. 10	3. 10	. 30 2. 70	. 60 5. 80	0.3 in. per day, 1:5. No aftershocks recorded; 0.3 in. per day, 1:10.
		'		Oklahoma	1	,			
Grady, 4N-8W-33	34°46′	98°03′	Rush Springs Sandstone	254	84.70			>1	No aftershocks recorded;
Pontotoc, 1N-6E-4	34°34′	96°40′	Arbuckle Limestone	1, 707	128.85			>1	0.3 in. per day, 1:1. Aftershocks recorded(?);
1N-5E-27			do						0.3 in. per day, 1:1. Beaded cable slipped at
Texas, 1N-12E-35	36°30′	101°44′	Ogallala Formation	386	192.37(13)			>1	0.3 in. per day, 1:1. Beaded cable slipped at time of quake. No aftershocks recorded;
Wagoner, 19N-16E-26	36°05′	95°34′	Alluvium	31	28, 40	0, 05	0, 05	.10	0.3 in. per day, 1:1. 0.3 in. per day, 1:1.
Washita, 10N-19W-10	35°21′	99°12′	Elk City Member of Quartermaster Forma- tion.	55	37. 17	. 18	. 18	.36	Do.
		1		Oregon					
Yamhill, 4W-24J1	45°12′	123°07′	Alluvium	114/94	5.44(+.01)	0. 045	0. 01	0. 055	Recorder tends to "hang up"; 2.4 in. per day, 1:
				Pennsylvai	ia				
Chester, Ch-152	40°08′	75°30′	Stockton Formation	750	LSD	Flow at	2. 18	2. 18	Flowing well; no after-
Cumberland, Cu-2	40°02′	77°18′	Ledger Dolomite	. 37	16. 30(06)	LSD.	. 08	. 08	shocks. 0.3 in. per day, 1:10.
Dauphin, 020-646-8	40°30′ 40°20′	76°46′ 76°46′	Martinsburg Shale	400 185	31, 51 (-, 37) 23, 96	. 10	. 71	. 81	Do. ¹ 0.3 in. per day, 1:1. ¹
020-646-10 020-646-2	40°20′ 40°21′	76°46′ 76°46′	Limestone	_ 225	18. 60 6. 29(17)	. 00	. 28	. 28 >1	Do. ¹ Do.
Franklin, Fr-2	39°59′	77°39′	Limestone Stones River Limestone	441/60	6. 29(17) 28. 0(?)	1. 10	.65(+?)	1.75(+?)	Tane came off nulley: 0.3
Fulton, Fu–1 Lackawanna, Dodge shaft Olyphant	40°03′ 41°23′ 41°27′	78°08′ 75°41′ 75°36′	Mauch Chunk Formation Coal minedo		3. 45 594. 78 716. 61(+. 03)	. 06 . 61 . 32	. 09 . 64 . 28	. 15 1. 25 . 60	in. per day, 1:10. 0.3 in. per day, 1:5. 0.3 in. per day, 1:10. Do.
shaft. Storrs 2 shaft Lancaster, Ln-32(Ln-242)	41°27′ 40°09′	75°38′ 76°33′	New Oxford Formation	300	604. 82 6. 14(?)	1. 30	. 88	2. 18 >2	0.3 in. per day, 1:5. Beaded cable came off pulley; 0.3 in. per day,
Luzerne, Lu-243	41°18′	76°15′	Catskill Formation	195	51, 70			>1	1:2. No aftershocks recorded;
Dennison St. Bore- hole.	41°19′	75°51′	Coal mine		512			>2	0.3 in. per day, 1:2. Aftershocks recorded (see table 6); 0.3 in. per day
Mercer, Mr-1364 Montgomery, Mg-225	41°22′ 40°08′	80°23′ 75°21′	Cussewago Formation Stockton Formation		5. 14 38. 20(+. 30)	. 13	. 12	. 25 , 308	1:2. Do. Quake recorded at bottom of "low" in water level
York, Yo-180	40°03′	76°45′	New Oxford Formation	490	21. 15(10)	. 61	. 97	1. 48	0.3 in. per day, 1:10. Quake recorded at botton of "low" in water level;
007-637-7	40°07′	76°37′	do	148	6. 33	. 49	. 96	1. 45	0.3 in. per day, 1:5. Quake recorded at bottom of "low" in water level
005-639-7	40°05′	76°39′	do	222	19. 30	. 24	. 53	. 77	0.3 in. per day, 1:2. Quake recorded at botton of "low" in water level; 0.3 in. per day, 1:1.

¹ Records atypical, but similar in all three wells.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

				First num- ber, depth		Water-le	evel fluctuat	ion (feet)	
County, well	Lati-	Longi-	Longitude, W.	of well; second, depth of casing to		From preearthquake level			Remarks
tadi	, and a second	tude, N. tude, W.		screen, per- forated casing, or open hole (feet)		Upward	Down- ward	Double amplitude	
				Puerto Rico					
Santa Isabel, Jauca 2	18°01′	66°22′	Tuffaceous clastics	300	82, 50(?)	1. 60	1.80	3. 40	No aftershocks recorded; beaded cable slipped on pulley. Well is in a graben of a fault zone;
Lajas, La Parguera Bayamón, Fort Buchanan_	17°58′ 18°24′	67°02′ 66°08′	LimestoneSand, limestone, and clay.	92 242	38. 30 37. 52	. 30	. 29	. 59 . 02	2.4 in. per day, 1:6. 2.4 in. per day, 1:1. Quake recorded at bottom of water-level "low"; 0.3 in per day 1:1.
Vega Alta, Sabana Hoyos	18°26′	66°20′		_ 94	+29. 22	. 08	. 06	. 14	0.3 in. per day, 1:1. Another quake(?) re- corded 60 hrs later; 0.3 in. per day, 1:1.
				Rhode Islar	nd				
			No wells	recorded the	earthquake.				
				South Carol	ina				
Beaufort, BFT-101BFT-304	32°10′ 32°08′	80°44′ 80°50′	Ocala(?) Limestonedo	649	15. 14(36) 7. 4(?)	4. 49	4. 49	>2 8. 98	Beaded cable thrown off pulley.
Florence, FLO-126 Jasper, J-46	32°18′	80°58′	Upper Cretaceous sand Ocala(?) Limestone		22. 83	1. 97	2.75	4.72	Aftershocks recorded (see table 6).
Lexington, LEX-79 Orangeburg, ORB-5	33°	81°	Upper Cretaceous sand Chlorite-hornblende schist.	280 1, 839				1. 17 . 90	
ORB-7Richland, RIC-200	33° 34°	81° 81°	Crystalline rocks	1, 969				. 30	
				South Dake	ota	,			
Beadle, Huron 2	44°	98°	Basal sand of glacial drift.	74	15. 27	0.06	0.06	0. 12	0.3 in. per day, 1:5.
113-63-2bbbb 111-63-15bc2 Lawrence, A-7-2-10badc	44°37′ 44°25′ 44°	98°22′ 98°23′ 103°	Glacial outwashdo Opeche and Minnelusa Formations.	155/71 52/2 1, 306/1, 266	28.52(05) 16.65(+.03) +121	. 05 . 053 12	. 07	. 12 . 053 23	0.3 in. per day, 1:2. 0.3 in. per day, 1:1. Pressure recorder; 51° per day, 1:558, no aftershocks recorded.
				Tennesse	e			1	
Campbell, Cb: 0-6	36°34′	84°07′	Rockcastle(?) Sandstone	620	74. 99	>1	>1	>2	0.3 in. per day, 1:1. Pen thrown off recorder by quake.
Crockett, Ck: B-5	35°42′ 36°04′	89°05′ 87°23′	Claiborne Group ("500"- foot sand). Fort Payne Chert	537 387	40. 16	. 27	. 34	. 61 1. 00	0.3 in. per day, 1:2. 0.3 in. per day, 1:10.
Dickson, Di: F-19 Fayette, Fa: W-1	35°22′	89°33′	Wilcox Group ("1400"- foot sand). Claiborne Group.	1,025	23. 00 (07) 73. 98	. 03	. 15	. 18	0.3 in. per day, 1:2.
Fa: W-2 Franklin, Fr: F-1 Humphreys, Hs: H-1	35°22′ 35°03′ 36°01′	89°33′ 86°16′ 87°57′	Fort Payne Chertdo	100	41. 36 30. 60 86. 70	>1 . 02 >1 . 105	. 04 . 105 >1	. 06 . 21 >2	Do. Do. 0.3 in. per day, 1:1. Pen thrown off recorder by quake.
Madison, Md: N-1- Shelby, Sh: J-1	35°42′ 35°00′	88°37′ 90°05′	Ripley FormationClaiborne Group		128. 6 42. 17	. 12 2. 86	. 14 1. 04	. 26 3. 90	0.3 in. per day, 1:1. One aftershock recorded (see table 6); 0.3 in. per day, 1:2.
Sh: K-75 Sh: L-1 Sh: L-15	35°05′ 35°03′ 35°04′	89°55′ 89°51′ 89°45′	Terrace deposit Claiborne Groupdo	578	44. 6 93. 58 74. 91(03)	. 12 . 005 . 145	. 09	. 21 . 005 . 185	0.3 in. per day, 1:2. 0.3 in. per day, 1:1. Do.
Sh: O-170 Sh: O-179	35°09′ 35°09′	90°01′ 90°02′	Wilcox Group Claiborne Group	1,387 472	73. 3 117. 5	. 05	. 05	. 10 . 80	0.3 in. per day, 1:10. Do.
Sh: Q-1	35°13′ 35°09′	89°54′ 89°48′	do	344 384	101. 62 90. 44	. 79	. 42	1.21 .147	0.3 in. per day, 1:2. 0.3 in. per day, 1:1. Wat level declined 0.13 ft in hrs after quake.
Sh: Q-24 Sh: Ŭ-1 Sh: U-2	35°13′ 35°21′ 35°21′	89°52′ 89°57′ 89°57′	Wilcox GroupClaiborne Group.	1,558	69. 36 53. 8 51. 1	. 10 . 47 . 96	. 04	. 14 . 96 >1. 00	0.3 in. per day, 1:2. 2.4 in. per day, 1:2.4. Do,
Tipton, Tp: E-3	35°26′ 35°55′	89°47′ 86°54′	Chandorne Group	496/466	197. 49 85. 37	. 24	. 30	. 54	1.2 in. per day, 1:2. 0.3 in. per day, 1:1. Wate level declined 0.27 ft in 16 hrs after quake.

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

			Longi- de, W.	First number, depth of well; second, depth of casing to screen, perforated casing, or open hole (feet)	Depth to water (feet)	Water-level fluctuation (fee			et)	
County, well Lat	Lati- tude, N.	Longi- tude, W.				From preearthquake level		Double	Remarks	
						Upward	Down- ward	amplitude		
				Texas						
Bexar, D-59	29°34′	98°41′	Edwards Limestone	400	279. 26	1. 31	1. 75	3.06	Float hung after major wave.	
F-172	29°32′	98°28′	do	500	124.10	>2.1	>2.9	>5	Aftershocks recorded (see table 6); 2.4 in. per day,	
F-214(169a) J-17	29°35′ 28°29′	98°30′ 98°26′	do	547/100 874	278. 45 76. 90	. 11 >2. 90	. 18 >2. 10	. 29 >5	1:6. Do.	
Comal, G-49 H-36	29°42′ 29°35′	98°08′ 98°19′	do	230/27 292/220		. 08 1. 14	. 05 1. 17	. 13 2. 31		
H-39 Dallas, C-19	29°37′ 32°59′	98°18′ 97°00′	Woodbine Sand	250 268/100	·- -	. 05	. 06	. 11		
Dallas, C-19 Cl Paso, Q-86(CR-1) Q-176(CR-2) Q-181	31°56′ 31°57′	106°36½′ 106°37′	Alluvium Bolson depositsdo	200/100 1,072/585 1,013/528	21. 54(10) 47. 27	. 11	. 21	. 32 1. 38	2.4 in. per day, 1:6. Do.	
Q-181 Q-182(CR-4) Q-203(CR-5)	31°58′ 31°56′	106°37′ 106°37′	Alluvium Bolson deposits	1,013/528 202/102 700/355	13. 39 31. 76	. 45 . 26 . 79	. 71 . 29 >. 25	1.16 .55 1.6E	Do.	
V-42 Falveston, E-93	31°47′ 29°23′	106°22′ 95°06′	do Sand	710/380	131.37	. 06	.00	.06	Do.	
L-61 Harris, W-109	29°22′ 29°54′	95°04′ 95°08′	Beaumont Clay	850 196	114.2 91.31(+.015)	. 82	. 38	1.20 .54	2.4 in. per day, 1:12. 0.3 in. per day, 1:1.	
895 Jackson, PP-80-03-101	29°42′ 28°59′	95°16′ 96°42′	Gulf Coast aquiferdo	1,651/1,027 590/150	59. 13(06)	. 40 2. 9	. 65 2. 9E	1.05 5.8E	Aftershocks recorded (see table 6); 2.4 in. per day, 1:6.	
PP-66-60-605 Medina, C-9-53	29°03′ 29°29′	96°31′ 99°08′	Edwards Limestone			. 21 . 52	. 17 . 36	. 38	1.0.	
I-2-25a J-1-82 Uvoldo H-2-23	29°27′ 29°21′ 29°27′	99°17′ 98°53′ 99°40′	do do Edwards Limestone	712		. 24	. 35	>5 >2 . 59		
Uvalde, H-2-23 H-2-30 H-3-23	29°22′ 29°26′	99°42′ 99°30′	dodo	721		. 25	.16	. 41 >5		
H-5-1 Val Verde, XV-1a	29°12′ 29°22′	99°46′ 100°48′	do	350 750	80. 2	. 55 . 65	. 50 . 65	1. 05 1. 30	2.4 in. per day, 1:6.	
				Utah						
Davis, (B-2-1) 24bad3	40°53′	111°54′	Alluvium	386	(+)29	1	1	2	Pressure recorder. 51° pe	
Juab, (D-11-1) 8aad1 Millard, (C-16-7) 12dcd	40°52′ 39°24′	111°59′ 112°35′	do	100	12.24(02) (+)21.2	. 27 1. 0	. 21 . 2	. 48 1. 2	day, 1:115. 0.3 in. per day, 1:1. Pressure recorder. 51° pe	
(C-16-8) 21 bcb1 (C-19-5) 4ddd1	39°25′ 39°11′	112°45′ 112°24′	do	988/118 521	3.97 32.74(06)	1.30 .04	1.20 .06	2. 50 . 10	day, 1:129. 2.4 in. per day, 1:6. Do.	
Salt Lake, (C-3-1) 32cad2 (C-4-1) 23dbd1	40°30′ 40°26′	111°58′ 111°54′	do	218 152	32. 74(06) 27. 06(06) 52. 87(?)	. 04	. 15 . 23	. 19 . 45	1.2 in. per day, 1:1. Float hung at time of quake; 0.3 in. per day, 1:1.	
Tooele, (C-2-6) 36dcc1 (C-3-2) 14bad1	40°35′ 40°33′	112°28′ 112°02′	do do	176 1,000	92. 36 324. 42	. 44 >. 58	. 58 >. 42	>1.02	0.3 in. per day, 1:2. Aftershocks recorded (see table 6); 2.4 in. per day,	
(C-7-8)10ebd1	40°13′	112°44′	Alluvium	175	90. 65	. 16	. 025	. 185	1:1.2. Float may not have moved freely; 2.4 in. per day, 1:1.2.	
(C-2-4)33aac1 Utah, (D-5-1)8dcc1	40°36′ 40°23′	112°17′ 111°51′	do	182 240	18. 50 14. 7(24)	1.48	1.66	>1 3.14	1.2 in. per day, 1:1. 1.2 in. per day, 1:6.	
Weber, (A-6-1)11cab1 (B-6-1)30cca1	41°16′ 41°13′	111°48′ 112°00′	do	354 756	14. 69(+. 01) 35. 17	. 04	. 024	>1 064	2.4 in. per day, 1:6. Aftershocks recorded (see table 6). Water level rose 0.08 ft in 8 hrs after quake; 1.2 in. per day, 1:1.	
				Vermont						
			No wells	recorded the	earthquake.					
				Virgin Islan	ds					
St. Thomas	18°21′	65°00′	Andesite volcanic breccia and tuff.	220		0. 02	0. 03	0. 05		
				Virginia		t .				
Page	38°	78°	Basalt of Catoctin Formation.	280	44.85(+0.50)	1.16	0. 45	1. 61	The aquifer of metamorphosed basalt is steeply dipping; 2.4 in. per day,	

Table 7.—Hydroseisms in wells in the United States caused by the Alaska earthquake—Continued

County, well t				First number, depth of well; second, depth of casing to screen, perforated casing, or open hole (feet)	Depth to water	Water-le	evel fluctuat	Remarks	
	Lati- tude, N.	Longi- tude, W.	Water-bearing formation			From preearthquake level			
						Upward	Down- ward	Double amplitude	
				Washingto	n				
Grant, 14/25–28E1 15/26–28Q1	46°40′ 46°45′	119°42′ 119°41′	Ringold Formation Yakima Basalt	648/492 892	472.83(03) 307.80(+.66)	0.05	0. 07 . 08	0.12	Water rose 1.20 ft within 4 hrs after first shock wave and stayed that way; 2.4 in. per day, 1:1.2.
Pierce, 20/3-18C1	47°10′	122°23′	Vashon outwash sand and gravel.	185/152	101.60	1. 97	1. 95	3. 92	Aftershocks recorded (see table 6); 1.2 in. per day 1:10.
Spokane, 25/43–13H126/45–32J2Thurston, 17/2–19M2Yakima, 12/17–8N1	47°40′ 47°44′ 46°56′ 46°32′	117°20′ 117°10′ 122°36′ 120°43′	Sand and graveldodo Basalt	71/65 155 97 212	66, 53 130, 42 23, 87 10, 79	. 01 . 57 . 03 . 12	. 01 . 65 . 03 . 22	. 02 1. 22 . 06 . 34	2.4 in. per day, 1:6.
'				West Virgini	ia				
Berkeley, 20-5-7	39°27′	77°58′	Beekmantown Limestone	250	38. 58	0.15	0. 15	0.30	1.2 in. per day, 1:5.
				Wisconsin	1				
Dane, Dn-9/11/34-4 Dn-8/6/26-11	43°12′ 43°08′	89°10′ 89°44′	St. Peter Sandstone Pleistocene sand and gravel.	70 59	51.11(+.01) 13.29	0. 09 . 17	0.00	. 09	0.3 in. per day, 1:1. Do.
Dodge, Dg-11/16/5-4	43°27′	88°37′	Cambrian and Ordovician sandstone	475	119.04(+.81)	. 81	. 19	1. 00	0.3 in. per day, 1:10.
Fond du Lac, Fl-15/17/11-12 Kenosha, Ke-2/20/18-19B Ke-1/22/13-46 Marinette, Mt-30/23/19-5	43°47′ 42°37′ 42°32′ 45°03′	88°25′ 88°11′ 87°50′ 87°44′	do Sand and gravel Dolomite Cambrian and Ordovician sandstone.	817 74 125 703	67. 97 (42) 1. 46 (+. 02) 21. 33 20. 35 (+3)	. 50 . 02 . 13 >3. 35	. 66 . 05 . 05 >. 17	1. 16 . 07 . 18 >3. 52	Do. Pen caught at edge of chart; 1.2 in. per day,
Milwaukee, Ml-7/22/29-45 Ml-7/22/17-120	43°02′ 43°04′	87°54′ 87°54′	Milwaukee Formation	1, 015 400	47.70 (+7.8)	. 43	. 34	.77 >12E.	1:10. 0.3 in. per day, 1:10. Aftershocks recorded (see table 6).
Ml-6/21/32-148_ Monroe, Mo-8/2W/29-17	42°56′ 44°00′	88°01′ 90°39′	Cambrian sandstone	179 192	35. 57 5. 26(+1. 43)	1. 43		>2 1.43	Permanent rise of 1.43 ft; 1.2 in. per day, 1:6.
Portage, Pt-23/8/13-410 Sauk, Sk-10/6/3-1 Sheboygan, Sb-15/21/28-19 Waukesha, Wk-6/19/2-14	44°28′ 43°22′ 43°44′ 43°00′	89°30′ 89°46′ 87°59′ 88°13′	Sand and gravel	90 426 450 1, 300	7. 52(01) 81. 78 3. 18 351. 45(+. 45)	. 09 . 10 1. 21 . 45	. 09 . 10 1. 29 . 00	. 18 . 20 2. 50 . 45	1.2 in. per day, 1:1. 1.2 in. per day, 1:6. 1.2 in. per day, 1:6. 1.2 in. per day, 1:10. 0.3 in. per day, 1:10.
Waupaca, Wp-21/11/9-63	44°18′	89°10′	Pleistocene sand and gravel.	94	21, 49(03)	. 37	. 35	. 72	1.2 in. per day, 1:6.
Wp-22/14/12-13	44°23′	88°44′	do	203	11. 25(01)	. 08	. 03	. 11	Do.
				Wyoming					,
Laramie, 14-67-18ddc	41°11′	104°56′	Siltstone of Brule Formation.	311	20. 26	0. 03	0. 03	0.06	
Platte, 29–69–24dbc2	42°28′	105°04′	Hartville Formation	840	15. 18	1.00	1E.	2E.	No aftershocks recorded; 2.4 in. per day, 1:6.